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1 **Very long-term stability of passive margin escarpment constrained by**
2 **^{40}Ar - ^{39}Ar dating of K-Mn oxides**

3

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23

24

25 **ABSTRACT**

26 The post-rift denudation history of high-elevation divergent continental margins is
27 central to decipher “source-to-sink” systems across such margins and to unravel the
28 topographic evolution of their escarpment. We perform ^{40}Ar - ^{39}Ar dating of supergene
29 cryptomelane (K-Mn oxide) from supergene manganese ore deposits to constrain the
30 age of in-situ formed laterites of the low-relief lowland and highland separated by the
31 Western Ghats Escarpment (WGE) of Peninsular India. Documentation of laterites as
32 old as 53 Ma on the highland and 47 Ma at the foot of the WGE shows that the
33 escarpment stabilized before 47 Ma (possibly 60 Ma ago). The geomorphic setting of
34 the dated weathering mantles further allows constraining post-Early Eocene denudation
35 of the lowland and the highland to maximums of 5 and 15 m/My, respectively. These
36 results allow refining apatite fission track thermochronology and cosmogenic
37 radionuclides studies that overestimate 4 to 20 times denudation rates, particularly after
38 escarpment stabilization.

39

40 **INTRODUCTION**

41 Constraining the denudation history of high elevation (i.e., escarpment bearing)
42 divergent continental margins documents their post-rift topographic evolution that has
43 implications to quantify their vertical motion, drainage pattern and sedimentary supply
44 over geological time scales. Current evolution models of high-elevation passive margin
45 topography predict contrasted denudation histories. Models based on apatite fission
46 track thermochronology (AFTT) predict a short period (≤ 10 -15 Ma) of high denudation
47 focused on a coastal strip resulting in escarpment formation shortly after the onset of
48 rifting (Brown et al., 2002; Braun and van der Beek, 2004; see also Matmon et al.,

49 2002). An alternative model integrating AFTT, the elevation of paleosurfaces and
50 indirect geological constraints argues for successive burials, uplifts and topographic
51 rejuvenations after continental break-up implying that escarpments result from the last
52 uplift episode (Japsen et al., 2012). However, possible thermal instabilities in the
53 shallow crust (< 2 km, i.e., $T^{\circ} < 60^{\circ}\text{C}$) and the failure of AFTT models to quantify
54 erosion of crustal sections < 1 km (Brown and Summerfield, 1997) result in large
55 uncertainties in the timing and magnitude of denudation predicted by these models (see
56 also Kohn et al., 2005). Whether and how escarpments are rejuvenated after their
57 formation (e.g., Moore and Blenkinsop, 2006) is also debated because the mode and
58 velocity of their retreat have crucial implications on the evolution of the drainage
59 networks of continental margins and their sedimentary budgets (e.g., Tucker and
60 Slingerland, 1994). Studies based on cosmogenic radionuclides (CRN) predict slow
61 denudation and retreat of escarpments, i.e., 3 to 16 m/My (Bierman and Caffee, 2001),
62 but are only valid over millennial time scales.

63 Once slow denudation regimes (< 20 m/My) install on continental margins,
64 remnant landscapes may be commonly preserved as a result of limited to moderate
65 relief inversion (Beauvais and Chardon, 2013). In the tropical belt, such paleolandscape
66 remnants are mantled by laterites that may host K-Mn oxydes (cryptomelane) datable
67 by $^{40}\text{Ar}/^{39}\text{Ar}$ geochronology that already proved beneficial (e.g., Vasconcelos, 1999;
68 Vasconcelos and Conroy, 2003; Beauvais et al., 2008). If properly mapped and dated,
69 lateritic paleolandscape remnants are very useful to calibrate erosion over geological
70 time scales (Beauvais and Chardon, 2013). Here we combine geomorphology with the
71 dating of cryptomelane from in-situ formed lateritic Mn ore deposits preserved at the
72 foot and above the Western Ghats Escarpment (WGE) of Peninsular India as a test of

73 post break-up denudation scenarios deduced from AFTT and CRN studies (Gunnell et
74 al., 2003; Mandal et al., 2015). Our results lead to very low denudation rates (< 5-15
75 m/My) since stabilization of the escarpment at least 50 Ma ago and argue for a great
76 stability of both the topography and relief of the margin since then.

77

78 **GEOMORPHOLOGICAL AND GEOCHRONOLOGICAL OUTLINES**

79 The WGE is up to 1000 m high and was carved into the Deccan traps and their
80 Archean basement (Fig. 1a). The escarpment separates the coastal lowland from the
81 highland plateau and coincides with the continental divide along most of its trace (Fig.
82 1a). The lowland consists of a weakly dissected lateritic pediment (Figs. 1b and 1c)
83 making the seaward piedmont of the WGE (Widdowson and Gunnell, 1999), on which
84 laterites have been loosely dated to Mid- and Late Tertiary by paleomagnetism
85 (Schmidt et al., 1983). The poorly defined highland laterites would have Late
86 Cretaceous-Eocene paleomagnetic ages (Schmidt et al., 1983). Those laterites are
87 mostly bauxites and mantle a low-relief relict landscape best preserved as 1000-1100 m
88 high mesas near the WGE (e.g., Krishna Rao et al. 1989). Within the highland plateau,
89 relicts of that paleolandscape are sparser (Fig. 1b). At Sandur, ^{40}Ar - ^{39}Ar dating of
90 supergene cryptomelane formed in Mn ore deposits carried by such an highland relict
91 (Figs. 1a and 1b) documents a period of intense lateritic weathering at c. 36-26 Ma
92 (Bonnet et al., 2014), after the formation of Eocene bauxites (Krishna Rao et al., 1989)
93 and before Neogene landscape dissection (Radhakrishna, 1993).

94 Inversion of apatite fission track data predicts higher denudation in the lowland
95 (up to 120 m/My) than in the highland (< 20 m/My) between 80 and 50 Ma, and low (<
96 20 m/My) denudation on either side of the escarpment after 40 Ma (Gunnell et al.,

97 2003). This suggests that the WGE formed at c. 50 Ma, after the extrusion of Deccan
98 traps (65 ± 2 Ma) and largely after rifting between India and Madagascar (88 ± 3 Ma).
99

100 **FIELD RELATIONSHIPS, MATERIAL AND METHOD**

101 The lowland pediment consists of a c. 40 km wide concave surface ranging in
102 elevation from 30 to 300 m, and incised by 25 to 100 m deep valleys (Figs. 1b-c). Field
103 observations indicate that the pediment has truncated a bauxitic profile before being in
104 turn weathered and cemented by a ferricrete, which caps a 25 to 60 m thick lateritic
105 weathering profile (Fig. 1c). The studied lowland Caurem and Naveli Mn ore deposits
106 are lentoid pockets within the pediment weathering profiles developed from Archean
107 manganiferous schists and phyllites. The two open cast pits are located at the foot of the
108 WGE at ~ 100 m and 140 m elevation, respectively (Fig. 1c), and four samples were
109 collected at altitudes of 68 m and 45 m in Caurem pit and 113 m and 85 m in Naveli pit
110 (see supplemental material). The highland Sandur Mn ore deposit formed upon Archean
111 manganiferous phyllites and is exposed on a relict lateritic paleosurface capped by a
112 ferricrete at 1012-1015 m elevation (Fig. 1b). Two samples were collected on benches
113 at altitudes of 890 to 975 m (see supplemental material). All samples are massive with
114 botryoidal or cavity filling structures enabling cryptomelane crystallization.

115 Thorough optical microscopy and X-ray micro-fluorescence analyses of 200-300
116 μm thin sections allowed separation of eight cryptomelane grains from 300-500 μm
117 symmetrical sections slabs using a binocular magnifier (see Bonnet et al., 2014). The
118 separated grains were characterized using X-ray diffraction, electron microprobe
119 analyses and scanning electron microscopy before irradiation. Cryptomelane ($\text{K}_x\text{Mn}^{\text{IV}}_{8-x}\text{Mn}^{\text{III}}_x\text{O}_{16}$)
120 crystallizes into a monoclinic prismatic system with a typical tunnel-type

121 crystal lattice framed by a double chain of MnO_6 octahedra and K^+ cations in the large
122 tunnel to insure the electronic neutrality of the lattice (Turner and Buseck, 1979). High
123 retentiveness of potassium (content up to 5.5 wt.% K) and radiogenic argon ($^{40}\text{Ar}^*$) in
124 the intra-crystalline tunnels warrants the suitability of cryptomelane for ^{40}Ar - ^{39}Ar
125 dating (see Vasconcelos, 1999).

126 Gas was extracted from the irradiated cryptomelane grains either from a step-wise
127 heating procedure in a double staudacher-type furnace, or from step incremental heating
128 of the grains with a CO_2 laser probe power. The gas fractions were then cleaned and
129 analyzed using a VG3600 or Argus IV mass spectrometer. The $^{40}\text{Ar}/^{39}\text{Ar}$ ages are
130 calculated from plateaus encompassing at least three consecutive ^{39}Ar release steps
131 comprising up to 50% of total $^{39}\text{Ar}_K$ released, and from best-fit inverse isochrones in
132 $^{36}\text{Ar}/^{40}\text{Ar}$ vs. $^{39}\text{Ar}/^{40}\text{Ar}$ correlation diagrams (see also supplemental material).

