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# **Suspended sediment and dissolved load budgets of two Amazonian rivers from the Guiana shield: Maroni river at Langa Tabiki and Oyapock river at Saut Maripa (French Guiana).**

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## **Abstract**

This work presents the data collected within the framework of an Observatory of research in Environment on the Amazonian basin – the ORE HYBAM. It relates to the dissolved and solid loads of the two main rivers of French Guiana, the Maroni and Oyapock rivers, running on the Guiana shield and draining respectively 64 230 and 24 630 km<sup>2</sup>. The low coefficient of variation of the average annual flows of the two rivers indicates an inter-annual hydrological stability probably related to the immediate vicinity of the Atlantic Ocean. The sedimentary load is mainly composed of quartz and kaolinite. The total suspended solid concentrations are among the world's lowest values; they range from 1 to 130 mg l<sup>-1</sup> during the hydrological cycle, with averages of 22 and 12 mg l<sup>-1</sup>, for the Maroni and Oyapock rivers respectively. The seasonal variability of these values is significantly higher than that of hydrologic flows, but without simple relationship with the discharge. Water chemical composition of the two rivers indicates a very weak mineralization, very similar to that found in the Amazonian rivers running on the Brazilian and Guianese shields and in the Congo River and its tributaries in the Central African Shield. Seasonal variations are observed in both basins; they correspond to higher concentrations during low water stage (from October to February) and to more diluted water during the flood, from April to July. A signature enriched in Cl<sup>-</sup> is present at the Saut Maripa station on the Oyapock River indicating a more marked influence of the trade winds in this basin. The computation of atmospheric contributions to ions budget indicated a weak contribution for Ca<sup>2+</sup> and Mg<sup>2+</sup>, which originates mainly from water-rock interactions in both stations, while more than half of Na<sup>+</sup> is derived from atmospheric inputs.

## **Key words.**

Hydrology, Suspended Sediment, Geochemistry, French Guiana, Amazon Region

## INTRODUCTION

The yield of suspended and dissolved materials is of fundamental interest for the computation of weathering and erosion rates. As runoff varies seasonally in tropical watersheds, regular sampling over some annual hydrological cycles should provide the best basis for estimating element budgets.

In French Guiana, the existing hydrological studies, conducted primarily by IRD (ex ORSTOM), focused on the recognition of potential sites for hydropower dam (Roche, 1978a; Fritsch, 1982, 1985), on the study of estuarine systems (Berthois & Hoorelbeck, 1964; Roche et al., 1974; Lointier, 1990; Vauchel, 1998), and on the effect of the rainforest clearing on soil erosion (Roche, 1978b, 1982; Fritsch, 1983, 1987, 1992, Fritsch & Sarrailh, 1986; Bariac et al., 1996). To date, major rivers in the region have been only preliminarily studied (Hiez, 1955; Hiez & Dubreuil, 1963) while the data on the dissolved and particulate flows essential for assessment erosion and weathering rates of these major tropical basins remains fragmented (Humbel, 1989; Lointier & Roche, 1989; Lointier 1995). A recent report of Monfort and Ruf (2005) assessed the main characteristics of the hydrological regimes of the rivers of French Guiana.

This paper reports the results of four years of a monthly survey in two forested tropical watersheds of French Guiana. The survey was performed in the frame of the Environment Research Observatory HYBAM (Geodynamical, hydrological and biogeochemical controls of erosion/alteration and material transport in the Amazon basin, <http://www.ore-hybam.org>). This observatory comprises 16 gauging and sampling stations distributed along the Amazon Basin and in the Andean piedmont; its main objective is the acquisition of long term data series to evaluate the climatic and human impacts. The two stations presented here are situated in French Guiana. One is located on the Maroni River, at Langa Tabiki, and the other one is situated on the Oyapock River at Saut Maripa. Both rivers correspond to the boundaries of French Guiana with a neighbouring country: in the West, the Maroni River follows the limit of Suriname and, in the East, the Oyapock River determines the border with Brazil. Sampling started in April, 2004 in both sites and was performed monthly.

## STUDIED AREA

The studied area corresponds to the drainage basins of the two main rivers of French Guiana: the Maroni and the Oyapock rivers (Figure 1 and Table 1).

The Maroni River is the largest river of this region. This river comes from the western Tumuc Humac hills located in Suriname, where it is locally named Itany River. Downstream the confluence with Inini River, its name changes to Lawa River, and finally downstream Tapanahony River confluence, it becomes the Maroni River. At the Langa Tabiki gauging station (**LT**), the Maroni river drains a 64 230 km<sup>2</sup> area, 43% located in French Guiana and 57% in Suriname (Hiez & Dubreuil, 1963).

The Oyapock River delimits the French Guiana border with Brazil. The river comes from the eastern Tumuc Humac hills, and flows downstream to the NE, despite more than 100 “sauts” (small waterfalls). With its main tributary (the Camopi River), the Oyapock River at the Saut Maripa gauging station (**SM**), drains a 24 630 km<sup>2</sup> area, 49% located in French Guiana and 51% in Brazil (Hiez & Dubreuil, 1963).

