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Sediment budget of the Napo River, Amazon basin, Ecuador and Peru

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Abstract

The upstream-downstream sediment budget along the Napo River (100,520 km², 6,300 m³ s⁻¹) was studied in the Andean Foothill of Ecuador, at the west of the Amazon basin. A comparative study was made, during 4 hydrological cycles (2001-2005) for 3 hydrological stations located upstream; and during 1 hydrological cycle (2004-2005) for the fourth one located near the mouth of the Napo River (region of Iquitos in Peru). This analysis showed an unusual increase in the concentration of suspended sediment recorded for the western part of the Amazon plain. These relatively high erosion rates for a floodplain basin (1,160 t km⁻² year⁻¹, i.e. 47% of TSS export at the exit of Ecuador) can be explained by a steeper slope than in the rest of the Andean foothills, where typically sedimentation phenomena are predominant.

Here, the tectonic uplift causes the Pastaza Megafan's existence. This progressively diverts the course of Napo River towards north and also provokes the remobilisation of fine fluvial deposits. Moreover, this geodynamic trend is completed by the impact of volcanic eruption, earthquakes and landslides. The combination of these phenomena, so common in the region, has provided a large sediment transfer, not only at present but also in the past, as can be confirmed by the presence of incised terraces, mainly formed by volcanic materials.

Then, these results were compared with a similar study carried out further South in the Madeira Basin at the foot of the Bolivian Andes. These studies show the spatio-temporal variability of the relation between sediment transfer and geodynamic processes at the Andean Piedmont.

Key Words : Amazon River basin, Sediment yield, Tectonics, Pastaza Megafan, Andes, Napo basin

I. INTRODUCTION

The Amazonian basin (Fig. 1), the largest on the planet with $6,2 \cdot 10^6$ km², provides $6,6 \cdot 10^{12}$ m³ of water (Molinier *et al.*, 1996) and $800 \cdot 10^6$ tons per year of sediment to the Atlantic Ocean (Guyot *et al.*, 2005). It represents, respectively, 16.6 and 6 % of the discharge and sediment planet contribution from the continents to the oceans, according to the world values of $39,7 \cdot 10^{12}$ m³ of water (Baumgartner et Reichel, 1975) and $13,500 \cdot 10^6$ tons of sediment (Milliman and Meade, 1983). Most of the sediment discharged (95%) to the Atlantic Ocean by the Amazon River comes from the Andes Mountains although the range only covers 12% of the surface area of the Amazon basin.

On the Brazilian side of the Amazonian basin, the sedimentary transport has already focused many works aimed at quantifying them (Gibbs, 1967, Meade *et al.*, 1985, Dunne *et al.*, 1998, Filizola, 2003, Moreira-Turcq *et al.*, 2004, Guyot *et al.*, 2005 and Laraque *et al.*, 2005). The sediment transport processes from the rivers draining the western part of the Amazonian plain located at the Andean Cordillera Piedmont are essentially reported in Bolivia at the southwest of the basin. Here, important sediment deposits are observed (Guyot *et al.* 1996), within a geodynamic context of subsidence (Horton et De Celles, 2001, Cornelius *et al.*, 2005, Baby *et al.*, in press).

Further north, in the Peruvian and Colombian part of the Amazonian basin, a lack of studies of the actual suspended sediment yield is evidenced, due to the distancing of these regions, at times of difficult access and /or to their insecurity.

However some recent works in tectonics, geodynamics and geomorphology have been carried out in the eastern ecuadorian and northern peruvian Andes piedmont, through some field campaigns and analysis of satellite images. These studies point out a tectonic uplift, that started at Neogene, and its consequences over the relief and erosion (Burgos *et al.*, 2004, Christophoul *et al.*, 2002a & b, Bès de Berc *et al.*, 2005). This tectonic deformation has led to an increase of the river's slopes. Consequently, these ones have tended to regain its equilibrium profile by increasing their valleys (Bès de Berc, 2003). In fact, a first exploratory work of the sediment flux through the Ecuadorian Amazon, during a single hydrologic cycle, evidenced an actual erosive process in that region (Laraque *et al.*, 2004). The same process was observed more downstream in the Peruvian part of Napo basin (Guyot *et al.*, 2007).

Along different mountain ranges of the planet in general (Hovius, 1998, Leeder *et al.*, 1998, Hsieh et Knuepfer, 2001), and in South America in particular, some works (Weng *et al.*, 2002, Keefer *et al.*, 2003, Aalto, *et al.*, 2003) have already put in evidence the dependance of fluvial transport with climate and tectonic. These relations complete those well-known processes of sediment transfer (erosion and/or transport and/or sedimentation) with the physiographic characteristics (climate, morphology, vegetation, pedologic coverings, geology, land usage, drainage network, ...) of mountainside basins, as illustrating by Restrepo *et al.*, (2006) in the Magdalena basin in Colombia. Notwithstanding, theses aspect are not still studied for the eastern piedmont of the Ecuadorian Andes and its plains, a region known for its great geodynamic activity, subject to frequent and continuous seismic and volcanic episodes (Yepes *et al.*, 1996) and numerous landslides (Zevallos *et al.*, 1993).

In order to resolve this problem, a first multiannual monitoring attempt of the sediment flow

was carried out on the hydrographic basin of the Napo River which drains the eastern piedmont of the Ecuadorian Andes. The data come from the HYBAM programme (www.mpl.ird.fr/hybam), which started in 2001 in Ecuador.

This work has been organized in the following way: - Part 2, a description of the studied region; - Part 3, data and methods used; Part 4, a result-based discussion. This last one addresses questions regarding spatio-temporal variability that evidenced sediment processes, at the Napo and at the Amazon River basins level.

II. STUDIED AREA

Geography

The whole Napo River drainage basin is situated between 00°10'N and 01°30'S, 75°20' and 78°40'W with an altitude ranging from 100 to 6,300 m.a.s.l. Its area covers 100,500 km², distributed among Ecuador (59.6%), Peru (40%) and Colombia (0.4%). On these latitude, the ecuadorian part of the Andean Cordillera is the narrowest within the Central Andes and presents important geodynamic activity (volcanic and seismic).

In Ecuador, where this study is more advanced, the Napo basin covers an area of 27,000 km², i.e. equivalent to 20% of the eastern part of this country. The Ecuadorian part of the Napo basin crosses the "Cordillera Real"; whose eastern side has steep slopes that descend from 6,000 to 500 masl over only 100 kilometers.

Geology

The Napo River drainage basin is located in the eastern sediment basin of Ecuador and Peru, which has sustained an uplift and tectonic deformation since Neogene (Bès de Berc, 2003). Drained lands in Ecuador, are essentially made up of a series of volcano-sediments; while in Peru drained lands are fluvial and tidal deposit (Räsänen *et al.*, 2005, Wesselingh *et al.*, 2006). The course of Napo River get around the northern limit of the "Pastaza Megafan" situated

between 0°00' and 5°00' S and 73°00' and 78°00' W with an altitude between 90-300 m.a.s.l. and an area of 40 300 km² (Bernal *et al.*, personal communication). It is the greatest tropical alluvial fan in the world according to Räsänen *et al.*, (1992).

Vegetation covers

Below 500 m high, at the exit of mountainous reliefs, a fairly flat and homogeneous forested ecosystem is found, which can be defined as part of the amazonian plains. This ecosystem covers the main part (82%) of the total surface of the Napo basin and it is located in Ecuador and Peru. The rest (18%) in Ecuador is constituted by the Andean Cordillera closing off the basin on the west (Fig. 1). The andean vegetation covers are composed by “puna” and “paramo”, which are typical short vegetation covers like a kind of highland tundra.

Climat

The whole drainage basin receives 3,200 mm yr⁻¹ of rainfall with a main rainfall season between May and June, with a monthly maximum of 330 mm, and a second rainfall season between September and November with a monthly maximum of 250 mm (Fig. 2).

The Ecuadorian Napo region receives high precipitation (2,900 mm.yr⁻¹) and presents very irregular runoff regimes, characterized by flash floods, reflecting their extreme sensitivity to rain events (Laraque *et al.*, 2007). The same authors showed that rainfall and discharge regimes are highly variable in space. They depend on altitude so that different regimes may be found in the same basin. At high locations, rainfall presents a marked bimode, with a maximum in April and October and a minimum in July-August and in January. This rainfall seasonality is associated with the annual progress of the sun and with the seasonal variability of the zonal wind that uplifts moist air from the Amazon basin. In small to medium upper basins, two discharge maxima occur along with the two rainfall maxima. However during the

July-August period, discharge values do not decrease, probably and mainly because of ground water running, of the dominating upper highlands "paramo" formation. In the lowlands, rainfall and discharge seasonal variability is very weak. Relative maxima are observed in April for rainfall, June for discharge; secondary small maxima are observed in October (rainfall) and November (discharge).

On the Peruvian part of the Napo River basin, the interannual rainfall average is $3,310 \text{ mm}\cdot\text{yr}^{-1}$.

Hydrology

Close to 6,000 m.a.s.l. of altitude, the origin of the Napo River is from the thawing of glaciers which covers the Cotopaxi (5,897 m.a.s.l.), Antisana (5,700 m.a.s.l.) and Llanganates (4,571 m.a.s.l.) volcanoes. Within the first half of the river course in ecuadorian territory, the Napo River presents a torrential to turbulent behavior. It leaves Ecuador at an altitude of 189 masl and at about 400 km from its headwaters, it receives the contribution of the Aguarico River and enters Peru to travel 600 km.