133

134 **RESULTS AND INTERPRETATIONS**

135 The ^{39}Ar release spectra of irradiated cryptomelane grains are stacked in the
136 Figures 2a and 2b. Each age spectrum allows calculation of a precise age from a plateau
137 encompassing heating steps overlapping at 2σ confidence level while degassing at least
138 5% of the total ^{39}Ar released from the highly retentive intra-crystalline sites of
139 cryptomelane structure. The calculated plateau ages also agree well with the isochron
140 ages (Table 1; see also data repositories DR1 to DR2 in supplemental material). The
141 ^{40}Ar - ^{39}Ar ages from the highland Sandur deposit document a weathering period from c.
142 53 to c. 50 Ma (Fig. 2a and 2c). These ages complement those previously obtained in
143 the same deposit from c. 36 to c. 26 Ma by Bonnet et al. (2014), which are also shown
144 in figures 2a and 2c. The old ages (53-50 Ma) in the highland are interpreted to reflect

145 (bauxitic) weathering of the low-relief relict landscape, whereas the 36-26 Ma
146 weathering period (Bonnet et al., 2014) is interpreted as that having led to the
147 geochemical reworking of the bauxites documented by Krishna Rao et al. (1989).
148 Lateritic profiles of the lowland pediment record weathering periods at c. 47-45 Ma, c.
149 24-19 Ma, and an episode at c. 9 Ma (Fig. 2b and 2c). The 47-45 Ma ages date the early
150 (bauxitic) weathering of the lowland, whereas the ages between 24 and 19 Ma
151 document renewed weathering and formation of the pediment's capping ferricrete, the 9
152 Ma age reflecting a discrete weathering pulse.

153

154 **DISCUSSION**

155 The preservation of in-situ formed laterites as old as 53-50 Ma above and 47-45
156 Ma below the WGE indicates that the current topographic envelope of the SW Indian
157 margin corresponds to a bauxitic paleosurface dating from the Early Eocene, which
158 already included the WGE with its present amplitude (Fig. 3). The 47-45 Ma old
159 laterites at 55 m depth underneath the lowland pediment belong to a bauxitic profile that
160 did not exceed 120 m in thickness (e.g., Bardossy and Aleva, 1990). Even if the lowland
161 bauxitic profile had been totally eroded (i.e., 120 m of stripping), the maximum
162 denudation rate of the piedmont between early (bauxitic) weathering (47 Ma) and
163 abandonment of the pediment (19 Ma) would be lower than 5 m/My. Given that
164 incision of the pediment does not exceed 100 m, the incision rate of the WGE'
165 piedmont is less than 6 m/My over the last 19 Ma. Anyhow, the remarkable
166 preservation of the pediment surface argues for negligible net denudation of the
167 piedmont's envelope after c. 20 Ma (Fig. 1c).

168 Different ages recorded at a same depth in the lowland weathering profiles
169 indicate that preserved old lateritic mantles of the escarpment piedmont have undergone
170 several weathering episodes under a slow mechanical denudation regime. Preservation
171 of 47 Ma old bauxitic weathering mantles at least 60 m thick thus attests that the current
172 escarpment established at least 47 Ma ago, and did not retreat since then. Therefore,
173 escarpment formation or retreat driven by Neogene rejuvenation of the lowland (e.g.,
174 Radhakrishna, 1993; Widdowson and Gunnell, 1999) is excluded.

175 Topographic inversion of the low-relief bauxitic paleolandscape throughout the
176 highland plateau rarely attains 450 m (e.g., Figs. 1b and 3). Therefore, the ages (53-50
177 Ma) obtained on laterites mantling this relict landscape imply a post-Early Eocene
178 incision rate of less than 9 m/My over the last 50 Ma. Given the very low preservation
179 rate of the old paleolandscape (Fig. 3), this incision rate approximates the denudation
180 rate (Beauvais and Chardon, 2013). Renewed weathering of the highland during the
181 Oligocene (i.e., 36-26 Ma; Bonnet et al., 2014) did not result in significant denudation
182 of the Eocene landscape (< 100 m at Sandur and Belgaum for instance; Fig. 1a; Krishna
183 Rao et al., 1989 and our own field observations). Therefore, erosion rates up to 15
184 m/My may be expected on the highland over the last 26 Ma.

185 Formation of 47 Ma-old lateritic bauxites at the foot of the WGE would be in
186 agreement with the last denudation pulse predicted at c. 50 Ma in the lowland (Fig. 2c)
187 from AFTT inversion model (Gunnell et al., 2003). But this also suggests that the
188 escarpment is even older as bauxitic profiles form slowly at 3 to 10 m/My in contexts of
189 tectonic quiescence and limited denudation (Boulangé et al., 1997). Therefore, if the
190 lowland bauxitic profile comprising 47 Ma-old laterite was 120 m thick, it would have

191 taken at least 12 My to develop under an escarpment that should have been stabilized at
192 least 60 Ma ago.

193 Our results imply denudation rates lower than 5 m/My in the lowland for the last
194 47 Ma, as opposed to 5-20 m/My (between 40 and 0 Ma) and even 20-70 m/My
195 (between 50 and 40 Ma) from the AFTT model (blue curve in Fig. 2c). Even erosion
196 rates of c. 10^{-2} to 10^{-1} m/ky (10-100 m/My) derived from CRN studies in riverine
197 sediments (Mandal et al., 2015) are also questionable given the remarkable preservation
198 of thick weathering profiles as old as 47 Ma on the WGE' piedmont. Likewise, the
199 preservation of c. 100 m thick highland Eocene lateritic profiles formed upon the latest
200 basaltic flow (63 Ma) of the Deccan traps (Widdowson and Gunnell, 1999) would be
201 unlikely under such erosion rates or those (15-25 m/My) derived from AFTT between
202 53 and 45 Ma (red curve in Fig. 2c). Furthermore, AFTT fails to precisely measure or
203 detect less than 1 km of denudation, which corresponds to the height of most passive
204 margin escarpments. In other words, our results suggest that AFTT- or CRN derived
205 denudation rates are largely overestimated and not realistic once the escarpment there.

206 Given the early installation of very slow denudation regimes across high-elevation
207 margins such as that of Peninsular India, we suggest that the combination of radiometric
208 dating ($^{40}\text{Ar}/^{39}\text{Ar}$ of K-Mn oxides, or $^4\text{He}/^3\text{He}$ of Fe-oxides) of lateritic landscape
209 remnants with apatite (U-Th)/He thermochronological models be a powerful tool to
210 refine the denudation history of divergent margins of the tropical belt.

211

212 CONCLUSION

213 ^{40}Ar - ^{39}Ar geochronology of supergene cryptomelane formed in situ in supergene
214 manganese ore deposits on either side of the Western Ghats escarpment indicate the

215 development and preservation of an Eocene weathering paleolandscape mantling the
216 escarpment, attesting of its stability since at least 47-50 Ma, and possibly since 60 Ma.
217 The ages obtained also constrain slow denudation rates of the lowland (< 5 m/My) and
218 the highland (< 15 m/My) of the escarpment after its stabilization. Our results allow
219 refining erosion rates estimated from AFTT and show that absolute dating of lateritic
220 paleolandscape markers may accurately constrain post break-up denudation history of
221 passive margins, with important implications on their sediment delivery histories on
222 geological time scales.

223

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228

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305

306 **Table caption**

307

308 **Table 1.** Plateau and isochrone ages obtained for the cryptomelane separated from
309 samples collected in highland Sandur and lowland Caurem (CAU) and Naveli (NAV)
310 Mn ore pits. Depth indicates the sampling level below the surface. MSWD = Mean

311 square weight deviation of isochron. ($^{40}\text{Ar}/^{36}\text{Ar}_{\text{atm}} = 298.56 \pm 0.31$ from Lee et al.,
312 2006).

313

314 **Figures caption**

315

316 **Figure 1.** (a) Topo-bathymetric setting of the Southwestern Indian margin. (b)
317 Topographic cross-section from Sandur to the escarpment edge in the highland and Goa
318 in the lowland (section trace on Fig. 1a). (c) Synthetic Cross-section of the lowland
319 piedmont passing through the two dated Mn ore pits.

320

321 **Figure 2.** Stacked ^{40}Ar - ^{39}Ar age spectra of cryptomelane from (a) the highland and (b)
322 the lowland Mn ore pits (located on Fig. 1). (c) Weathering periods derived from series
323 of ^{40}Ar - ^{39}Ar plateau ages including σ errors, with denudation rate curves derived from
324 inversion model of apatite fission track data (Gunnell et al., 2003), both for the highland
325 (HL) and lowland (LL). (The 36-26 Ma ages are from Bonnet et al., 2014).