The Guiana region can be divided into three distinct morphological units (Barret, 2002):

- i. the low-land zone corresponds to the littoral plain with an altitude lower than 30 m a.m.s.l.; its extension varies from 5 to 40 km wide,
- ii. the “Central Massif” extends over  $\pm$  100 km and corresponds to the Green Belt domain of the Guiana Shield. It has a general WNW/ESE orientation and comprises deeply eroded massifs with an altitude ranging from 400 to 850 m a.m.s.l.,
- iii. the southern peneplain consists of convex hills less than 250 m high. Many “inselbergs” are present in this area, contrasting with the monotone landscape of the low hills.

The studied area is situated in the northern and eastern parts of the Guiana Shield, which extends over the north of the Brazilian Amazon Basin, the West of Colombia and Venezuela and covers the French Guiana, the Surinam and the Guiana. In the Northeast of Brazil and south of French Guiana, it consists of igneous rocks (granite, gabbro, diorite) of Paleo-Proterozoic ages. These rocks are locally metamorphized and amphibolitized. The centre and north of French Guiana is composed of greenstone belts oceanic rocks and has been affected by the Trans-Amazonian tectonic and metamorphic event (Gibbs & Barron, 1993; Gibbs et al., 1986)

Intense weathering of the bedrocks since the Cretaceous in this very stable area has induced the development of a thick layer of residual alterite masking the substratum (Edmond et al., 1995 ; Negrel & Lachassagne, 2000). However, a recent (Quaternary) uplift is observed, probably associated with the peripheral bulge engendered by sediment loads of the Amazon fan (Driscoll & Karner, 1994). It could explain the drainage of unweathered material in some parts of the basins.

The climate of this region is governed by the displacements of the Intertropical Convergence Zone (ITCZ) resulting from the convergence of the Trade Winds blowing from the high-pressure areas in the subtropical latitudes towards the low-pressure areas around the equator. Four seasons can be distinguished (Groussin, 2002). From mid-November to mid-February, the ITCZ moves southwards and passes over the land, giving rise to a little rain season. In March, the ITCZ reaches its maximum southern position (between 1°S and 2°N); this period corresponds to a diminution of average rainfall. From April on, the ITCZ moves slowly northward and consequently an intense rainy season develops until July. From August to the end of October, the ITCZ is in a more northern position and a dry season occurs. As a consequence of the direct relationship between rainfall and the displacement of the ITCZ, a clear east-west trend appears in average total annual rainfall: it ranges from 3500–4 000 mm in the east to 1700–2 000 mm in the west. The mean annual rainfall for Langa Tabiki and Saut Maripa drainage basins is respectively 2 520 mm and 2 680 mm. The mean annual temperature is 26.5°C and is relatively constant all along the year (Hiez & Dubreuil, 1963).

The studied area is mainly covered by a dense tropical forest with high ecological biodiversity. Along the coast, some areas of mangrove and swamp exist.

## **MATERIAL AND METHODS**

The present study concerns hydrology, suspended sediment and dissolved yields of the Maroni and Oyapock rivers, French Guiana. Suspended sediment and geochemical

data comes from samples collected in these two rivers between April, 2004 and January, 2008, in the frame of the ORE HYBAM observatory.

The Langa Tabiki gauging station on the Maroni River was installed by IRD (ex-ORSTOM) in 1952, and the Saut Maripa station on the Oyapock River in 1953. These two gauging stations are not influenced by oceanic tidal cycles (Hiez, 1955; Hiez & Dubreuil, 1963). In the past, water levels were monitored two times a day, and today, every half hour by limnometric “Cloe” type sensors. The Langa Tabiki and Saut Maripa gauging stations present good and stable rating curves, calibrated on the basis of respectively 77 and 80 discharge measurements from 1953 to 2001 at each station. Daily discharges have been calculated using HYDRACCESS software (available by free download at <http://www.mpl.ird.fr/hybam/outils/hydraccess.htm>).

The limit and extent of the basins were obtained by extracting the topographic limits river basins from Shuttle Radar Topography Mission (SRTM) map.

The total suspended solid material (TSS) was collected monthly in the center of the river, and then measured by filtration of 1 liter of water on 0.45  $\mu\text{m}$  cellulose acetate filters. After weighting, the suspended sediments were separated from the filter by agitation in an ultrasonic bath. They were transferred in Teflon® containers and subjected to a total dissolution procedure using hot HF – HNO<sub>3</sub>, after a short preliminary H<sub>2</sub>O<sub>2</sub> digestion eliminating organic material. Due to the nature of the filters used (nitrocellulose), no POC determination was undertaken. The TSS yield computation is discharge-weighted, using mean monthly discharge and TSS values of the month.