At the river mouth near Bellavista gauging station, the Napo River drains an area of 100,520 km^2 and presents a mean annual discharge of $6,300 \text{ m}^3 \text{ s}^{-1}$ (that mean $62 \text{ l s}^{-1} \text{ km}^2$) which corresponds to 4% of the Amazon River recorded at Obidos, an hydrometric station located the furthest downstream in Brazil.

The Napo River is the main eastern river of Ecuador (Fig. 1) with an annual discharge of $2,210 \text{ m}^3 \text{ s}^{-1}$ at Nuevo Rocafuerte station on the outlet of this country. Its area is $26,860 \text{ km}^2$, with a solid discharge equal to $82 \text{ l s}^{-1} \text{ km}^2$. In Ecuador, the Napo River drains almost 3% of the whole Amazonian catchment.

The Napo is a «white water» river type. Typically, these rivers have high levels of total suspended sediments, due to the mechanical erosion of the Andean chain. In fact, according to Guyot (2007), the sediment transport's Napo River is estimated at $49 \cdot 10^6 \text{ ton year}^{-1}$, which

corresponds to 6% of the sediment flow passing through Obidos station in Bresil.

III. METHODS AND DATA

To understand the sediment dynamic at the Andean Piedmont of the Napo River basin, an integrated approach using hydrology and geomorphology was used. The hydrology was studied through out the implementation of hydrologic and sedimentologic measuring stations, along the course of the Napo River, which drift off perpendicularly to the Andes to penetrate the Amazonian plains. Then, a comparative upstream-downstream study of sediment budgets was performed using the collected data.

Figure 1 shows the location of the stations used on the main node of this schematic fluvial system. San Sebastian (SEB) and Francisco de Orellana stations (FDO) control the output of the Andes, and Nuevo Rocafuerte station (ROC) closes the basin at the outlet from Ecuador. In Peru, Santa Clotilde (CLO) is located after the confluence with Curaray River and, finally, Bellavista station (BEL) controls the whole basin (Table 1). Besides, a field scientific crossing along the Napo River during the falling stage (October, 2004) has enabled the study of the longitudinal evolution of the TSS discharges, concentrations and flows (Fraizy et al., 2005).

The geomorphology was studied through out the longitudinal profile and the spatial variations of percentage in sand, water and vegetated island. The construction of longitudinal profile and stream slope is another result of the field scientific crossing along the Napo River on October 2004. We used for this DGPS survey, a base station and cinematic station with bifrequencies AShtech ZXTREM and Trimble 5700 receptors (Bourel et al, 2007). This survey covered the Napo River from Puerto Misahualli, in the Subandean Zone, to the confluence with the Marañon River along 1100 km (Fig. 3), 13500 measurements were realized with an elevation accuracy smaller than 0.5 m and a mean RMS (Root Mean Square) of 0.1 m.

Spatial variations of percentage in sand, water and vegetated island was evaluated using

Landsat images (TM and ETM+, spatial resolution 30 m) of 1986, 1989, 1995, 1996 and 2002. They were collected during the falling stage on the reach between FDO and ROC stations. The limitations of satellite coverage required us to divide the segment ROC -FDO in two areas (ROC-X / X-FDO, Fig. 3). Both set of images were studied for the same months to point out and quantify the sand islands changes in term of sedimentary supply. To quantify these variations of percentage on the Napo River, we classified satellite images through the parallelepiped threshold. In order to do that, three homogenous classes, taking into account spectral signature, were defined as: water, sand and vegetated island.

III. 1 – WATER LEVEL AND DISCHARGE MEASUREMENTS

In these gauging stations, the observers carried out daily or twice a day limnometric measurements. Frequent gaugings were carried out along five hydrological cycles using a 1,200 kHz ADCP (Acoustic Doppler Current Profiler) to establish gauging curves.

III. 2 – SUSPENDED SEDIMENT SAMPLING AND FLOW CALCULATIONS

At the three Ecuadorian stations water surface samples were taken in the middle of the river by an observer, every 10 days since 2001. The samples were filtered using 0.45 μm cellulose acetate filters to separate the TSS. Cerón (2004) and Laraque *et al.*, (2004) studied the relations between these surface TSS concentrations and the average of TSS concentrations in the cross section. As described by Filizola & Guyot, (2004), we used a point sampling protocol to determinate the TSS concentrations in different points of the cross sections. This was made at different stages during several complete annual hydrological cycles.

Figure 4 shows the established relation for SEB and FDO stations, located at the exit of the Andean Piedmont. The calculated mean TSS concentrations in these stations are only 5% higher than those TSS concentrations taken by the observers at the surface. This is due to the

weak heterogeneities of velocities and specific TSS concentrations within the measured profiles, which illustrate much more turbulent flows marked by numerous swirls, assuring a continuous mixing of particles in suspension. Further downstream in the Amazonian basin near the ROC station, the river sections become much more heterogeneous with profiles characterized by higher velocity gradients and TSS load (Filizola, 2003, Laraque *et al.*, 2005). During the hydrologic cycles, the peaks of TSS concentrations precede flood peaks at times, overlapping them or occurring afterwards. The shift of flood and TSS peaks is larger in the middle of the Amazonian plains where the former always precedes the latter, at times for many months (Laraque *et al.*, 2005). This observation invalidates the use of relations (TSS = f(Q)) to calculate sediment discharge. To overcome this obstacle, a frequency of 10-days samplings, much more adapted to the behaviour of these rivers, has been implemented. Such a frequency, applied over many hydrologic cycles, enables the obtaining of more accurate calculation of the TSS yields.

As the frequency of the discharge daily values for the studied stations is higher than those of TSS (10 days), the intermediate TSS values were calculated by linear interpolation and multiplied by the discharge to obtain the suspended load. It is possible to obtain a daily suspended sediment yield using the Hydraccess software (www.mpl.ird.fr/hybam/outils/hydraccess.htm).

The monthly and annual sediment flows come from the arithmetic mean of corresponding time series. Monthly and annual TSS concentrations are calculated by dividing the corresponding suspended sediment yield by discharges.

The 4 annual April-March cycles of the 2001–2005 period studied in Ecuadorian station and the April-March cycles of the 2004–2005 period studied in Peru (BEL station) , were used to compare the stations budget from upstream to downstream Napo River. The results are presented in tables 2 and 3.

IV. RESULTS & DISCUSSION

IV. 1 - Temporal Variability of suspended sediment yields

For the studied stations of the Napo basin, the monthly hydrograms are slightly bimodal with a main period of high waters (from February to September) finishing between May and July, and a short second season of floods during, in general, from mid-November to January.

Table 2 shows the range of variation of discharges, TSS concentrations and flows studied, at daily, ten-days and monthly time spans throughout the entire group of stations during the studied period. The 10-day instantaneous TSS concentrations vary between 2 and 2,806 mg l⁻¹ according to the station and season. As much as for the extreme values as for the mean ones, those of SEB are the highest of the four studied stations. Regarding the reports of the extreme concentrations for each station, those at ROC and at SEB are identical for the max/min ratios (281). This ratio reports along with the one at FDO (478) are much more superior to that of BEL (11). The higher ratios of the three first stations are due to their weak minimal TSS concentrations in low water periods, between 2 and 10 mg l⁻¹, which are common in mountainous rivers. The extreme daily discharges ratios, which are smaller than those of TSS concentrations, are in a diminishing order equal to 13 (FDO), 10 (BEL), 7 (SEB) and 5 (ROC). At a monthly time span, these extreme ratios are slightly reduced.

The two annual cycles from april 2001 to march 2002 and from april 2004 to march 2005 are similar with important increase (around 50%) of upstream-downstream TSS concentrations and suspended sediment yields (Table 3). On the other hand, the two other cycles from april 2002 to march 2003 and from april 2003 to march 2004 show a weaker increase of upstream-downstream TSS concentrations (10 to 30%) and suspended sediment yield.

Figure 5a shows the monthly variability of the upstream-downstream suspended sediment yield in the ecuadorian part of the Napo basin (ROC-SEB-FDO) during the 2001-2005 period studied. It shows the important temporal variations marked essentially by two irregular

erosion phases (April, 2001-October, 2002 and May, 2003-April, 2005) alternated with two sedimentation phases of weaker amplitude (November, 2002-April, 2003, and May, 2005-September, 2005).

Figure 5b shows the monthly variability of the upstream-downstream suspended sediment yield in the Peruvian part of the Napo basin (BEL-ROC) during the common July 2004 - September 2005 period studied. This positive budget probably comes from the two main tributaries: the Aguarico and Curaray rivers, unfortunately poorly studied. On the left bank, at the north part of the Napo basin, the Aguarico River drains in its upstream reach, the Andean cordillera. On the right bank, the Curaray River drains a vast ancient alluvial plain in uplift process, which originates the Pastaza Megafan.

IV. 2 – Spatial variability of actual sediment transfers

During the studied period, the Andean part of the Napo basin controlled by FDO and SEB stations, produced an interannual mean of $12 \cdot 10^6$ ton year⁻¹ of TSS, but a total of $22.6 \cdot 10^6$ ton year⁻¹ left the Ecuadorian Napo Basin downstream at ROC (Table 3). The area between these stations covers $9,150 \text{ km}^2$ and provided a net gain difference of $10.6 \cdot 10^6 \text{ t year}^{-1}$ of TSS. This amount corresponds to 47% of the Napo River sediment yield at ROC. Also, the specific sediment yield from the lower part of the basin is about $1,160 \text{ t km}^{-2} \text{ year}^{-1}$. This upstream-downstream increase of suspended sediment yield stems from an increase of Q and TSS concentrations. These variations seem very high, considering there are almost no uplands in this intermediary area, which is protected by a dense forest cover.