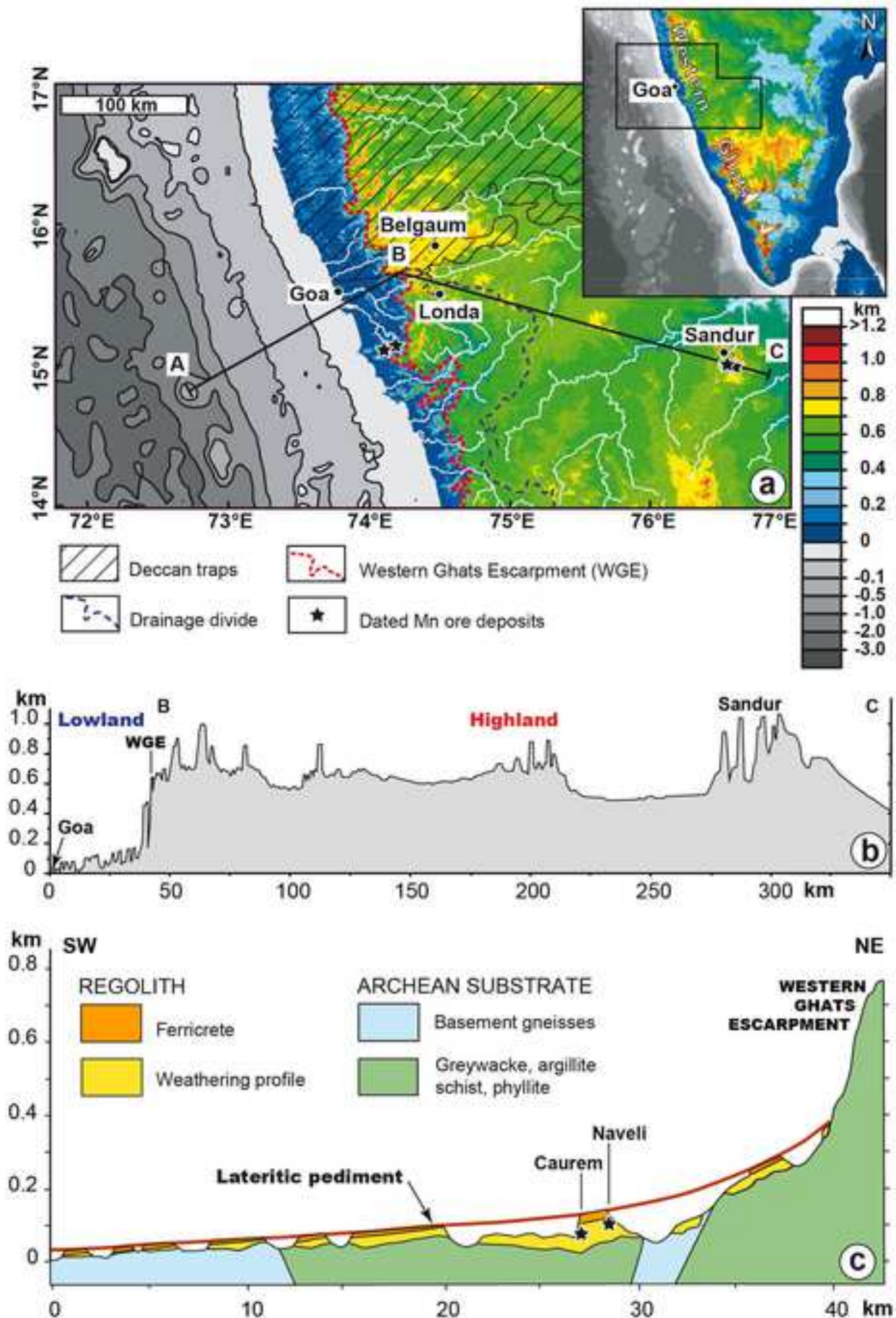
326

327 **Figure 3.** Cross-section of the Southwestern Indian divergent margin from offshore
328 basin (A) to Goa, escarpment edge (B) and Sandur (C) (section trace on Fig. 1a). The
329 offshore section including sedimentary limits is adapted from Chaubey et al. (2002),
330 and proportions of clastics are derived from Campanile et al. (2008). Very low offshore
331 accumulation of clastic sediments and correlative Eocene to Mid-Miocene carbonate
332 production (Whiting et al., 1994; Chaubey et al., 2002; Campanile et al., 2008) agrees
333 with onshore weathering and very slow mechanical denudation in the lowland.

Table 1

Sample ID	Depth ± 2 m	Plateau age $\pm 2\sigma$ (Ma)	% ^{39}Ar released	Isochron age $\pm 2\sigma$ (Ma)	$^{40}\text{Ar}/^{39}\text{Ar}$ intercept $\pm 2\sigma$	MSWD
Highland						
KMK-2	-37	49.60 \pm 1.34	100	49.60 \pm 1.41	302.0 \pm 25	0.10
KPA-8	-125	53.21 \pm 3.43	76	53.01 \pm 3.45	296.4 \pm 13	0.40
Lowland						
NAV-4	-27	20.84 \pm 1.68	54	21.12 \pm 0.45	288.0 \pm 6	0.33
CAU-1c	-32	21.92 \pm 0.23	56	22.07 \pm 0.24	294.0 \pm 1	2.28
CAU-1a	-32	23.67 \pm 0.29	57	23.67 \pm 0.29	294.9 \pm 2	2.31
NAV-3f	-55	8.65 \pm 0.16	84	8.57 \pm 0.59	308.7 \pm 94	0.55
CAU-3f	-55	19.35 \pm 1.52	58	20.22 \pm 0.35	283.9 \pm 18	0.05
CAU-3a	-55	23.65 \pm 0.21	54	23.71 \pm 0.21	295.5 \pm 2	1.85
NAV-3b	-55	45.08 \pm 0.78	57	45.15 \pm 0.74	293.7 \pm 3	2.04
NAV-3c	-55	47.07 \pm 0.54	60	47.11 \pm 0.53	293.5 \pm 3	1.25