The samples for geochemical analysis of dissolved material were collected monthly in acid-washed HD polyethylen containers, near the water surface, in the middle of the river by IRD hydrologists of Cayenne. They were filtered in the IRD laboratory at Cayenne through 0.22  $\mu\text{m}$  DURAPORE® filters. Samples were stored in acid-washed HD polyethylen bottles. The samples for anions determinations were not acidified while those for major cations and trace element determination were acidified with ultrapure HNO<sub>3</sub>. The samples for dissolved organic carbon (DOC) determination were stored in glass bottles previously burned in an oven at 450°C during two hours; they were acidified with ultrapure H<sub>3</sub>PO<sub>4</sub>.

The major anions (F, Cl, NO<sub>3</sub>, PO<sub>4</sub>, SO<sub>4</sub>) were determined using ion chromatography and Ca, Mg, Na, K, Fe, Si, Al, Mn, Ti and P were determined by ICP-OES. For DOC analysis, a high temperature catalytic oxidation (HTCO) method was used. The TDS and elemental yields computation are also discharge-weighted.

The analytical precision, evaluated using repeated standard reference materials analysis, is generally better than 10% while reproducibility was determined using replicate sample analysis and is better than 5%.

## **RESULTS AND DISCUSSION**

### **Hydrology**

For the whole dataset, from 1951 to 2007 (Table 1 and Figure 2), annual mean discharge data for the Maroni River is 1 700 m<sup>3</sup> s<sup>-1</sup>, with a minimum of 820 m<sup>3</sup> s<sup>-1</sup> (1964) and a maximum of 2 640 m<sup>3</sup> s<sup>-1</sup> (1971). For this river, the variation coefficient (or RSD: standard deviation / mean value) is 0.28. At Saut Maripa, the Oyapock River

presents a mean annual discharge of  $830 \text{ m}^3 \text{ s}^{-1}$ , a minimum of  $360 \text{ m}^3 \text{ s}^{-1}$  (1983), and a maximum of  $1\,440 \text{ m}^3 \text{ s}^{-1}$  (1989). The variation coefficient is 0.24, indicating a relative stability of the inter-annual hydrological regime, probably due to the proximity of the Atlantic Ocean. The extreme values present a good similitude with the Amazon River discharge at the Óbidos gauging station in Brazil, where large flooding were observed in 1974-1976, 1989 and 1999, and low water in 1964, 1980, 1983, 1992 and 1998 (Callède et al., 2004).

Mean discharge observed during the 2004-2007 period (sampling period) is  $1\,840 \text{ m}^3 \text{ s}^{-1}$  and  $790 \text{ m}^3 \text{ s}^{-1}$  respectively for Langa Tabiki and Saut Maripa (Table 2), these values are very close to the mean value of the long time observation period (Table 1). The seasonal variability of the discharge is similar for the two stations (Figures 3a and b), with the beginning of the hydrological cycle in November, a first flooding period from December-January to February-March, and a second larger flood from April to July-August, as observed by Hiez and Dubreuil (1963).

### **Suspended Sediment load**

On the basis of the 47 samples collected monthly at each of the two stations, an initial characterization of sedimentary flow has been achieved. The Total Suspended Solid (TSS) concentrations range from 1 to  $130 \text{ mg l}^{-1}$  during the hydrological cycle, with discharge – weighted averages of  $21.9 \pm 1.6$  and  $12.1 \pm 1.2 \text{ mg l}^{-1}$ , for the Maroni and Oyapock rivers respectively (Table 2). This last value is very similar to the result of Lointier (1995) at Saut Maripa who obtained a mean value of  $12.7 \pm 3.4 \text{ mg l}^{-1}$ , in his one year survey with sampling every two days. It should be noted that part of the TSS is organic but due to sampling conditions the particulate organic carbon was not determined. The seasonal variability of these values is significantly higher than that of hydrologic flows (RSD from 0.89 for LT and 1.17 for SM), but without simple relationship with the flow (Figure 4a). The solid flow obtained show a similar variability, with peaks of sediment flows quite well related with the hydrological maxima (Figures 3a and b). These flows, with interannual average value of  $3\,750 \text{ t day}^{-1}$  at Langa Tabiki and  $940 \text{ t day}^{-1}$  at Saut Maripa correspond to very low physical erosion rates, which are respectively  $21$  and  $14 \text{ t km}^{-2} \text{ year}^{-1}$  for the Oyapock and the Maroni. These values are comparable to those observed on the rivers of Bolivia (Guyot et al., 1996) and of Brazil over the Brazilian and Guiana shields (Filizola, 2003) under tropical and equatorial climates. This is also observed in the Congo basin which is also located on an old shield and presents similar climatic conditions (Laraque et al., 1998, Laraque et al., 2009).