If on one hand, the interannual budgets of the studied period indicate a sediment production (erosion) of $10.6 \cdot 10^6 \text{ ton yr}^{-1}$ for the Andean Piedmont located between SEB+FDO and ROC, on the other hand, the annual budgets are very variable, with erosion ratios which are consecutively in the following order 16, 5, 6.6 et $14.5 \cdot 10^6 \text{ ton yr}^{-1}$ for the 4 annual cycles

studied.

The interannual erosion rate (TSS in $\text{t km}^{-2} \text{ yr}^{-1}$) corresponds to the discharge or sediment yield normalized for the drainage area. These are the only units which enable the comparison among them of the basin transport dynamics, whether it is due to erosion or sedimentation. Although the Coca basin has the lowest specific discharge ($66 \text{ l s}^{-1} \text{ km}^2$) of the sub basins studied in Ecuador, and the smallest watershed area (one third of the watershed area controlled by FDO), it presents the maximum specific sediment rate ($1,062 \text{ t km}^{-2} \text{ year}^{-1}$), twice the Napo basin at FDO ($515 \text{ t km}^{-2} \text{ year}^{-1}$), and the interannual suspended sediment yield are similar at the two stations. This result stems from the high TSS concentrations of SEB (487 mg l^{-1}), three times the value of FDO (166 mg l^{-1}) (Tables 1 & 3).

Such a difference encountered in a specific sediment yield between these two rivers can be explained from their longitudinal profile. The Coca River rises in the eastern flanks of Reventador volcano, and flows for 209 km on the Subandean with a steepest slope of 38.8 cm km^{-1} . The Napo River is 1,075 km long, ranging 301 m in altitude (404–103 m above sea level) from Puerto Misahualli, in the Subandean, toward the confluence with Marañón River, with a strong variation of its slope (Fig. 3). Notwithstanding, upstream crossing the foothills of the Subandean Zone, both rivers profile (Coca & Napo river) are very steep and characterized by a strong slope ranging from 1.7 m km^{-1} to 35 cm km^{-1} . The difference begins at FDO, when the Coca River debouches into Napo River, and the last one shows a pronounced break in the profile that marks the beginning of Amazonian basin with a drop in the mean slope of 35 cm km^{-1} (Table 1). This slope stays roughly constant for the first 300 km in a transition area between FDO and ROC, even if some slope breaks can be observed due to the transverse profiles for navigation channel optimisation (Fig. 3). Likewise, this slope is more than 10 times higher than in the rest of the Amazon basin, and therefore it promotes a water velocity that is strong enough to create a relatively rectilinear course of the Napo River

as it leaves the Andes (Fig. 1). The increase in velocity results in an increase in the erosive and transport capacity of the river. Consequently, in the plains, the value for specific erosion is only a little smaller than in the mountains (1,160 against 1,577 t km⁻² year⁻¹) (Table 3). This sedimentary remobilization is applied on fluvial deposits essentially constituted by fine sand particles like shown by the d50 (0,2 to 0,5 mm) of the Hybam Napo'04 field trip (Fraizy *et al.*, 2005). These values are intermediates between those measured by Guyot *et al.*, (1999), on the Andean outlet (d50 ~30 mm) and the very flat plain “Llanos” (d50 ~11 mm) area in Madeira drainage basin in Bolivian Amazonia.

These observations were also confirmed by the results of the **scientific field trip** along the Napo River, which took place during the falling stage of October, 2004 (Table 4 and Fig. 1). In Ecuador, the TSS concentrations and suspended sediment yield increased upstream-downstream respectively from 45 mg l⁻¹ and 991 ton d⁻¹ at FDO to 132 mg l⁻¹ and 13,800 ton d⁻¹ at ROC. The suspended sediment yield gained 12,809 ton d⁻¹ between these two stations, stem partially from contributions made from the Coca River (1,778 ton d⁻¹). The rest (11,031 ton d⁻¹) is due to the remobilization of old alluvial deposits of the Napo River. This phenomenon's river basin is so relevant in this area considering the weaker TSS concentrations observed in its main tributaries (33.2 mg l⁻¹ for the Coca River and 24.3 mg l⁻¹ for the Tiputini River), which in reality might have been diluted in the Napo River waters.

But, since a flood was occurring during the field trip, it is necessary to study with caution the measured values. It could be the reason for the relatively high SM concentration of 199 mg l⁻¹. The suspended sediment yield gained 61, 550 ton d⁻¹ between BEL-ROC, stem mainly from the increasing of SM concentration during the flood rather that from the tributaries input that reached only 13,300 ton d⁻¹.

The next reach of Napo profile, between BEL and ROC (Fig. 3), are characterized by a general diminishing in the mean slope (11 cm km⁻¹), which is three time smaller that the one

at ROC-FDO section. Nonetheless, we obtain always the same conclusions about a mean positive sediment budget encountered during an entire yearly cycle (Fig. 5b). It amounts to 19.8 t day^{-1} of TSS, coming from a stream of 540 km long which drained an area of 73,660 km^2 (73.3 % of the total basin area). These values correspond to a specific discharge of $64.8 \text{ l s}^{-1} \text{ km}^2$ and a specific solid discharge of $268 \text{ t km}^{-2} \text{ year}^{-1}$. The fluxes sediment contribution of this intermediate area (BEL-ROC) corresponds to 42.5% of the Napo River sediment yield at BEL (table 3). This contribution is more related to the high annual discharge of this zone ($4,777 \text{ m}^3 \text{ s}^{-1}$ or 68.5% of BEL annual discharge) than to its weaker mean SM concentration (13 mg l^{-1} , usual feature value in the central Amazonian plain).

As we noted in § IV. I, after leaving Ecuador, the Napo River receives the contribution of two important tributaries in terms of discharges and TSS, the Aguarico and Curaray rivers. For example, during the scientific field trip of october 2004 (Tab. 4), each one of them presented a discharge of about $800 \text{ m}^3 \text{ s}^{-1}$ that correspond both to 36% of Napo River discharge at BEL, with respectively 110 and 60 mg l^{-1} of TSS. The Curaray catchment and the rest of the drained area of the Napo River in Peru correspond to a relatively flat area covered with rain forest. The other little tributaries totalised less that $500 \text{ m}^3 \text{ s}^{-1}$ or 12% of Napor River at BEL, with weak TSS concentrations (17 to 78 mg l^{-1}).

These considerations prove that the sediment load which exit the Ecuador country at ROC are totally transported to the mouth of the river near BEL with an increase due to the input coming essentially from the two main tributaries, the Aguarico and Curaray rivers. The Curaray River can present relatively high TSS concentrations. On june 2002, at the border between Ecuador and Peru at an altitud of 200 masl (similar with the one of ROC), Laraque *et al.*, (2002) measured a discharge of about $800 \text{ m}^3 \text{ s}^{-1}$ for about 200 mg l^{-1} of TSS during a flood in Pavacachy on the Curaray River. These values are similar with the annual mean ones encountered on the Napo River at FDO at 260 masl and correspond to TSS fluxes of 13,800 ton

d⁻¹.

Finally, in the last 100 km, until its confluence with Marañón River, the Napo flows with a weak slope close to 5 cm km⁻¹ (Fig. 3), and no erosional power.

IV. 3 – Impact of volcanism and geodynamic on TSS

To better illustrate the lack of interannual reproductivity of the observed sediment transport dynamic, cumulative curves of monthly TSS concentration, discharge and suspended sediment yield have been analysed (Fig. 6). Within a stable basin, the accumulation of suspended sediment yield are similar from one year to the next and regularly increases with some cyclic changes corresponding to alternations of wet and dry seasons (Laraque et al., 1995), like is the case in FDO basin (Fig. 6b). In SEB (Fig. 6a) basin, contrary to the discharges, the TSS accumulations (concentrations and suspended sediment yield) show a sharp rise from November, 2002, affecting ROC station located downstream (Fig. 6c). This slight drop returns to normality 6 months later in May, 2003 (Fig. 6a). This abnormality may be explained due to the El Reventador volcano eruption located in the upstream part of the Coca basin. In fact, on 11/03/2002 the volcano expelled in a few hours an ash cloud of 15 km of height estimated at 200 million tons of material (Eissen and Le Pennec, 2002). Studies made on dissolved elements from the Coca River to SEB, Gondran (2004) and Bernal *et al.*, (2007) show that El Reventador ash leaching might have lasted many months. It was increased by the glowing ash clouds and post - eruptive lahars, which are frequent in that wet and sloping region. The latter might be the origin of the succession of the 3 important peaks of TSS to SEB in 2003, which do not always correspond to the strong floods (Fig. 7a & b). On the contrary of those taking place in May 2001 and in May 2004, which have also provoked curve jolts shown in Figure 6a, correspond to relatively important floods. However, according to the periodic report's "Instituto de Geofísica - Ecuador" (Bernal, personal

communication), this last one (May 2004) was a secondary lahar produced by important precipitations which have remobilized the ash of El Reventador volcano eruption in 2002. It is quite likely that a great part of the volcanic material carried by the Coca River was deposited downstream from the SEB, an area marked by the rupture of the Andean – Piedmont river slope. Downstream, a number of sandbanks begin to appear and become more frequent along the entire Napo River until ROC station (Fig. 3). Thus, the main erosion process in the downstream basin of the Ecuadorian mountainside temporarily inverted itself in 2003 under the effect of a major and isolated volcanic activity. It then returned to « normality » with a restart of the erosion process showed by an upstream – downstream increase of TSS concentrations.