TABLE 1

**FIGURE 1**

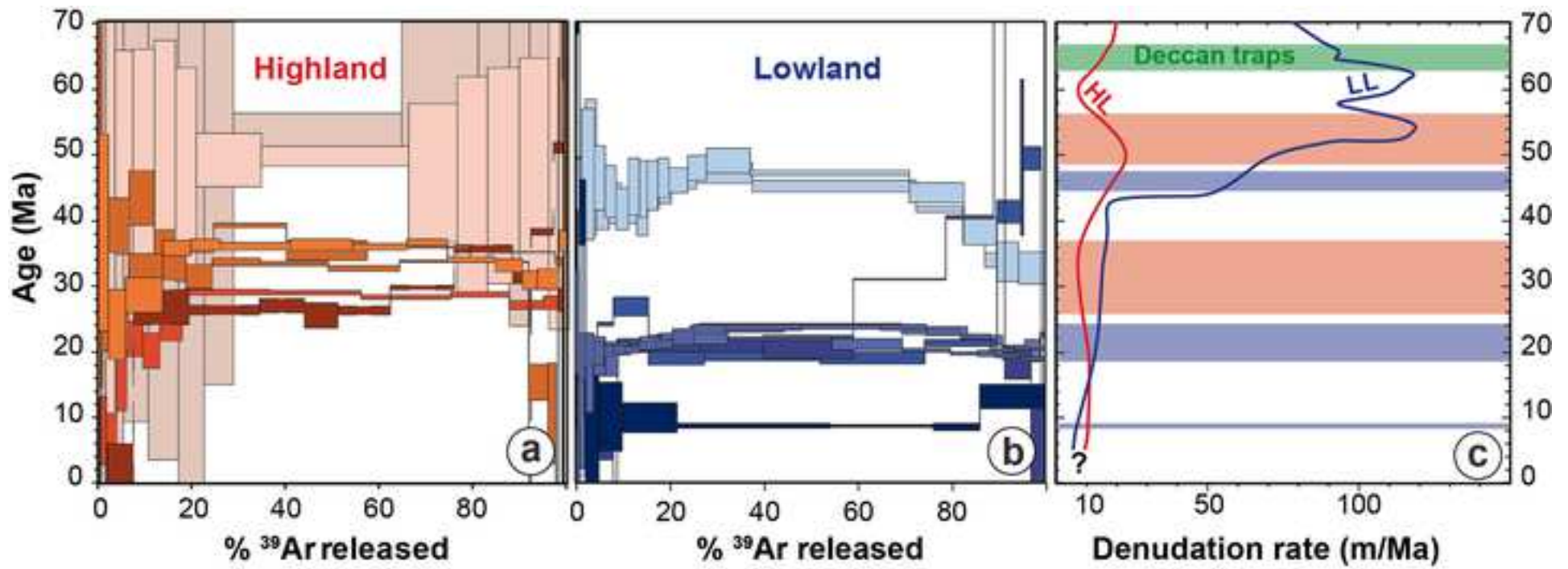
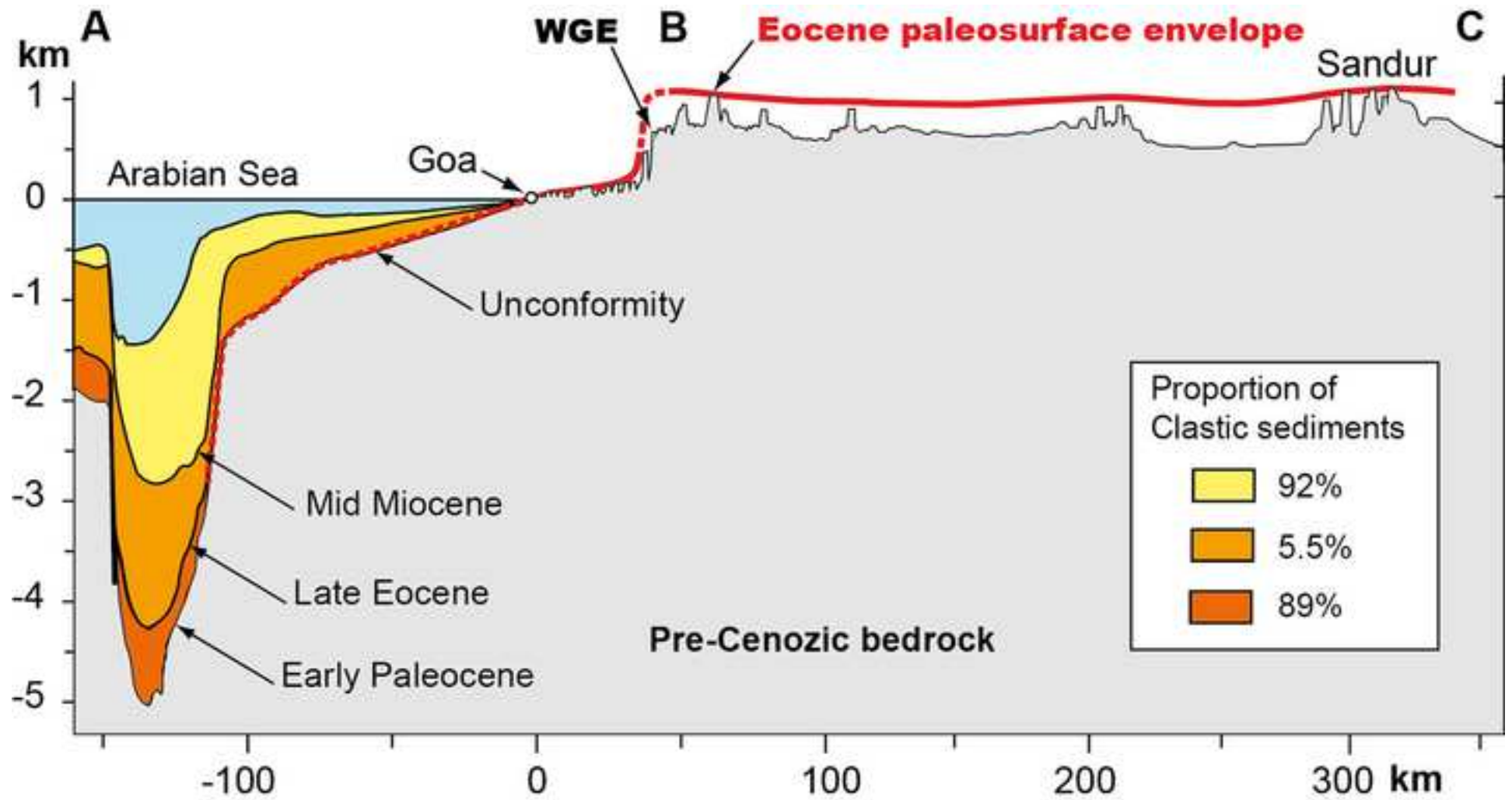


FIGURE. 2

**FIGURE 3**

Beauvais et al., 2015, Very long term stability of passive margin escarpment constrained by ^{40}Ar - ^{39}Ar dating of K-Mn oxides

Supplementary material

Manganiferous ore samples have been sampled in upland Sandur ore deposits (KPA-8: $15^{\circ} 0' 2.27''\text{N}/76^{\circ} 32' 41.6''\text{E}$; KMK-2: $14^{\circ} 59' 48.34''\text{N}/76^{\circ} 34' 40.34''\text{E}$) and in lowland ore deposits, Naveli (NAV-3: $15^{\circ} 07' 55.10''\text{N}/74^{\circ} 09' 49.72''\text{E}$; NAV-4: $15^{\circ} 08' 07.12''\text{N}/74^{\circ} 09' 32.26''\text{E}$), and Caurem (CAU-1: $15^{\circ} 07' 2.19''\text{N}/74^{\circ} 08' 39.66''\text{E}$; CAU-3: $15^{\circ} 07' 3.47''\text{N}/74^{\circ} 08' 38.97''\text{E}$).

Cryptomelane grains were separated from 300-500 μm thick slabs by hand picking. The separated grains were ultrasonically cleaned in absolute ethanol and conditioned in aluminium foil packets, to be irradiated for 50 hours in the TRIGA Mark-II reactor of Pavia University (Italia). The Factor J was determined from the analysis of the standard Taylor Creek Rhyolite sanidine-2 (TCRs-2) monitor, with an age of 28 ± 0.08 Ma (Baksi et al., 1996). The standard was analyzed after every ten unknown samples. After a two-month “cooling” period, the irradiated cryptomelane grains were loaded in a double vacuum Staudacher-type furnace for step heating Ar isotopes measurements. The furnace temperature was calibrated by means of a classical thermocouple, and the gas purification was accomplished using a cold trap with liquid air and Al-Zr AP10 getters (one hot, one cold) for 8 minutes before the introduction into the VG3600 mass spectrometer. One minute was allowed for equilibration before analysis. ^{40}Ar and ^{39}Ar were measured on a Faraday cup with a resistor of 10^{11} ohm, while ^{39}Ar , ^{38}Ar , ^{37}Ar , and ^{36}Ar were analyzed using a scintillator and photomultiplier after interaction on a Daly plate. The analytical data are reported in data repositories (Figures DR1 and Table DR2), and the errors are quoted at the 1σ level. Plateau ages are calculated from at least three consecutive ^{39}Ar release steps comprising up to 50% of total $^{39}\text{Ar}_K$ released and overlapping at the 2σ confidence level (Fleck et al., 1977). Isochrone

ages are accepted when mean square weighted deviation (MSWD) is less than 2.5 and the $^{40}\text{Ar}/^{36}\text{Ar}$ intercept within 2σ from the $(^{40}\text{Ar}/^{36}\text{Ar})_{\text{atm}}$ value of 298.56 ± 0.31 (Lee et al., 2006 ; Renne et al., 2009).

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Figure. DR1. ^{39}Ar releasing spectra showing well defined plateau ages with K/Ca (grey) and Ar^* (black) step curves (left) and inverse isochrone diagrams (right) of cryptomelane grains from the open pits samples of (A) Sandur (B) Naveli and (C) Caurem Mn ore deposits (MSWD = mean square weighted deviation).

Tables. DR2-DR3. Analytical results obtained for highland and lowland cryptomelane grains, either from Laser energy (spectrometer Argus IV) or double vacuum Staudacher-type furnace temperature, T °C, (spectrometer VG 3600) for each irradiated crypromelane grain. The

concentrations of ^{36}Ar , ^{37}Ar , ^{38}Ar , ^{39}Ar and ^{40}Ar with their respective 1σ error are provided for each step heating. The amount of $^{40}\text{Ar}^*$ (%), of ^{39}Ar released ($\%^{39}\text{Ar}$) and the K/Ca ratio (derived from $^{39}\text{Ar}/^{37}\text{Ar}$) are also given. Finally, this table show ratios $^{40}\text{Ar}^*/^{39}\text{Ar}_k$ used to determine the corresponding apparent ages, which are presented in the last column with their associated 2σ error. The different J-Factor values are also provided for each irradiated grains.