**Mineralogical and chemical composition.** X-ray diffraction was performed on some samples and indicated the presence of kaolinite as the main mineralogical component with low (less than 5%) quartz and amorphous phases. As observed on the whole Amazonian basin, kaolinite seems to be the characteristic clay mineral of river sediments issued from Brazilian and Guianese shields (Guyot et al., 2007). No noticeable variations were encountered during the hydrological cycles. The Table 3 presents the mean, minimum and maximum values of major elements in the suspended sediments from both stations. Due to analytical procedure,  $\text{SiO}_2$  content was not determined, and the dilution effect of quartz was not evaluated, since, as stated before, quartz concentrations are generally low. Ti and P were generally below detection limit, probably because of the low amount of material present on the filter and of the final dilution of sample before ICP-OES determination. The

sediments are mainly composed of Al and Fe, which correspond to a classical weathering product in tropical areas. No obvious differences are apparent between the stations, with the exception of K which presents a threefold increase at Langa Tabiki (mean: 1.23 %) with respect to Saut Maripa (0.48%).

By multiplying the relative concentrations (%) of the various elements by the sediment yield (Table 2), the flow of each element of the particulate phase is obtained (Table 3). The main exported element, apart from SiO<sub>2</sub> not determined, is the Al (507 t day<sup>-1</sup> at LT and 139 t day<sup>-1</sup> at SM), followed by Fe (234 t day<sup>-1</sup> to LT and 68 t day<sup>-1</sup> to SM).

### **Dissolved load**

The whole geochemical data base of the survey is available on the project website (<http://www.ore-hybam.org>). The Table 4 presents the mean values of the major elements and dissolved organic carbon (DOC) concentrations, the pH, and the conductivity during the four years of the survey for both rivers. This table also gives the minimum and maximum encountered values.

**Major elements.** The geochemistry of the dissolved material from the rivers of Guiana has been established on the basis of the 53 samples collected monthly since April, 2004.

Both rivers are characterized by dilute waters as indicated by a mean total cationic charge ( $\Sigma^+$ ) of  $\sim 200 \mu\text{eq/L}$  (Table 4). They present slightly acid to near neutral pH, low conductivity values and molar Ca<sup>2+</sup>/Na<sup>+</sup> ratio < 1. Although very diluted with regards to the Amazon river and its main tributaries, the Solimões and the Madeira rivers (Table 5), these waters are however more concentrated than the black water rivers found in the Amazon basin, such as the Rio Negro river (Gaillardet al.,1997), or the Caroni River (Edmond et al.,1995), in the Orinoco basin. Their total cationic and anionic charges are more similar to those found in the Congo basin (Dupré et al, 1996) but they differ by their lower dissolved silica content.

For SM, our values are very close to those found by Lointier (1995) who reported a mean conductivity of  $20.9 \pm 2.3 \mu\text{S cm}^{-1}$ , mean value of  $0.73 \pm 0.10 \text{ mg l}^{-1}$  for Ca<sup>2+</sup>,  $2.47 \pm 0.30 \text{ mg l}^{-1}$  for Cl<sup>-</sup> and  $4.74 \pm 0.72 \text{ mg l}^{-1}$  for Si. As already observed for the suspended load, it appears thus that the monthly time scale used in our study leads to the acquisition of results very similar to those acquired with a much higher frequency (every two days in the case of the work of Lointier).

The mean normalized inorganic charge balance ( $\text{NICB} = (\Sigma^+ - \Sigma^-) / \Sigma^-$ ) is very low (SM = 4%, LT = 5%). However, it has a large range of variation along the hydrological cycle (SM: from +44 to -24%, LT: from +44 to -19%). The Figure 5 indicates that these variations are clearly linked to the DOC concentration, as the samples with high NICB value present high DOC content and conversely. This confirms thus the observations reported by numerous other authors in tropical river basins like the Congo Basin (Dupré et al., 1996; Probst et al., 1992), the Amazon Basin (Gaillardet et al., 1997) and the Orinoco Basin (Edmond et al., 1995), who assumed that the imbalance could be linked to the dissociation of organic acids producing H<sup>+</sup> ions and organic anions. The dominant anionic species is HCO<sub>3</sub><sup>-</sup> followed by Cl<sup>-</sup>. The Maroni River has a 10% higher TDS value than the Oyapock River.

At Langa Tabiki, the major ions are generally inter-correlated, except NO<sub>3</sub><sup>-</sup> which is not interconnected with other ions (Table 6a). At Saut Maripa (Table 6b), Ca<sup>2+</sup> is

weakly correlated only with  $\text{Na}^+$  and  $\text{HCO}_3^-$  while other cations are well inter-correlated, and correlated with  $\text{Cl}^-$ . The very significant correlation between  $\text{Na}^+$ ,  $\text{K}^+$  and  $\text{Cl}^-$  should be emphasized and will be discussed further.