The hypothesis of an unstable basin, where the sediment budget is disturbed due to catastrophic events, like earthquakes, volcanic eruptions and related eruptions events, is supported by the satellite image analyses (Fig. 8), which allows evaluating the variation percentage of sand on the Napo River. In these areas, the Napo River has only one thalweg but with the presence of multiple moveable sand bars. In the area I (FDO-X, Fig. 3), which is closer to foothills, it is characterized by roughly constant percentages in sand, water and vegetated islands, during the 3 years analysed (1995, 1996, and 2002); with a slight tendency to decline in the sand (drop of 5%, Fig 8). This fact suggests the presence of an erosion process in the nearest Amazonian plain. The area II (X-ROC, Fig. 3) is characterized by a stronger variation of the percentages in sand, water and vegetation for the 3 years studied (1986, 1989, and 2002; Fig. 8). In the case of the sand, for example, in 1986 the river has an average of 15%, after three years it almost doubled the sand amount (27%), and finally after 13 years the percentage diminishes to half (8%). This variability in the sediment supply from the watershed can be explained by active tectonics and seismicity characteristics of the zone, which in march 1987 provoked two major earthquakes (M_s 6.1 and 6.9), near Reventador

Volcano. These quakes were the triggering of one of the most catastrophic landslides of the 20th century – worldwide (Schuster, 1996). About 600 mm of rain fell in the region in the months preceding the earthquakes (Crespo, 1987), and thus land sliding mainly in saturated, residual soils on steep slopes. The slope failures commonly started as thin slides, which rapidly turned into debris avalanches and debris flow that flowed down the slopes into the tributaries and on into major rivers (Schuster, 1996). Estimates of the total volume of earthquake-induced mass wastage ranged from 75-110 million m³ (Crespo et al., 1987; Okusa et al., 1989).

The important geodynamic activity of this part of the Andes, marked by an exceptional density of volcanoes, is directly affected by the tectonic slip of the Nazca Plate under the South American Continent, which caused an uplift of the Eastern Andean foreland. In Napo basin, according to Christophoul *et al.*, (2002a & b), this previous process has shown an effect in the eastern side of the Ecuadorian Cordillera (at the Miocene-Pliocene limite), by the uplift of ancient alluvial plains originating the Pastaza Megafan. This is evidenced geomorphically, by river incisions and presence of old elevated fluvial terraces upstream FDO, dating from the Neocene period (Burgos *et al.*, 2004, Bès de Berc *et al.*, 2005). The courses of Napo had then progressively deviated towards the northern border. The influence of that morphologic structure and the relatively high slope for a drainage of alluvial plain formations, is the cause for the regular tracking and little meandering of Napo at the entrance of the Amazonian plains. This maintains once again an important hydraulic dynamic.

The observation of the variability of sediment transfer process confirms the preliminary interpretations of Laraque *et al.*, (2004) announcing an erosive process in the region of the Andes which present the most important tectonical dynamic characterized by intensive volcanic and seismic activity.

IV. 5 - Regional comparisons

The results found on the Peruvian Napo basin were unexpected because its slope indice (11 cm km^{-1}) is similar to the one encountered at the foot of the Bolivian Andes (slope of 8 cm km^{-1}). In this last region, the river (Ex. of Beni/Mamoré River) meanders with low water velocity through an area undergoing subsidence, where they deposit of more that half of their suspended sediment yield (Guyot, 1993, Guyot, *et al.*, 1996, Baby *et al.*, in press). From the analysis of the ^{210}Pb activity of sedimentary cores, Aalto, *et al.*, (2003) evaluated a rate of sedimentation at 100 Mt.yr^{-1} in the floodplains of Beni foreland basin which covers around $70,000 \text{ km}^2$.

More downstream, in central depression of the Amazon basin in Brazil, Laraque *et al.*, (2005), quantifies important sedimentation processes of $0.4 \cdot 10^6 \text{ t yr}^{-1} \text{ km}^{-1}$ in a region where the fluvial slope varies between 3 and 0.3 cm km^{-1} . The nature of these processes was already shown by Meade *et al.* (1985) and Dunne *et al.*, (1998), and illustrates fluvial transport marked by cyclic storage and remobilization stage of sediment, before reaching the Atlantic ocean.

V. CONCLUSION

This paper based on collected field data and geomorphological interpretation, enables the proposal of a current sediment transport model for the Andean Piedmont and plains of the Napo basin in Ecuador and Peru. The sediment transfer along the river axis of the Napo River have been studied and quantified: in Ecuadorian part of the Napo basin during four hydrologic cycles from 2001 to 2005; in its Peruvian part during one hydrologic cycle from 2004 to 2005; and during a field scientific trip in October 2004. Regarding the annual budget, erosion processes dominate, with an important year-to-year variability. For the 4 studied cycles in

Ecuadorian part, they are 16, 5, 6.6 et $14.5 \cdot 10^6 \text{ ton yr}^{-1}$, and 19.8 ton yr^{-1} for the only one cycle studied in Peruvian part. These values show that during the studied period, approximately 47% of the TSS flux exited from the ecuadorian basin of Napo stemming from the erosion of andean foreland ($1,160 \text{ t km}^{-2} \text{ year}^{-1}$) located between FDO+SEB and ROC station.

Then, the Napo River enters the flat Amazon Plain where a transport process of TSS dominated, increase by the input coming essentially from the two tributary of Aguarico and Curaray River. The last one tributary drains the Pastaza Megafan.

A discussion based on the comparing of these results with the geomorphologic and geodynamic characteristics of this Andean region, enables us to understand the origin and variability of the mass transfer processes. In the amazonian plain of Ecuador, the basin presents a relatively high slope (35 cm km^{-1}). It is made up of very fine fluvial sediments. Taking in consideration the active and current uplift of the western foreland belt of the Amazon plain, these results confirm those other studies on geodynamics (conducted over a longer period). Since the end of Neogene and during Quaternary period, under the influence of the Andes uplifts, the foreland basin has presented erosion processes of the old sediments, which were deposited at a time when the slopes were less steep. Further downstream in the peruvian part of the Napo Basin, the index slope diminish significantly ($10\text{-}20 \text{ cm km}^{-1}$) between ROC and BEL and the sediment contribution of the intermediated drained area corresponds to 42.5% of the Napo River sediment yield at BEL. Here the sediment budget stays always high, with an excess of 19.8 ton yr^{-1} for the only one year studied. This positive budget probably comes from the two main tributaries of Aguarico and Curaray River, unfortunately poorly studied, and maybe also from the continuity of an erosion and/or transport process of TSS.

The opposite condition has been observed further South in the Madeira foreland basin (Guyot *et al.* 1996), where half of the sediment suspended load of the Beni/Mamoré River settles

between the Andean Piedmont of Bolivia and the Brazilian border. This basin presents a weaker slope of 8 cm km^{-1} in a region still marked by an active flexural subsidence. Further east, in the Central Amazon Flood Plain in Brazil, sedimentary processes are also observed with weaker slopes ($3 \text{ to } 0.3 \text{ cm km}^{-1}$, Laraque *et al.*, 2005).

The combination of geodynamic phenomena (tectonism, volcanoes, earthquakes and landslides) common in this region, favours large sediment transfers, not only at present but also in the past. This is confirmed by the presence of incised terraces, mainly formed by volcanic materials (Bès de Berc, 2003).

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Legend: continuous line = limit of Amazon basin; dot line = limit between Andes and Amazonian plain

Fig. 2 Mean rainfall regime in the Napo basin (1964-1998)

Fig. 3 Longitudinal profile of Napo River is represented by a grey line; while slope curve is represented by a black line. The proportion of sand, water and vegetated island in the Napo River is presented in two reaches FDO – X (top) and X – ROC (down). X's coordinate: $0^{\circ} 26' 39.16'' \text{ S} / 76^{\circ} 14' 51.05'' \text{ W}$

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legend : TSS histogram in 10^3 ton d^{-1} (positive values = erosion process, negative values = sedimentation process); monthly discharge curve (Q) in $\text{m}^3 \text{ s}^{-1}$; abscissas in month-year.

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Tab. 3 Annual Discharges (Q in $\text{m}^3 \text{s}^{-1}$), TSS load (in mg l^{-1} , ton d^{-1} , ton yr^{-1} , $\text{ton km}^2 \text{yr}^{-1}$), at the Gauging Station studied.

Tab. 4 Data from October 2004 field sampling crossing along Napo River: latitude, longitude, distance from the Napo mouth (km) and altitude (m), TSS load concentration (mg l^{-1}), measured discharges (Q in $\text{m}^3 \text{s}^{-1}$), TSS fluxes & annual upstream-downstream budget (TSS in ton d^{-1}) between the gauging stations studied. For sampling code, see fig. 1.