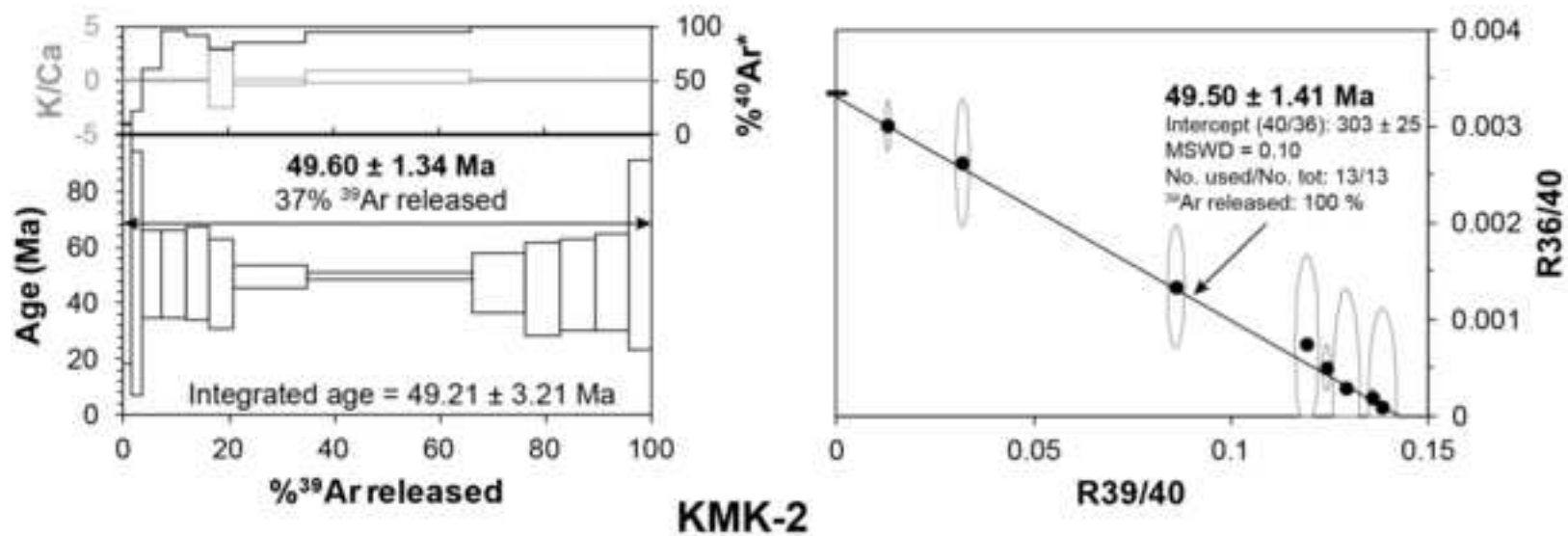
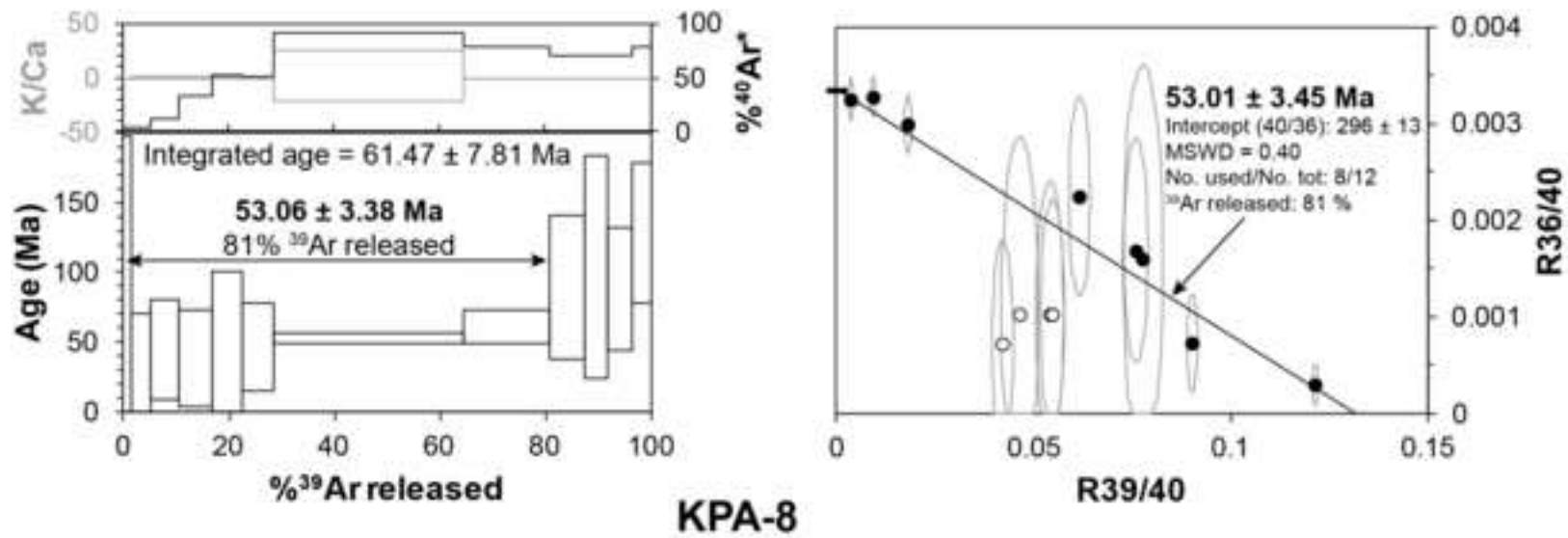