In terms of dissolved material discharge, the Maroni River has a significant higher exportation value with more than  $3\,910\text{ t day}^{-1}$  compared to  $1\,430\text{ t day}^{-1}$  in the Oyapock River (Table 2). As the drainage area at LT is much greater ( $64\,230\text{ km}^2$ ) than at SM ( $24\,630\text{ km}^2$ ), the calculated weathering rate is very low, nevertheless similar for both stations ( $22$  and  $21\text{ t km}^{-2}\text{ year}^{-1}$  respectively). Considering physical erosion rate and weathering rate, total denudation rate of these two equatorial drainage basins range from  $44$  (LT) to  $35$  (SM)  $\text{t km}^{-2}\text{ year}^{-1}$ . On the Maroni River, TDS and TSS yields are equivalent, whereas on the Oyapock River TDS represent 60% of the total exportation yield. Looking at the elemental composition of the total exported material, Al and Fe are the main components of the particulate phase while  $\text{Na}^+$  and  $\text{Ca}^{2+}$  dominate in the dissolved load (Table 7). The Soluble Transport Index ( $\text{STI} = \text{QD} / (\text{QD} + \text{QS})$ , Martin and Meybeck, 1979) is very high for the major cations, even if their absolute concentrations are relatively low.

**Seasonal variability.** All the major elements show a regular seasonal pattern (Fig. 6): the higher concentrations are related to the low water stage, from October to February, while more dilute waters occur during the flood from April to July. Figure 4b presents a slight negative relationship between discharge and TDS load for both gauging stations. Monthly discharge and TDS yield variability are very similar for the two rivers (Figure 3 a and b) but it should be noted that the TDS are higher than TSS nearly 95% of time for the Oyapock river whereas this occurs only nearly 50% for the Maroni river; this is probably due to an oceanic influence at SM (see below).

On a ternary cation ( $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Na}^+ + \text{K}^+$ ) diagram (Fig. 7), all the samples fall on a linear trend close to the ( $\text{Na}^+ + \text{K}^+$ ) apex; a slight displacement towards  $\text{Ca}^{2+}$  appears for some samples from both rivers. This pattern was found also by Edmond et al. (1995). As the computed estimated chemical composition of average igneous rocks falls in the same envelope, they concluded that, in first approximation, the weathering processes go to completion without retention of the cations in secondary phases. This should thus also be the case for the rivers of French Guiana.

The ternary anion ( $\text{Si}$ ,  $\text{HCO}_3^-$ ,  $\text{Cl}^- + 2\text{SO}_4^{2-}$ ) diagram (Fig. 8) shows that the LT samples have a more alkaline composition than the SM samples which are richer in  $\text{Cl}^-$  (mean value: SM =  $64.3\ \mu\text{mol L}^{-1}$ , LT =  $51.9\ \mu\text{mol L}^{-1}$ ). This higher  $\text{Cl}^-$  content probably reflects a greater contribution of sea-salt aerosol conveyed by the dominant winds blowing from east to west. On a Na-Cl diagram (Fig. 9), the SM samples are closer to the sea-salt line than the LT samples; their mean Na:Cl ratio is also lower (1.34 and 1.76 respectively). This ratio is however far above the value of sea-salt (0.86) indicating that Na has also a continental source. As can be seen on figure 10, the Cl:Na ratio at SM is generally well correlated with discharge, which is not the case at LT. This pattern is thus clearly related to the greater influence of the trade winds bringing the rain to that station.



## Atmospheric contribution

As atmospheric inputs can contribute significantly to the chemical composition of both rivers, we try to evaluate this impact on the dissolved load budget. The composition of the rain water has mostly an oceanic origin but can also be influenced by other sources.  $\text{Na}^+$  and  $\text{Ca}^{2+}$  concentration can be influenced by continental inputs, the  $\text{NO}_3^-$  have a mainly continental influence and  $\text{K}^+$  concentration can be influenced by the vegetation (Gaillardet et al., 1997; Negrel et al., 1997; Negrel and Lachassagne, 2000). According to Meybeck (1983), when  $\text{Cl}^-$  in river is exclusively originating from rainfall, the correction of atmospheric inputs can be estimated for any element X referring to the  $\text{Cl}^-$  concentration in the stream multiplied by the X/Cl ratios in the rainwater. In rivers, the main sources of  $\text{Cl}^-$  are the rain contribution, the halite dissolution (evaporitic rocks), the anthropogenic inputs (Meybeck, 1983; Stallard and Edmond, 1981) and the leaching of the canopy (Millet et al., 1999). In the studied area, no evaporitic rocks are present (Edmond, 1995). The vegetation impact is difficult to quantify (Chave, 2005), and we will consider this contribution as negligible compared to the great influence of sea proximity to the basins. The human impact has also been considered as negligible as the population density in both basins is less than 1 people per  $\text{km}^2$  (Guillemet, 2002). All the  $\text{Cl}^-$  exported by the Maroni and the Oyapock rivers has thus been supposed to come from atmospheric inputs.