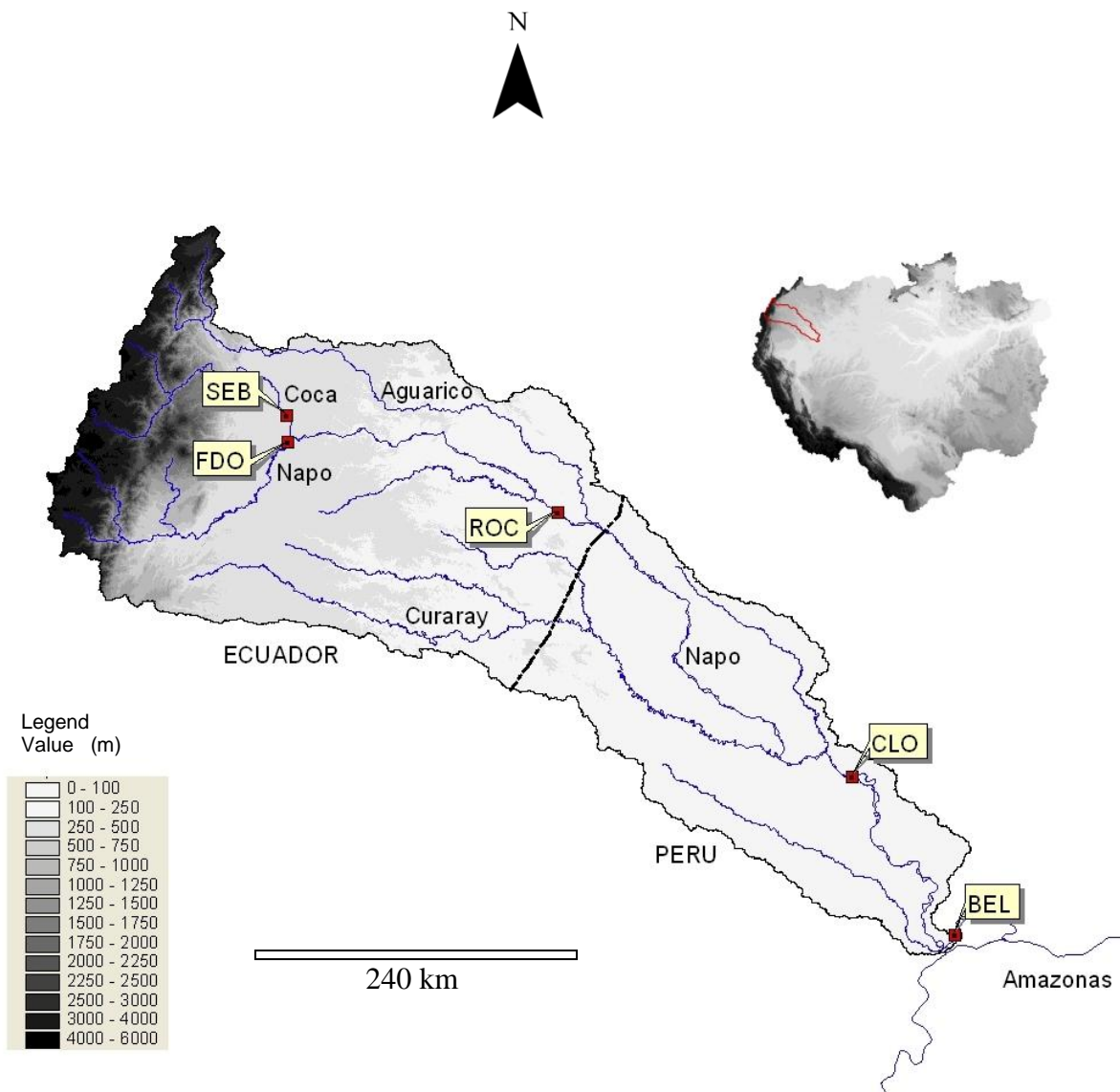


Figure 1 – Location of the Napo basin and its reference stations (SEB = Coca river at San Sebastian, FDO = Napo river at Francisco de Orellana, ROCA = Napo river at Nuevo Rocafuerte, CLO = Napo river at Santa Clotilde, BEL = Napo river at Bellavista). Legend: continuous line = limit of Amazon basin; dot line = limit between Andes and Amazonian plain.

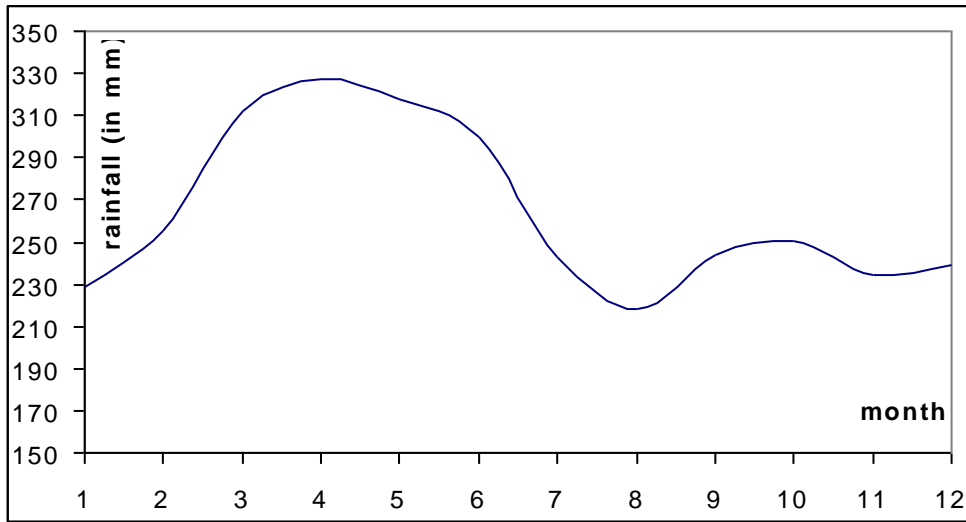


Figure 2 – Mean annual rainfall regime in the Napo basin (1964-1998)

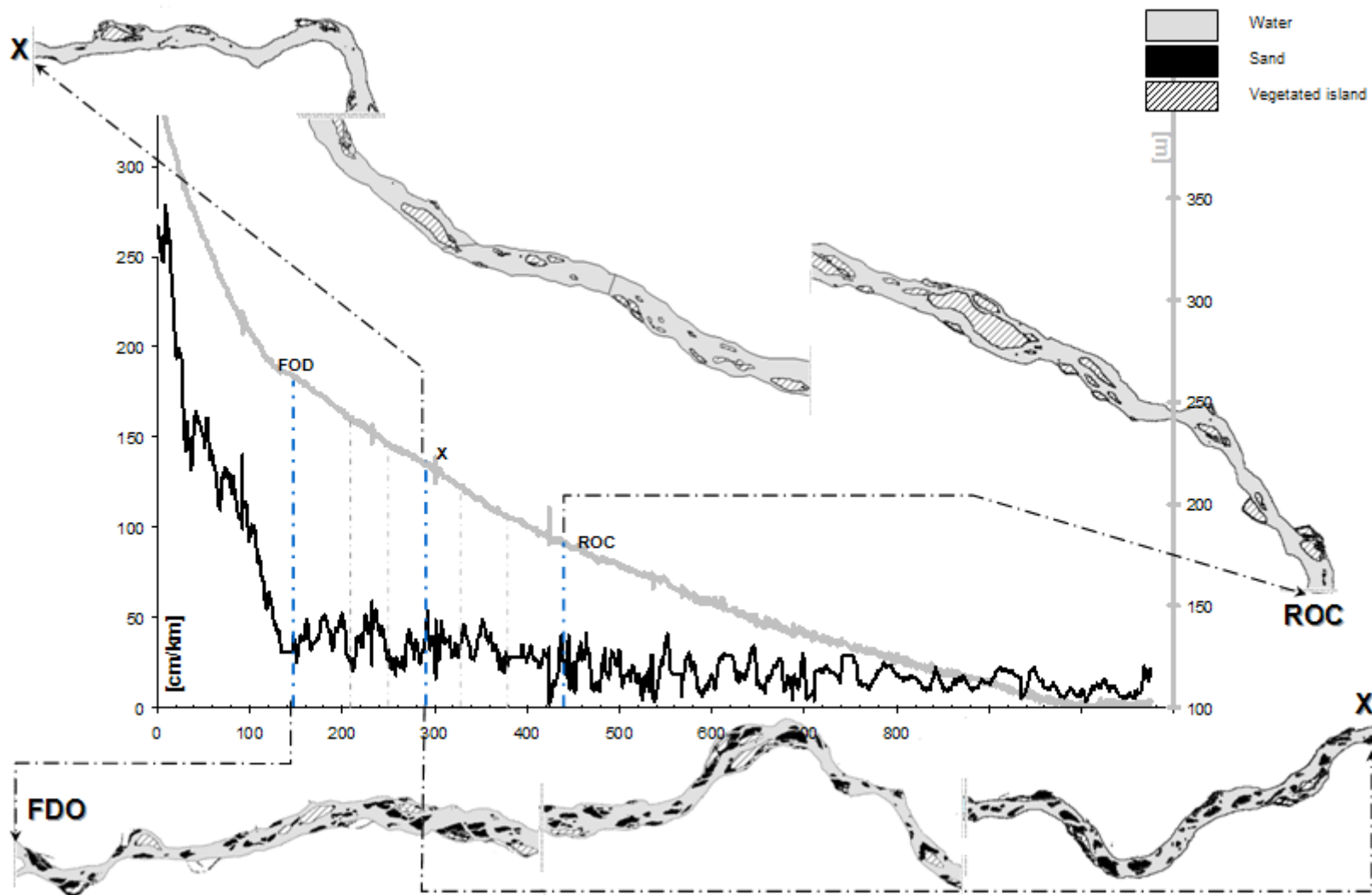


Figure 3 - Morphology of Napo River in the Amazonian Plain. Longitudinal profile is represented by a gray line; while slope curve is represented by a black line; X = (0° 26' 39.15" S; 76° 14' 51.05"W).