Figure. DR1- A

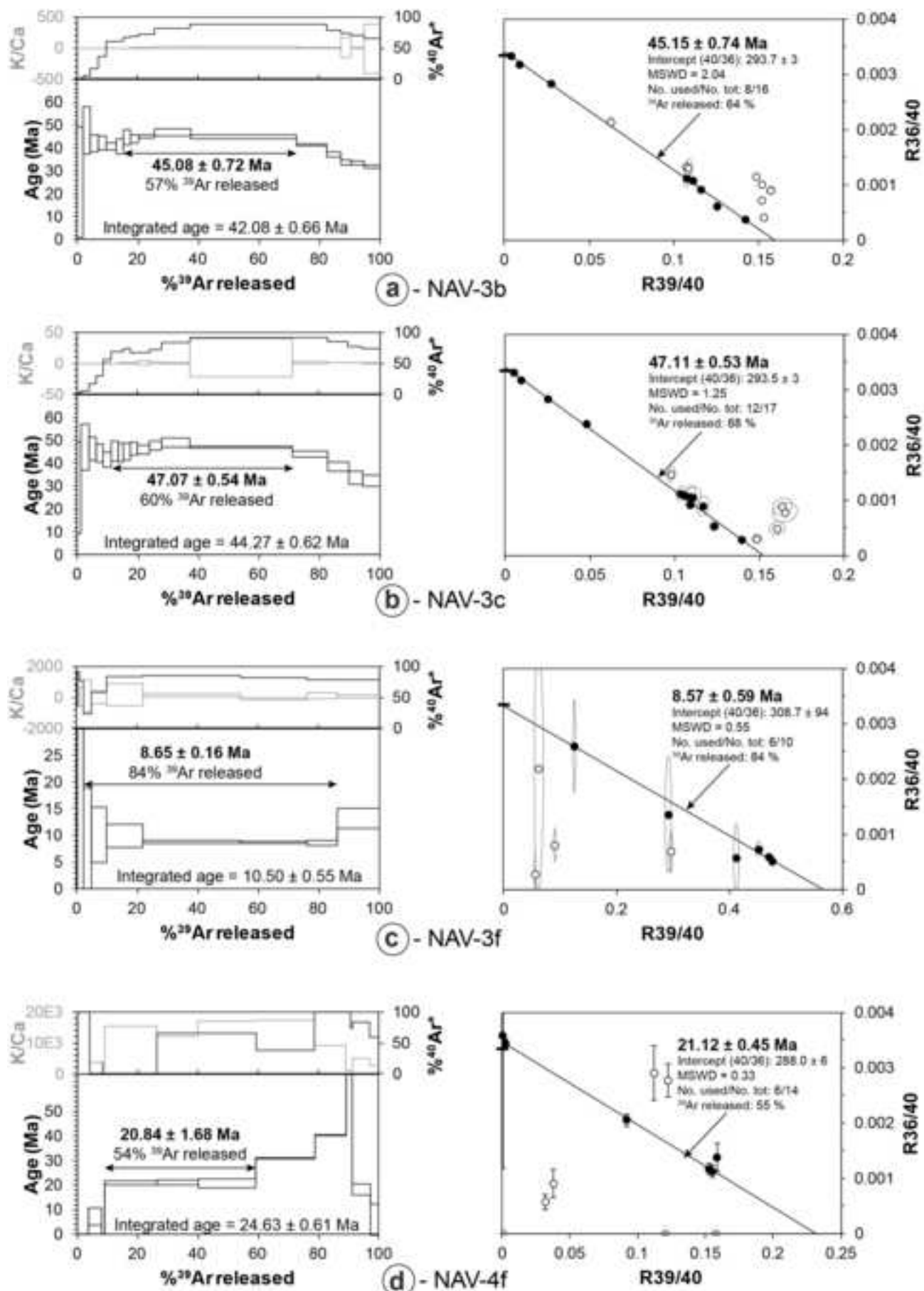


FIGURE. DR1-B

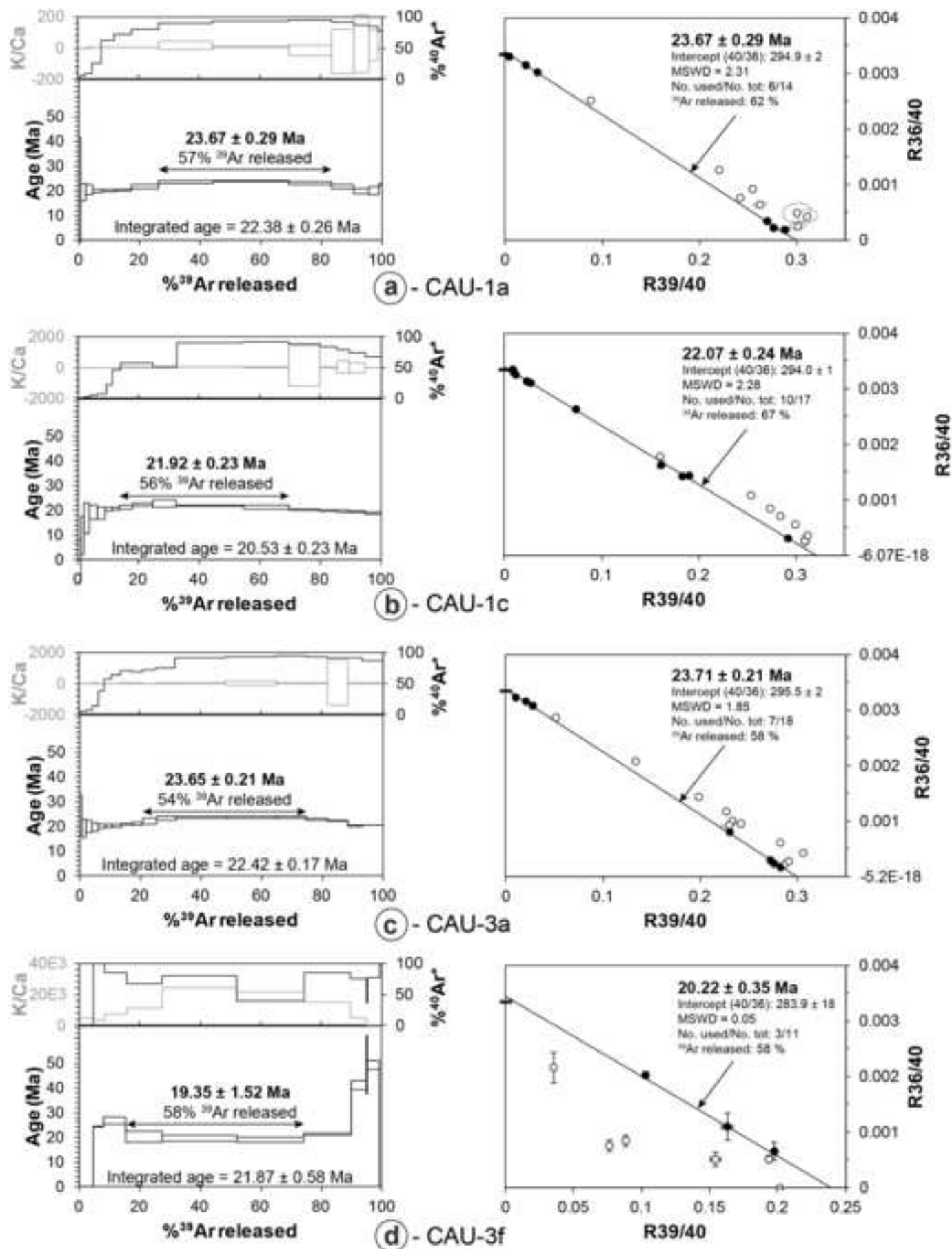


FIGURE. DR1-C

KMK-2		J = 0.00402950 ± 0.00000363															
Relative Abundances		36Ar [V]	%1σ	37Ar [V]	%1σ	38Ar [V]	%1σ	39Ar [V]	%1σ	40Ar [V]	%1σ	40(r)/39(k) ± 2σ	Age ± 2σ (Ma)	40Ar(r) (%)	39Ar(k) (%)	K/Ca ± 2σ	
48_KMK-2	4 °C	4	1.1839652	0.511	231.8581	72.426	0.3215167	6.386	4.95915	0.668	371.0028	0.252	8.52452 ± 6.07881	61.07 ± 42.82	11.01	1.46	0.0089 ± 0.0129
49_KMK-2	4 °C	4	0.5680986	0.691	134.9113	172.430	0.2579779	9.837	6.52910	0.550	201.6829	0.468	7.03545 ± 6.19151	50.55 ± 43.87	22.43	1.95	0.0205 ± 0.0707
51_KMK-2	5 °C	4	0.2762329	1.149	265.9626	60.978	0.2666950	9.407	13.01527	0.628	148.9358	0.627	7.00618 ± 2.17753	50.34 ± 15.43	60.32	3.89	0.0207 ± 0.0253
52_KMK-2	5 °C	4	0.0984086	3.023	303.8510	62.665	0.2356036	8.707	15.12151	0.301	107.7680	0.866	7.00719 ± 2.19461	50.35 ± 15.55	96.89	4.53	0.0211 ± 0.0264
53_KMK-2	5 °C	4	0.0963079	3.187	216.0023	93.560	0.2716634	7.769	14.61645	0.193	114.3836	0.819	7.05201 ± 2.34370	50.67 ± 16.61	91.08	4.49	0.0288 ± 0.0539
55_KMK-2	6 °C	4	0.0998989	2.979	29.7181	649.669	0.2428391	8.840	14.53952	0.185	121.8410	0.766	6.52232 ± 2.27418	46.91 ± 16.15	77.72	4.41	0.2101 ± 2.7294
56_KMK-2	6 °C	4	0.1380666	2.228	-169.60	90.823	0.6075632	4.609	44.83913	0.130	369.8846	0.260	6.83980 ± 0.57628	49.16 ± 4.09	84.95	13.95	-0.1140 ± 0.2071
57_KMK-2	6 °C	4	0.1912850	1.639	143.1068	85.817	1.3379733	1.752	100.95234	0.115	758.6071	0.123	6.92242 ± 0.20665	49.75 ± 1.46	94.03	31.30	0.3030 ± 0.5201
4_KMK-2	7 °C	4	0.0770691	11.247	294.5770	98.924	0.4479906	5.973	33.31179	0.128	219.7423	1.161	6.55654 ± 1.48708	47.15 ± 10.56	100.90	10.27	0.0483 ± 0.0956
5_KMK-2	7 °C	4	0.0584562	14.822	403.3393	72.429	0.2896831	11.745	21.71829	0.152	117.2863	2.176	6.26068 ± 2.35533	45.05 ± 16.74	114.37	6.51	0.0228 ± 0.0331
6_KMK-2	8 °C	4	0.0724425	12.081	617.6113	48.098	0.3092851	10.266	22.86000	0.140	114.7255	2.224	6.49998 ± 2.29767	46.75 ± 16.31	126.97	6.81	0.0156 ± 0.0150
8_KMK-2	10 °C	4	0.0912916	9.525	704.0786	40.752	0.3388404	7.878	21.13750	0.219	103.3896	2.469	6.60114 ± 2.41926	47.47 ± 17.17	131.69	6.27	0.0126 ± 0.0103
9_KMK-2	12 °C	4	0.0525406	16.497	573.1992	66.571	0.2170933	14.152	14.13494	0.357	76.3872	3.342	7.98376 ± 4.79143	57.26 ± 33.83	143.38	4.17	0.0103 ± 0.0137
Σ			3.0041547	0.734	3748.6166	23.075	5.1447248	1.833	327.73499	0.060	2825.6367	0.223					
KPA-8		J = 0.0039438 ± 0.00000355															
Relative Abundances		36Ar [V]	%1σ	37Ar [V]	%1σ	38Ar [V]	%1σ	39Ar [V]	%1σ	40Ar [V]	%1σ	40(r)/39(k) ± 2σ	Age ± 2σ (Ma)	40Ar(r) (%)	39Ar(k) (%)	K/Ca ± 2σ	
31_KPA-8l	3 °C		2.8089159	0.536	234.32	130.296	0.6593686	3.542	3.20068	0.820	843.8298	0.072	11.13103 ± 17.59558	77.69 ± 120.21	4.00	1.52	0.0056 ± 0.0145
32_KPA-8l	4 °C		2.6898971	0.550	195.28	150.750	0.6711098	4.116	7.48457	0.411	829.3414	0.068	7.10266 ± 7.13672	49.96 ± 49.51	6.15	3.61	0.0162 ± 0.0488
33_KPA-8l	4 °C	4	1.7738239	0.533	261.41	118.841	0.5120498	5.930	10.48252	0.495	574.4086	0.096	7.18749 ± 5.30643	50.55 ± 36.80	12.60	5.06	0.0169 ± 0.0402
35_KPA-8l	5 °C	4	0.4962284	0.761	97.74	380.046	0.2574942	9.157	13.08041	0.307	212.9395	0.259	5.85080 ± 4.97662	41.25 ± 34.69	34.98	6.40	0.0572 ± 0.4350
36_KPA-8l	5 °C	4	0.2897844	0.903	213.48	232.238	0.2457546	6.535	11.22189	0.282	142.6239	0.386	6.78306 ± 7.64231	47.74 ± 53.08	52.63	5.56	0.0223 ± 0.1035
37_KPA-8l	5 °C	4	0.3112543	0.847	130.12	250.994	0.2640977	12.342	12.42831	0.238	162.4586	0.337	6.60376 ± 4.51844	46.49 ± 31.41	50.14	6.20	0.0408 ± 0.2046
39_KPA-8l	5 °C	4	0.1868086	1.810	23.81	913.916	0.9692660	2.109	71.99564	0.134	593.5181	0.148	7.49733 ± 0.51736	52.69 ± 3.58	90.92	36.17	1.2999 ± 23.7605
40_KPA-8l	6 °C	4	0.0683805	3.439	-745.35	46.385	0.3885421	7.195	32.44277	0.200	364.3965	0.454	8.68685 ± 1.84452	60.91 ± 12.72	76.96	16.22	-0.0190 ± 0.0177
41_KPA-8l	6 °C		0.0430701	2.523	-813.85	69.654	0.1578973	15.361	12.94441	0.371	248.4135	0.063	12.55212 ± 7.34728	87.37 ± 49.93	66.94	6.66	-0.0072 ± 0.0100
43_KPA-8l	6 °C		0.0425698	2.407	-572.74	96.309	0.1136482	26.458	8.10058	0.637	181.0269	0.085	14.35041 ± 11.36444	99.55 ± 76.71	66.08	4.19	-0.0064 ± 0.0123
44_KPA-8l	7 °C		0.0562862	1.242	-483.32	65.192	0.1757371	11.013	9.27267	0.519	174.7342	0.091	12.14450 ± 5.62442	84.60 ± 38.28	66.89	4.84	-0.0086 ± 0.0112
45_KPA-8l	8 °C		0.0601309	1.322	-240.19	109.146	0.1056815	20.501	7.09495	0.579	172.0936	0.089	18.80735 ± 6.36101	129.39 ± 42.23	77.76	3.58	-0.0130 ± 0.0284
Σ			8.8271500	0.273	-1699.29	77.343	4.5206470	1.934	199.74940	0.087	4499.7847	0.052					