As showed by Stallard and Edmond (1981), the  $\text{Cl}^-$  concentration in the rain decreases with the distance from the ocean. The greater influence of the Atlantic inputs on the Oyapock basin than on the Maroni basin induces higher  $\text{Cl}^-$  concentration at SM (mean value :  $2.28 \pm 0.16 \text{ mgL}^{-1}$ , Table 4) than at LT (mean value :  $1.84 \pm 0.4 \text{ mgL}^{-1}$ ). Without fractionation of the rain composition, the X/Cl ratio is constant in space and during time. Negrel et al. (1997) studied the time variability of rain composition in Cayenne (French Guiana) and Negrel and Lachassagne (2000) studied the space variability of rain composition on the Maroni basin. The X/Cl measured during those studies confirmed a spatial and time variability. The time variability of X/Cl ratio at Cayenne can be explained by the variation of the origin of air masses during the year while the spatial variability can be explained by the variation of continental inputs on rain water. To integrate these variabilities, we decided to calculate an average of X/Cl rain from the Negrel et al. (1997) and Negrel et Lachassagne (2000) data. On the Cayenne data, we excluded the September data because of the high influence of continental dust contamination on the sample reported by the authors. The standard deviation calculated on those values was used to evaluate the accuracy of the data. The data of table 8 were used to calculate:

$$X_{\text{rain}} = X/\text{Cl}_{\text{rain}} * \text{Cl}_{\text{riv}} \quad (1)$$

$$X\% = X_{\text{rain}}/X_{\text{riv}} \quad (2)$$

where :  $X_{\text{rain}}$  is the contribution of the atmospheric inputs to the element budget;

$X/\text{Cl}_{\text{rain}}$  is the ratio X/Cl of the rain;

$\text{Cl}_{\text{riv}}$  is the  $\text{Cl}^-$  concentration of the river;

X% is the relative atmospheric contribution to the element X.

The equations (1) and (2) were applied on each monthly value. The table 9 presents the average of the relative atmospheric contribution, the RSD calculated on the monthly relative atmospheric contribution and the variability of the X/Cl rain applied on the monthly values.

Due to weak contribution from rain inputs, the  $\text{Ca}^{2+}$  and  $\text{Mg}^{2+}$  mean values are  $15 \pm 3\%$  and  $15 \pm 5\%$ , respectively, on the Maroni river, and  $27 \pm 6\%$  and  $23 \pm 4\%$  on the Oyapock river. The source of those cations is thus mainly the water-rock

interaction. Conversely, a great rainfall contribution is observed for  $\text{Na}^+$  ( $48 \pm 6\%$  and  $63 \pm 8\%$  for the Maroni and for the Opayock respectively). It should be noted that Maroni river results are consistent with the conclusions of Negrel and Lachassagne (2000) who obtained similar proportions for  $\text{Ca}^{2+}$  and  $\text{Mg}^{2+}$  rain inputs (average  $10 - 14\% \pm 1.3 - 3.4\%$ ) and for Na ( $48.2 \pm 3.9\%$ ). Similar proportions were also recorded in the Congo Basin (Negrel et al, 1993). In the Amazon Basin, Gaillardet et al., 1997, observed similar proportion for  $\text{Na}^+$  and  $\text{Mg}^{2+}$ , but a lower contribution for  $\text{Ca}^{2+}$ . For  $\text{K}^+$ ,  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$ , it is difficult to evaluate the atmospheric contribution because of the great X/Cl variability observed in the rain for those ions, especially at SM. Furthermore, these ions present low concentrations in the river waters and are consequently very sensitive to the application of atmospheric correction. For  $\text{K}^+$ , the influence of the vegetation is another factor adding uncertainties in evaluating the atmospheric input (Negrel et al., 1993, Gaillardet et al., 1997).

## CONCLUSIONS

This study allowed to obtain for the first time continue (4 years) and reliable data on both the geochemical and sedimentary fluxes of the two main rivers of French Guiana.

Both rivers present relatively stable inter-annual hydrological regime, probably due to the proximity of the Atlantic Ocean. During the hydrological cycle, the observed variability of the suspended material as well as of the TDS is low, although it exceeds that of the water flow. As observed on other rivers flowing on the Brazilian and Guiana shields, the sedimentary fluxes are very low and correspond to physical erosion rates ranging from 14 (LT) to 21 (SM)  $\text{t km}^{-2} \text{ year}^{-1}$ . The mineralogical and chemical composition of the transported sediments is also typical of the material usually found in tropical areas, with kaolinite as the main mineralogical species and Al and Fe as the main elements exported in the suspended load.

The chemical composition of the waters reflects the geological settings of the area, characterized by the alteration of the igneous rocks of the Guiana shield. This results in very low concentrations of major elements and thus in relatively limited dissolved material discharge for both rivers. The clear relationship observed between the mean normalized inorganic charge balance and the DOC results from the dissociation of organic acids. The greater influence of trade winds at the Saut Maripa station induces a different Cl:Na pattern with respect to the one observed at Langa Tabiki.