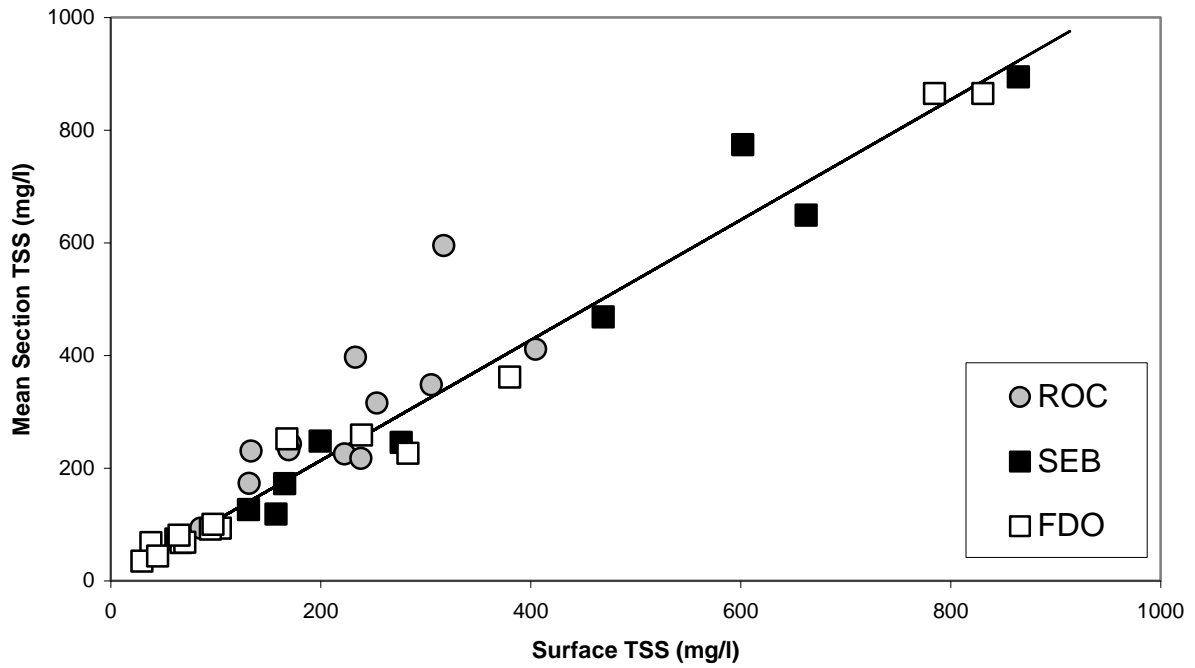


Figure 4 – Relation between superficial and mean TSS concentration in the section (in mg l^{-1}) at the 3 gauging station studied (ROC, SEB, FDO).

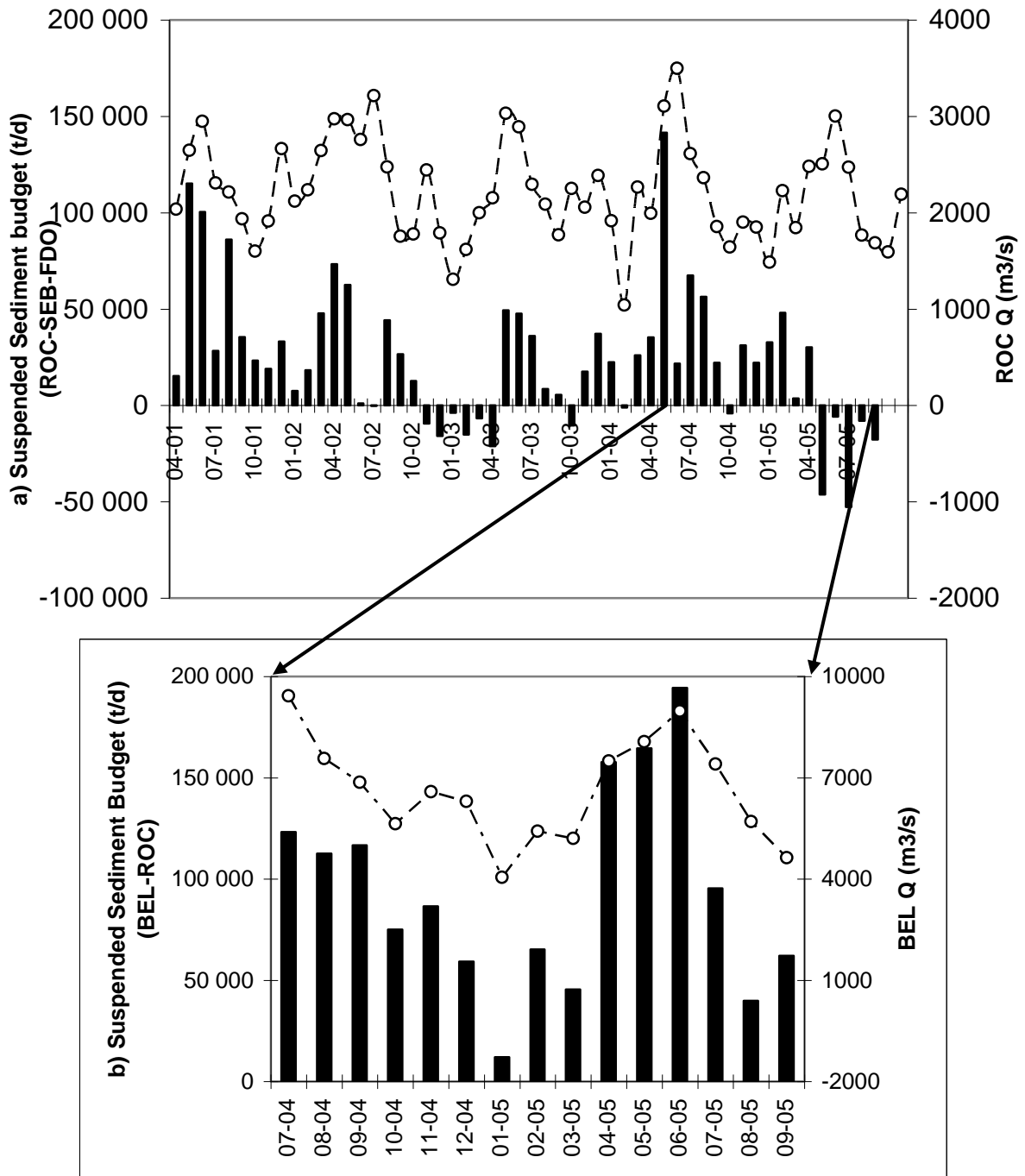


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 legend : TSS histogram in 10^3 ton d^{-1} (positive values = erosion process, negative values = sedimentation process); monthly discharge curve (Q) in $\text{m}^3 \text{ s}^{-1}$; abscissas in month-year.

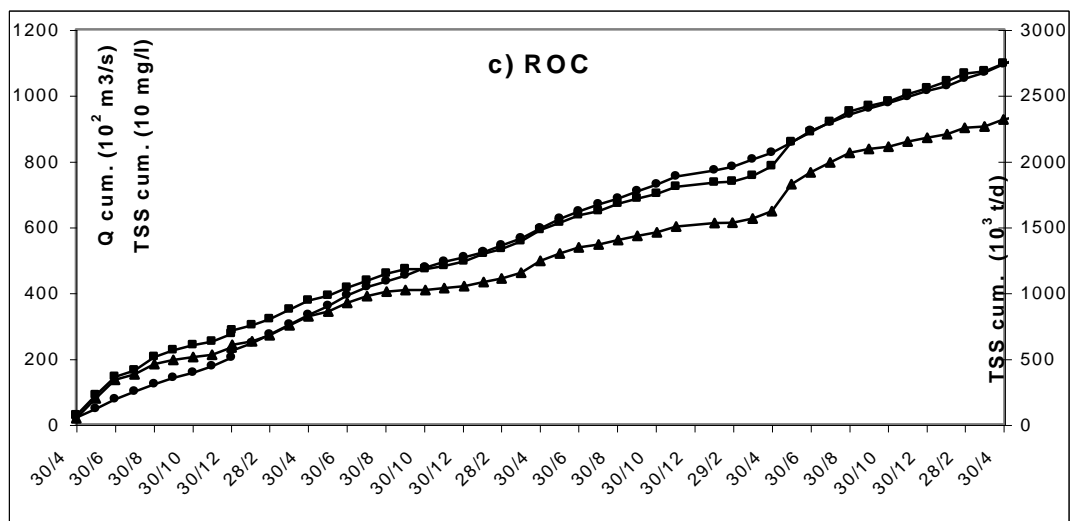
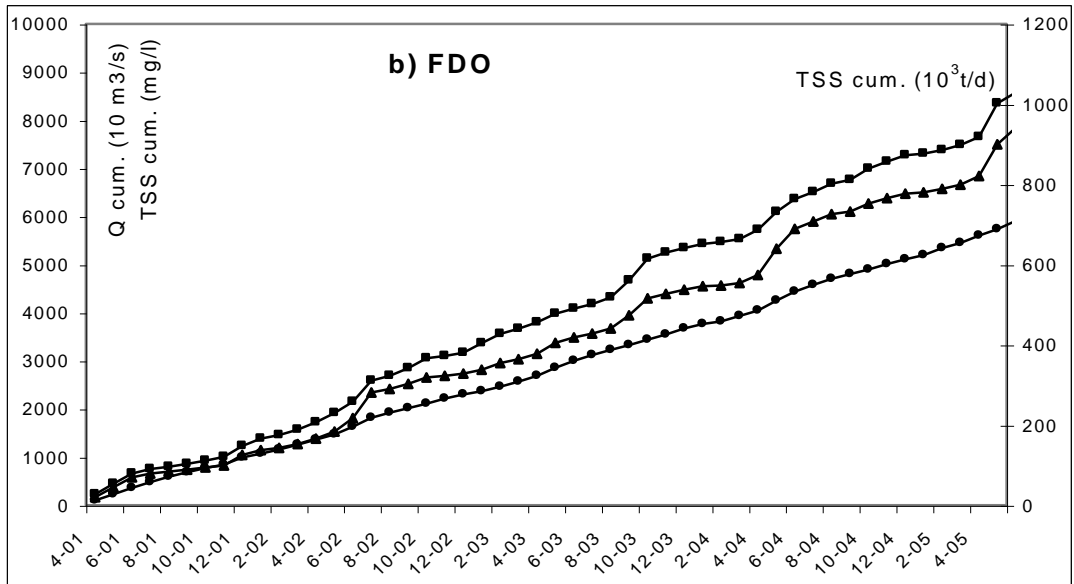
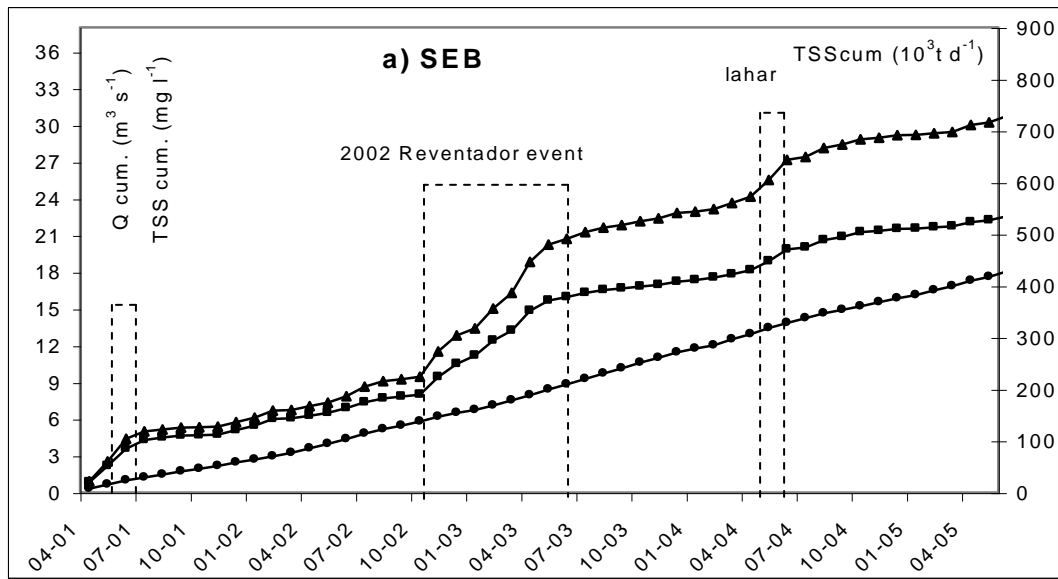