TABLE DR2

NAV-3b J=0.00402950±0.00000f Spectro : ARGUS IV																
Relative Abundances	36Ar [V]	%1s	37Ar [V]	%1s	38Ar [V]	%1s	39Ar [V]	%1s	40Ar [V]	%1s	40*/39(k) ± 2s	Age ± 2s (Ma)	40Ar* (%)	39Ar(k) (%)	K/Ca ± 2s	
9_NAV-3B	4.0%	24.071859	0.571	974.8500	3.800	6.452027	0.489	29.2382	0.264	7129.981	0.048	3.49467 ± 3.35953	25.29 ± 24.14	1.40	1.97	0.013 ± 0.001
10_NAV-3	5.0%	11.580366	0.574	648.5615	5.972	3.480601	0.789	32.7046	0.237	3580.652	0.083	6.62952 ± 1.45997	47.67 ± 10.36	5.97	2.23	0.021 ± 0.003
12_NAV-3	5.5%	3.903645	0.616	440.0984	9.459	1.590267	1.424	37.1487	0.213	1332.082	0.227	5.85733 ± 0.52065	42.18 ± 3.71	16.19	2.55	0.036 ± 0.007
13_NAV-3	6.0%	1.468001	0.768	261.4635	33.774	0.998112	1.998	41.1356	0.212	651.098	0.458	5.84172 ± 0.43820	42.07 ± 3.12	36.74	2.83	0.067 ± 0.045
14_NAV-3	6.3%	0.626130	1.393	225.9253	12.936	0.901243	2.375	45.2578	0.204	420.622	0.721	5.64217 ± 0.21635	40.65 ± 1.54	60.49	3.12	0.086 ± 0.022
16_NAV-3	6.6%	0.450185	1.795	79.9672	112.109	0.649674	3.459	35.6219	0.236	327.276	0.914	5.64410 ± 0.48233	40.67 ± 3.44	61.33	2.46	0.191 ± 0.429
17_NAV-3	6.9%	0.391032	2.037	142.1313	56.491	0.620744	3.355	33.6207	0.222	311.508	0.958	6.19825 ± 0.47011	44.61 ± 3.34	66.69	2.32	0.101 ± 0.115
18_NAV-3	7.2%	0.383583	2.104	36.5946	75.686	0.679554	3.692	38.3892	0.211	344.616	0.885	6.10043 ± 0.23878	43.91 ± 1.70	67.91	2.65	0.451 ± 0.682
20_NAV-3	7.5%	0.617538	1.367	-31.8060	80.778	1.253744	1.678	78.9865	0.162	680.972	0.443	6.26551 ± 0.11657	45.09 ± 0.83	72.70	5.46	-1.068 ± 1.726
21_NAV-3	7.8%	0.845306	5.636	17.6699	160.546	2.497727	1.983	170.3464	0.155	1357.707	0.693	6.50367 ± 0.20211	46.78 ± 1.44	81.59	11.77	4.145 ± 13.310
22_NAV-3	8.1%	1.375291	5.032	30.9040	95.668	6.908670	1.099	507.1068	0.144	3567.849	0.541	6.23009 ± 0.11296	44.84 ± 0.80	88.55	35.05	7.056 ± 13.500
24_NAV-3	8.4%	0.427832	1.741	123.3786	33.397	1.980591	1.118	144.0581	0.143	941.179	0.185	5.72255 ± 0.06529	41.22 ± 0.47	87.54	9.95	0.502 ± 0.335
25_NAV-3	8.7%	0.347291	2.174	49.9188	77.368	1.043066	2.383	69.6635	0.161	458.658	0.478	5.16493 ± 0.13201	37.25 ± 0.94	78.41	4.81	0.600 ± 0.928
26_NAV-3	9.0%	0.273815	2.708	2.3805	946.327	0.670710	3.440	47.3429	0.189	301.250	0.728	4.64885 ± 0.15548	33.56 ± 1.11	73.06	3.27	8.551 ± 161.850
28_NAV-3	10.0%	0.424388	1.799	43.4034	57.291	0.923435	2.180	61.7075	0.173	406.348	0.539	4.60544 ± 0.12494	33.25 ± 0.89	69.90	4.26	0.611 ± 0.700
29_NAV-3	12.0%	0.597971	1.363	-2.0230	1208.846	1.190922	2.100	76.8151	0.165	517.295	0.424	4.42211 ± 0.10297	31.94 ± 0.74	65.67	5.31	-16.328 ± 394.761
S		47.784230	0.373	3043.4182	6.214	31.841089	0.397	1449.1435	0.060	22329.092	0.107					