The atmospheric inputs are higher for the Opoyack river than the Maroni river.  $\text{Na}^+$  is most impacted by rain contribution,  $\text{Ca}^{2+}$  and  $\text{Mg}^{2+}$  in rivers are mainly related to the rock-water interaction. It is difficult to give conclusions about the atmospheric contribution on the  $\text{K}^+$ ,  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  rivers contents because of the low concentrations found in the river and of the great variability observed in the rain for these ions.

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## List of tables

Table 1: Main characteristics of the studied basins

Code	River	Latitude	Longitude	Drainage basin (km <sup>2</sup> )	Rainfall (mm.year <sup>-1</sup> )	Period	Discharge			Ke (%)
		(degree N)	(degree W)				(m <sup>3</sup> .s <sup>-1</sup> )	(l.s <sup>-1</sup> .km <sup>-2</sup> )	(mm.year <sup>-1</sup> )	
LT	Maroni River at Langa Tabiki	5.1401	54.3551	64 230	2 520	1951-2007	1 700	26.5	840	33.3
SM	Oyapock River at Saut Maripa	3.7995	51.8883	24 630	2 680	1953-2007	830	33.7	1 060	39.6

Table 2: Discharge, TDS and TSS yields for the 2004-2007 period

	Unit	Maroni river				Oyapock river			
		Mean	Max	Min	RSD	Mean	Max	Min	RSD
Daily discharge	m <sup>3</sup> s <sup>-1</sup>	1 840	6 160	60	0,69	790	3 270	70	0,77
TSS	mg l <sup>-1</sup>	21,9	132,2	1,6	0,89	12,1	102,4	1,9	1,17
Sediment yield (QS)	t day <sup>-1</sup>	3 750	14 730	24	0,86	940	7 140	13	1,04
TDS	mg l <sup>-1</sup>	25,8	33,8	17,5	0,17	21,6	34,0	13,0	0,17
Dissolved yield (QD)	t day <sup>-1</sup>	3 910	10 650	160	0,58	1 430	6 310	150	0,72
QS+QD	t day <sup>-1</sup>	7 660				2 370			
QS/(QS+QD)	%	49				40			
Solid transport yield (QS)	(t km <sup>2</sup> .yr <sup>-1</sup> )	21				14			
Dissolved transport yield (QD)	(t km <sup>2</sup> .yr <sup>-1</sup> )	22				21			
Total (QS + QD)	(t km <sup>2</sup> .yr <sup>-1</sup> )	44				35			

Table 3: Mean, minimum and maximum value for major elements in the suspended sediments (in %), and elemental fluxes (QS\_el in t km<sup>2</sup> y<sup>-1</sup>)

Element	Maroni river					Oyapock river				
	Mean	Min	Max	RSD	QS_el	Mean	Min	Max	RSD	QS_el
Al	13.51	5.2	27.81	0.39	2,88	14.78	7.54	24.67	0.37	2,06
Fe	6.24	2.45	11.98	0.43	1,33	7.25	2.37	14.86	0.45	1,01
Ca	0.63	0.13	3.58	1.24	0,13	0.45	0.18	1.09	0.55	0,06
Mg	0.27	0.1	0.75	0.54	0,06	0.22	0.12	0.49	0.46	0,03
Na	0.23	0.04	0.62	0.69	0,05	0.20	0.03	1.11	1.16	0,03
K	1.23	0.22	7.64	1.37	0,26	0.48	0.26	1.06	0.49	0,07
Mn	0.14	0.04	0.56	0.81	0,03	0.14	0.05	0.51	0.89	0,02

Table 4: Mean, maximum and minimum values for the major elements in the dissolved load, and elemental fluxes (QD\_el in t day<sup>-1</sup>)