Figure 6 - Cumulative (Cum) curve of Q (in $m^3 s^{-1}$ - circle), TSS concentration (in $mg l^{-1}$ - square) and TSS fluxes (in $10^3 ton d^{-1}$ - triangle) during the period studied at SEB.

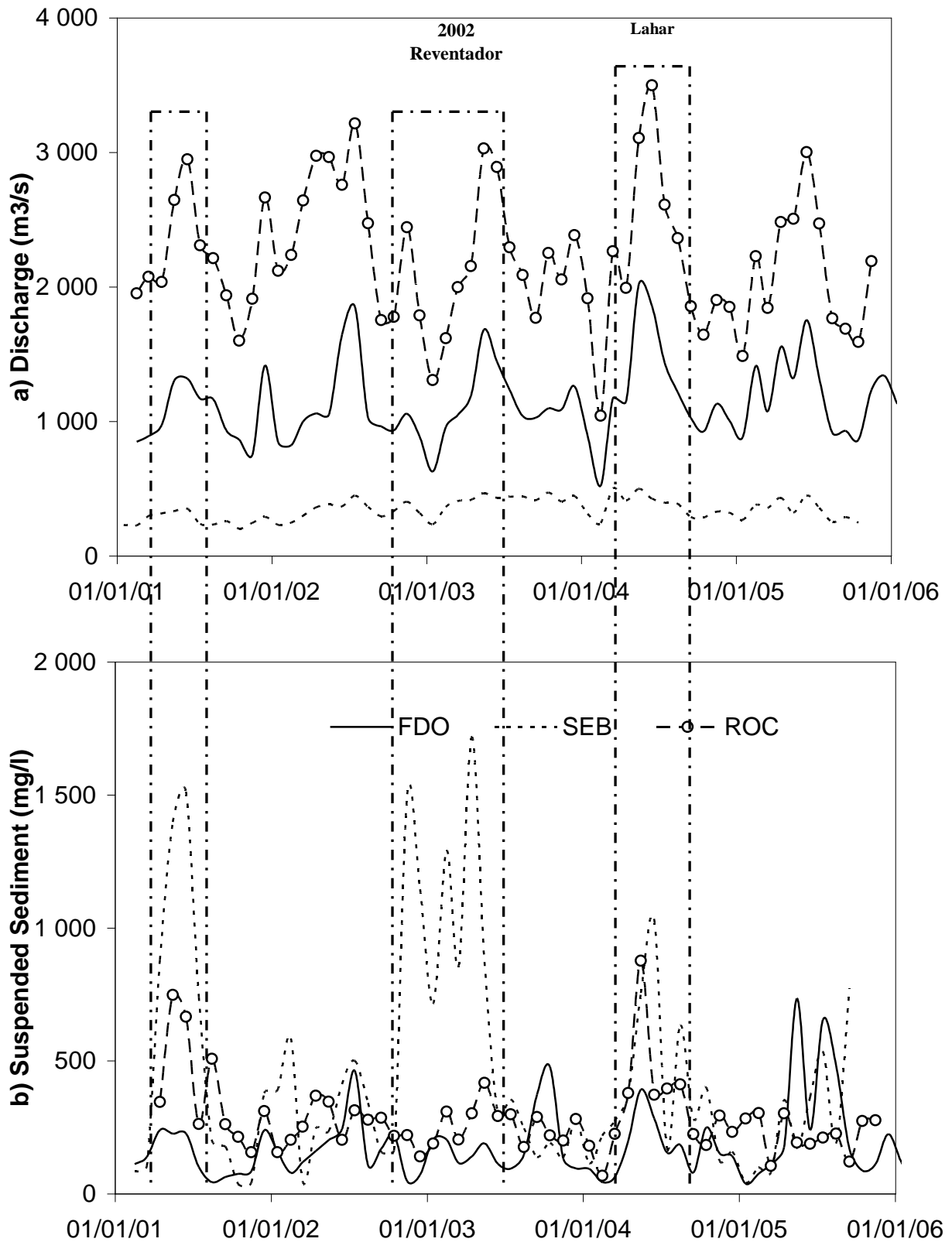


Figure 7 Monthly **a)** hydrographs (in $\text{m}^3 \text{s}^{-1}$) and **b)** TSS concentration (in mg l^{-1}) at the 3 gauging station studied.

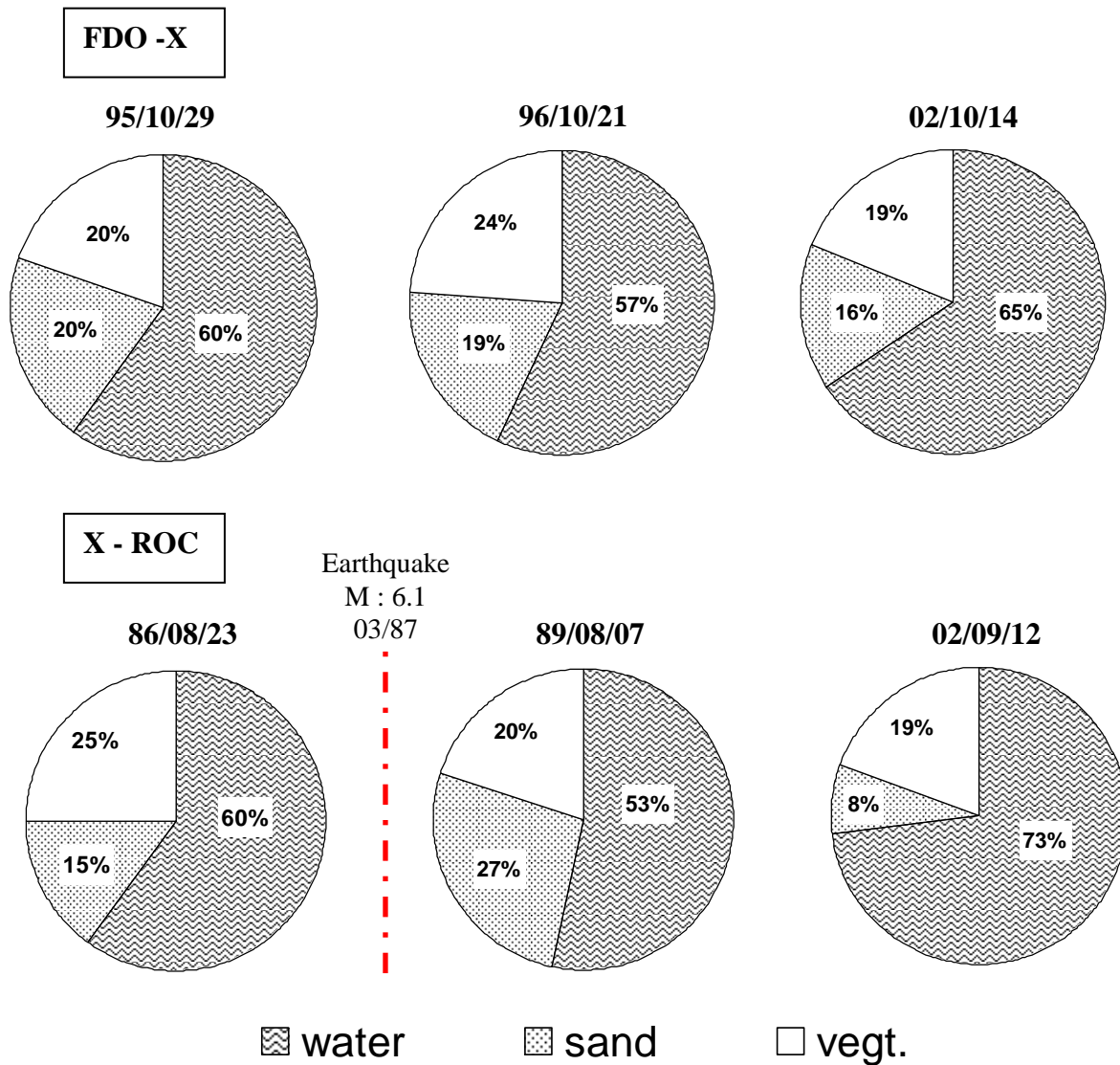


Figure 8 – Spatial variation of percentage in sand, water and vegetated island for two reach: FDO-X and X-ROC (see Fig. 3)

Table 1 - Presentation of the drainage basin studied with location of gauging stations.