NAV-3c J=0.00402950±0.00000f Spectro : ARGUS IV																
Relative Abundances	36Ar [V]	%1s	37Ar [V]	%1s	38Ar [V]	%1s	39Ar [V]	%1s	40Ar [V]	%1s	40(r)/39(k) ± 2s	Age ± 2s (Ma)	40Ar(r) (%)	39Ar(k) (%)	K/Ca ± 2s	
30_NAV-3	4.0%	8.752647	0.596	426.7565	8.165	2.514326	1.690	13.5829	0.722	2603.852	0.128	4.05000 ± 2.78278	29.27 ± 19.95	2.06	1.15	0.013 ± 0.002
33_NAV-3	5.0%	11.214253	0.583	585.3519	5.546	3.508387	1.328	33.1208	0.329	3477.948	0.095	6.53924 ± 1.40857	47.03 ± 10.00	6.15	2.83	0.024 ± 0.003
34_NAV-3	5.5%	2.908345	0.749	363.0345	9.858	1.302245	2.816	25.2080	0.399	989.550	0.338	6.44893 ± 0.68821	46.39 ± 4.89	16.26	2.16	0.030 ± 0.006
35_NAV-3	6.0%	1.394728	1.125	296.6431	15.241	0.865514	4.489	26.5601	0.387	549.448	0.614	6.16229 ± 0.54708	44.35 ± 3.89	29.55	2.28	0.038 ± 0.012
37_NAV-3	6.3%	0.530153	2.671	210.5173	22.268	0.707024	4.914	31.4310	0.456	319.566	1.060	5.77253 ± 0.43768	41.58 ± 3.12	56.50	2.71	0.064 ± 0.028
38_NAV-3	6.6%	0.335936	4.083	277.1481	24.569	0.528844	6.699	26.4947	0.450	241.116	1.412	6.28306 ± 0.60876	45.21 ± 4.33	68.52	2.28	0.041 ± 0.020
39_NAV-3	6.9%	0.241142	5.628	141.6432	27.439	0.412575	9.544	21.1454	1.127	190.018	1.741	6.20784 ± 0.60618	44.68 ± 4.31	68.74	1.82	0.064 ± 0.035
41_NAV-3	7.2%	0.242656	5.559	143.1868	33.818	0.448551	8.689	26.2855	0.391	224.302	1.491	6.28464 ± 0.51255	45.22 ± 3.64	73.36	2.27	0.079 ± 0.053
42_NAV-3	7.5%	0.329281	4.115	92.9742	36.342	0.553563	6.936	27.9930	0.392	269.873	1.226	6.45414 ± 0.43010	46.43 ± 3.05	66.78	2.42	0.129 ± 0.094
43_NAV-3	7.8%	0.479685	2.867	19.0977	187.128	0.831439	4.518	46.1147	0.251	435.650	0.758	6.40131 ± 0.26571	46.05 ± 1.89	67.74	4.00	1.038 ± 3.885
45_NAV-3	8.1%	0.382928	3.562	27.6454	116.220	0.714079	5.614	44.3478	0.257	406.026	0.814	6.65059 ± 0.26851	47.82 ± 1.91	72.61	3.84	0.689 ± 1.603
46_NAV-3	8.3%	0.441581	7.202	-103.9544	75.409	1.527758	2.373	108.2819	0.178	879.165	0.975	6.81805 ± 0.26662	49.01 ± 1.89	84.03	9.39	-0.448 ± 0.676
47_NAV-3	8.6%	0.825150	3.863	17.7697	164.163	5.198367	0.801	390.9425	0.145	2804.071	0.284	6.54328 ± 0.06727	47.06 ± 0.48	91.22	33.89	9.460 ± 31.059
4_NAV-3C	9.0%	0.296675	10.789	66.4975	76.241	1.840361	1.671	134.9982	0.165	909.669	0.845	6.12346 ± 0.19278	44.08 ± 1.37	90.84	11.70	0.873 ± 1.331
5_NAV-3C	9.5%	0.274793	9.852	105.6833	37.942	1.096781	2.486	79.8466	0.224	497.948	1.437	5.32747 ± 0.28379	38.41 ± 2.02	85.34	6.91	0.325 ± 0.246
6_NAV-3C	10.0%	0.265903	10.089	38.1470	42.728	0.770335	2.495	53.8095	0.283	325.930	2.196	4.65000 ± 0.40155	33.57 ± 2.87	76.73	4.66	0.606 ± 0.518
8_NAV-3C	12.0%	0.370818	7.520	48.8411	48.762	1.023809	2.959	65.5177	0.249	401.063	1.788	4.50543 ± 0.34034	32.53 ± 2.44	73.56	5.68	0.577 ± 0.562
S		29.286675	0.408	2756.9828	6.462	23.843957	0.634	1155.6803	0.071	15525.196	0.140					

NAV-3f J=0.00272700±0.00000f Spectro : VG 3600																
Relative Abundances	36Ar [V]	%1s	37Ar [V]	%1s	38Ar [V]	%1s	39Ar [V]	%1s	40Ar [V]	%1s	40(r)/39(k) ± 2s	Age ± 2s (Ma)	40Ar(r) (%)	39Ar(k) (%)	K/Ca ± 2s	
33260_AR	600 °C	0.0000303	53.662	0.0003365	92.443	0.0000662	26.204	0.0008544	5.743	0.0137850	2.924	5.68804 ± 11.30657	27.83 ± 54.90	35.25	0.15	1 ± 2
33261_AR	650 °C	0.0000194	75.313	0.0004149	74.920	0.0001216	13.749	0.0038414	0.692	0.0686538	0.572	16.37933 ± 2.27331	79.02 ± 10.73	91.64	0.66	4 ± 6
33262_AR	700 °C	0.0000745	19.418	0.0000597	521.472	0.0001942	8.529	0.0083573	0.612	0.0925542	0.431	8.43301 ± 1.03260	41.11 ± 4.98	76.15	1.44	60 ± 628
33263_AR	750 °C	0.0003261	15.991	0.0000765	589.195	0.0003058	8.150	0.0156988	0.216	0.1257348	0.765	1.86222 ± 1.96744	9.16 ± 9.65	23.25	2.70	88 ± 1040
33264_AR	800 °C	0.0001342	38.761	-0.0001787	252.192	0.0004710	5.334	0.0288385	0.219	0.0990994	0.956	2.05119 ± 1.06820	10.09 ± 5.24	59.69	4.95	-69 ± 350
33265_AR	850 °C	0.0000968	53.933	0.0001908	236.068	0.0010055	3.966	0.0689251	0.184	0.1679506	0.563	2.01224 ± 0.44874	9.90 ± 2.20	82.58	11.84	155 ± 733
33266_AR	900 °C	0.0002073	10.139	0.0007232	42.946	0.0023209	1.764	0.1894340	0.207	0.3999653	0.176	1.77859 ± 0.06645	8.75 ± 0.33	84.24	32.53	113 ± 97
33267_AR	950 °C	0.0001615	4.771	-0.0005931	35.101	0.0012941	1.698	0.1271947	0.164	0.2720278	0.281	1.75335 ± 0.03830	8.63 ± 0.19	81.98	21.84	-92 ± 65
33268_AR	1000 °C	0.0000944	9.413	0.0002402	87.036	0.0008270	2.667	0.0582583	0.180	0.1295069	0.491	1.73497 ± 0.09302	8.54 ± 0.46	78.05	10.00	104 ± 182
33270_AR	1450 °C	0.0001905	27.384	0.0005923	76.000	0.0010664	2.254	0.0809093	0.187	0.2736496	0.432	2.67738 ± 0.38234	13.15 ± 1.87	79.16	13.89	59 ± 89
S		0.0013351	8.254	0.0018623	60.878	0.0076728	1.084	0.5823119	0.087	1.6429275	0.150					

TABLE DR3 (continued)

Step	Temperatur °C	40Ar/39Ar	38Ar/39Ar	37Ar/39Ar	36Ar/39Ar (E-3)	39Ar (E-14) moles	%39Ar released	%40Ar*	40Ar*/39K	Age ± s.d. Ma
NAV-4f										
Spectro : VG 3600										
J= 0.002727										
1	350	233329.16	1.04	66.39745	1039.77	0	0	99.87	0	0 ± 0
2	450	477.918	0.409	0.01935	1647.584	0.015	3.91	-1.88	-8.96	0 ± 0
3	550	8.185	0.118	0.0005	22.722	0.016	8.2	17.72	1.45	7.121 ± 3.503
4	600	1644181.3	1.04	66.39745	5898367.2	0	8.2	-6.01	-103237.8	0 ± 0
5	650	8.153	0	0.00269	77.269	0.003	9.01	-180.3	-14.7	0 ± 0
6	700	6.563	0.021	0.00012	7.719	0.067	26.42	64.94	4.26	20.848 ± 0.988
7	750	6.456	0.008	0.00016	7.304	0.052	40.09	66.26	4.28	20.923 ± 1.073
8	800	10.924	0.015	0.00011	22.598	0.073	59.13	38.69	4.23	20.676 ± 1.731
9	850	6.353	0.014	0.00011	0.002	0.075	78.77	99.68	6.33	30.89 ± 0.089
10	900	8.322	0.012	0.00021	0.003	0.039	89.07	99.75	8.3	40.385 ± 0.211
11	950	26.545	0.032	0.00157	24.211	0.005	90.45	72.97	19.37	92.865 ± 18.47
12	1000	31.473	0	0.00246	18.469	0.003	91.33	82.6	26	123.561 ± 8.969
13	1200	6.321	0.002	0.00038	8.767	0.022	97.05	58.7	3.71	18.162 ± 2.383
14	1450	8.936	0	0.00073	25.99	0.011	100	13.83	1.24	6.07 ± 6.384
CAU-3f										
Spectro : VG 3600										
J= 0.002727										
1	550	5.946	0.211	0.00042	27.703	0.019	4.68	-38	-2.26	0 ± 0
2	600	4.964	0.071	0.00056	0.009	0.014	8.2	99.54	4.94	24.146 ± 0.216
3	650	6.494	0.026	0.00027	3.348	0.029	15.57	84.46	5.49	26.786 ± 1.404
4	700	6.145	0.021	0.00017	6.785	0.047	27.41	67.05	4.12	20.159 ± 2.29
5	750	5.065	0.02	0.00008	3.36	0.097	52.01	80	4.05	19.826 ± 1.234
6	800	9.738	0.021	0.00009	19.76	0.088	74.3	39.84	3.88	18.985 ± 0.934
7	850	5.171	0.02	0.00013	2.684	0.061	89.72	84.27	4.36	21.315 ± 0.494
8	950	11.343	0.048	0.00039	9.709	0.02	94.84	74.53	8.45	41.119 ± 1.734
9	1000	28.501	0	0.00398	61.908	0.002	95.34	35.75	10.19	49.439 ± 11.841
10	1200	13.136	0.009	0.53894	10.137	0.015	99.15	77.29	10.16	49.29 ± 1.805
11	1450	52.209	0	30714.419	59.822	0.003	100	99.96	-99.93	0 ± 0

TABLE DR3 (continued)