Code Station	Gauging station	River	Latitude (degree)	Longitude (degree)	Altitud Basin area		Discharge		Distance from mouth (km)	Study period	slope ⁽³⁾ (cm km ⁻¹)	Section for slope data
					⁽¹⁾ (m)	⁽²⁾ (km ²)	$\text{m}^3 \text{s}^{-1}$	$\text{l s}^{-1} \text{km}^2$				
SEB	San Sebastian	Coca	-0.343	-77.007	290	5,270	350	66	870	01/2001-10/2005	38.8	ROC-SEB
FDO	Francisco de Orellana	Napo	-0.473	-76.983	262	12,440	1,130	91	850	01/2001-11/2005	30.4	ROC-FDO
ROC	Rocafuerte	Napo	-0.917	-75.396	189	26,860	2,210	82	610	01/2001-12/2005	17.0	CLO-ROC
CLO	Santa Clotilde	Napo	-2.486	-73.674	129	85,770	5,700	66	280	01/2004-12/2005	12.0	BEL-CLO
BEL	Bellavista	Napo	-3.478	-73.076	106	100,520	6,270	62	70	01/2004-02/2006		

(1) = DGPS data from Fraizy *et al.*, (2005).

(2) = from SRTM extraction.

(3) = Bourrel, L., *et al.*, (2007).

Table 2 - Interannual Values : minimum, mean and maximum of daily and monthly discharge ($\text{m}^3 \text{s}^{-1}$), 10 days TSS load and monthly TSS load (mg l^{-1}), daily and monthly TSS yield (Ton d^{-1}), at the gauging stations studied

		SEB	FDO	ROC	BEL
Daily Discharge ($\text{m}^3 \text{s}^{-1}$)	Mean	345	1,134	2,207	6,265
	Max	724	5,055	5,087	12,360
	Min	103	381	954	1,285
	Max / Min	7	13	5	10
	Values number	1,765	1,793	1,801	789
Monthly Discharge ($\text{m}^3 \text{s}^{-1}$)	Mean	344	1,140	2,217	6,267
	Max	502	2,020	3,499	11,090
	Min	202	523	1,043	1,426
	Max / Min	2	4	3	8
	Values number	58	58	58	25
10 day TSS load (mg l^{-1})	Mean	431	161	218	125
	Max	2,806	956	1,403	341
	Min	10	2	5	32
	Max / Min	281	478	281	11
	Values number	167	175	165	76
Monthly TSS load (mg l^{-1})	Mean	467	190	291	235
	Max	1,766	728	912	348
	Min	37	37	69	139
	Max / Min	48	20	13	3
	Values number	56	57	56	19
Daily TSS yield (ton d^{-1})	Mean	14,235	19,982	58,861	133,395
	Max	116,330	355,510	551,590	424,410
	Min	168	229	3,385	36,086
	Max / Min	694	1,552	163	12
	Values number	1,751	1,773	1,739	618
Monthly TSS yield (ton d^{-1})	Mean	14,607	20,146	59,332	132,762
	Max	63,447	83,259	244,780	246,440
	Min	639	2,059	6,222	48,476
	Max / Min	99	40	39	5
	Values number	56	57	56	19

Table 3 - Annual Discharges (Q in m³ s⁻¹), TSS load (in mg l⁻¹, ton d⁻¹, ton yr⁻¹, ton km² yr⁻¹), at the gauging station studied

	Studied Period	SEB	FDO	ROC	BEL	SEB+FDO	ROC-FDO-SEB	BEL-ROC
Discharge (m ³ s ⁻¹)	04/2001-03/2002	270	1,047	2,272		1,317	955	
	04/2002-03/2003	357	1,093	2,256		1,449	807	
	04/2003-03/2004	416	1,138	2,179		1,554	624	
	04/2004-03/2005	363	1,261	2,199	6 976	1,624	575	4,777
	<i>Mean value</i>	351	1,135	2,227	6 976	1,486	740	4,777
TSS (mg l ⁻¹)	04/2001-03/2002	545	140	344		180	53	
	04/2002-03/2003	635	179	260		230	20	
	04/2003-03/2004	415	168	251		178	34	
	04/2004-03/2005	352	177	352	224	179	80	13
	<i>Mean value</i>	487	166	302	224	192	45	13
TSS (ton d ⁻¹)	04/2001-03/2002	14,050	13,404	71,571		27,454	44,117	
	04/2002-03/2003	19,890	18,617	52,590		38,507	14,083	
	04/2003-03/2004	15,318	16,708	50,158		32,026	18,132	
	04/2004-03/2005	12,081	21,519	73,422	127 562	33,600	39,822	54,140
	<i>Mean value</i>	15,335	17,562	61,935	127 562	32,897	29,038	54,140
TSS (x 10 ⁶ ton yr ⁻¹)	04/2001-03/2002	5,1	4,9	26,1		10,0	16,1	
	04/2002-03/2003	7,3	6,8	19,2		14,1	5,1	
	04/2003-03/2004	5,6	6,1	18,3		11,7	6,6	
	04/2004-03/2005	4,4	7,9	26,8	46,6	12,3	14,5	19,8
	<i>Mean value</i>	5,6	6,4	22,6	46,6	12,0	10,6	19,8
TSS (ton km ² yr ⁻¹)	04/2001-03/2002	973	393	973		1,366	1,760	
	04/2002-03/2003	1,378	546	715		1,924	562	
	04/2003-03/2004	1,061	490	682		1,551	723	
	04/2004-03/2005	837	631	998	463	1,468	1,589	268
	<i>Mean value</i>	1,062	515	842	463	1,577	1,158	268

Table 4 - Data from october 2004 field sampling crossing along Napo river : latitude, longitude, distance from the Napo mouth (km) and altitude (m), TSS load concentration (mg l⁻¹), measured discharges (Q in m³ s⁻¹), TSS fluxes & anual upstream-downstream budget (TSS in Ton d⁻¹) between the gauging stations studied. For sampling code, see Fig. 1.

Sampling code	Date	River	Station	Latitude	Longitude	distance km	altitude m	TSS mg l ⁻¹	measured Q m ³ s ⁻¹	TSS Ton d ⁻¹	TSS budget* Ton d ⁻¹
NA_01	18/10/2004	Misahualli	A.J. Napo	-1.03225	-77.6642			13.4			
NA_02	18/10/2004	Napo	A.J. Misahualli	-1.03664	-77.66713	994	404	28.8			
NA_03	18/10/2004	Arajuno	AJ Napo	-1.05060	-77.52500			30			
NA_04	19/10/2004	Suno	AJ Napo	-0.69470	-77.13600			6.8			
NA_05	19/10/2004	Payamino	AJ Napo	-0.47290	-76.99970			22.4	40	77	
NA_06	19/10/2004	Napo	FDO	-0.47396	-76.98009	850	262	44.8	256	991	991
NA_07	19/10/2004	Coca	SEB	-0.45377	-76.98652			33.2	620	1,778	
NA_08	20/10/2004	Napo	Itaya	-0.43200	-76.53600	789	242	74.4	1,175	7,553	6,562
NA_09	21/10/2004	Panayacu	AJ Napo	-0.43154	-76.10489			107.6	26	242	
NA_10	21/10/2004	Napo	Panacocha	-0.46941	-76.06965	718	223	86	1,200	8,916	1,363
NA_13	22/10/2004	Tiputini	AJ Napo	-0.82886	-75.55802			24.3		1,896	
NA_14	22/10/2004	Napo	ROCA	-0.87787	-75.45855	608	189	132	1,210	13,800	3,006
NA_15	23/10/2004	Yasuni	AJ Napo	-0.94500	-75.39600			78.4	135	914	
NA_16	23/10/2004	Aguarico	AJ Napo	-0.96200	-75.19800			109.8	793	7,522	
NA_17	25/10/2004	Aushiri	AJ Napo	-1.78600	-74.71400			18.2	130	204	
NA_18	26/10/2004	Napo	D.J.Aushiri	-1.85900	-74.64800	425	154	127.0	2,680	29,407	13,334
NA_19	27/10/2004	Curaray	AJ Napo	-2.34997	-74.09310			60.8	820	4,307	
NA_20	28/10/2004	Tambayacu	AJ Napo	-2.46400	-73.61200			26.2			
NA_21	28/10/2004	Napo	CLO	-2.48181	-73.68194	253	128	80.8	3,800	26,528	11,101
NA_22	28/10/2004	Tacsha	AJ Napo	-2.78825	-73.54963			17.2	46	68	
NA_23	29/10/2004	Yanayacu	AJ Napo	-3.08080	-73.15320			18.4			
NA_24	30/10/2004	Mazan	AJ Napo	-3.50450	-73.10850	73	107	27.8	117	281	
NA_25	30/10/2004	Napo	BEL	-3.47826	-73.07606	70	106	199.2	4,378	75,349	9,158