	Unit	Langa Tabiki					Saut Maripa				
		Mean	Max	Min	RSD	QD_el	Mean	Max	Min	RSD	QD_el
pH		6.2	7.0	5.1	0.07		6.2	6.9	5.5	0.06	
Conductivity	μS cm <sup>-1</sup>	28.6	82.0	21.0	0.38		23.3	42.0	17.0	0.24	
Fe	mg.l <sup>-1</sup>	0.13	1	0.01	1.48	18.7	0.1	0.6	0.01	1.05	6,8
Si	mg.l <sup>-1</sup>	4.83	6.22	3.12	0.16	730.1	4.36	6.01	2.73	0.18	304,4
Al	mg.l <sup>-1</sup>	0.09	0.59	0.01	1.16	16.9	0.12	0.55	0.02	1.03	10,8
Ca	mg.l <sup>-1</sup>	1.17	2.28	0.58	0.18	166.2	0.86	2.04	0.48	0.36	62,6
Mg	mg.l <sup>-1</sup>	0.64	0.84	0.36	0.20	93.4	0.5	0.67	0.33	0.18	33,6
Na	mg.l <sup>-1</sup>	2.08	3.43	1.27	0.20	300.0	1.96	2.43	1.19	0.18	134,7
K	mg.l <sup>-1</sup>	0.86	1.86	0.46	0.32	121.2	0.63	1.05	0.28	0.30	42,9
F	mg.l <sup>-1</sup>	0.02	0.02	0.02	0.03	2.9	<0.01	-	-	-	-
Cl	mg.l <sup>-1</sup>	1.84	2.28	1.45	0.14	284.4	2.28	3.14	1.49	0.16	157,8
NO <sub>3</sub>	mg.l <sup>-1</sup>	0.51	1.39	0.14	0.58	88.3	0.36	0.8	0.12	0.43	27,0
PO <sub>4</sub>	mg.l <sup>-1</sup>	<0.1	-	-	-	-	<0.1	-	-	-	-
SO <sub>4</sub>	mg.l <sup>-1</sup>	0.5	1.37	0.17	0.58	70.4	0.46	1.49	0.1	0.65	28,6
HCO <sub>3</sub>	mg.l <sup>-1</sup>	8.82	14.45	4.25	0.30	1260.8	6.29	13.6	1.65	0.37	427,0
DOC	mg.l <sup>-1</sup>	6.44	13.87	2.75	0.40	1020.0	8.21	50.43	2.46	0.39	359,0
Σ cations	μe.l <sup>-1</sup>	223.3	348	160.7	0.19		184.9	253.6	122.4	0.18	
Σ anions	μe.l <sup>-1</sup>	213.6	308.5	141.3	0.21		182.0	302.4	99.0	0.24	
Ca/Na		0.324	0.437	0.174	0.16		0.254	0.488	0.134	0.30	
NICB		5%	44%	-19%	2.14		4%	44%	-24%	2.99	

Table 5 : Comparison of the composition of the dissolved phase at Langa Tabiki and Saut Maripa with data from other tropical shield basins

	Unit	Maroni R. (a)	Oyapock R. (a)	Negro R. (b)	Solimões R. (b)	Madeira R. (b)	Amazon R. at Obidos (b)	Caroni R. at Capaura (c)	Oubangui R. (d)	Kasai R. (d)	Congo R. (d)
pH		6.2	6.2	4.9	7.1	6.7	6.8	4.63	6.3	6.4	6.4
Si	mg.l <sup>-1</sup>	4.83	4.36	1.68	4.21	3.93		0.29	6.36		5.32
Ca	mg.l <sup>-1</sup>	1.17	0.86	0.36	7.45	2.65	3.93	0.11	2.60	1.76	2.08
Mg	mg.l <sup>-1</sup>	0.64	0.50	0.10	1.22	1.22	0.80	0.37	1.00	0.92	1.36
Na	mg.l <sup>-1</sup>	2.08	1.96	0.32	2.41	1.40	1.47	0.35		1.22	1.13
K	mg.l <sup>-1</sup>	0.86	0.63	0.47	0.90	0.90		0.27	1.10	1.29	1.41
Cl	mg.l <sup>-1</sup>	1.84	2.28	0.57	2.20	0.46	1.17		0.77	0.91	1.16
NO <sub>3</sub>	mg.l <sup>-1</sup>	0.51	0.36	0.31				1.08			
HCO <sub>3</sub>	mg.l <sup>-1</sup>	8.82	6.29		29.40	15.60	17.40	31	14.34	9.46	11.22
Σ cations	μe.l <sup>-1</sup>	223	185	66	600	293	326	26	303	208	301
Σ anions	μe.l <sup>-1</sup>	214	182	39	592	428	355	4.63	283	164	239
NICB		5%	4%	41%	1%	-16%	9%		7%	21%	21%

(a) : this study

(b) :Gaillardet et al., 1997

(c) : Edmond et al, 1995

(d) : Dupré et al., 1996



Table 6 : Correlation coefficients between major ions at Langa Tabiki (a) and Saut Maripa (b)

(in bold : significant coefficient,  $n = 41$ ,  $\rho_0 = 0.40$  for  $t = 99.5\%$  )

(a)	Mg	Na	K	Si	Cl	NO <sub>3</sub>	SO <sub>4</sub>	HCO <sub>3</sub>
Ca	<b>0.40</b>	<b>0.80</b>	<b>0.73</b>	0.29	<b>0.46</b>	0.22	<b>0.40</b>	<b>0.49</b>
Mg		0.37	0.16	<b>0.64</b>	0.13	0.03	0.11	<b>0.72</b>
Na			<b>0.80</b>	<b>0.40</b>	<b>0.78</b>	0.01	<b>0.52</b>	<b>0.44</b>
K				0.11	<b>0.67</b>	0.31	<b>0.83</b>	<b>0.40</b>
Si					0.02	0.02	0.31	0.37
Cl						0.28	<b>0.69</b>	0.26
NO <sub>3</sub>							<b>0.49</b>	-0.21
SO <sub>4</sub>								0.01

(b)	Mg	Na	K	Si	Cl	NO <sub>3</sub>	SO <sub>4</sub>	HCO <sub>3</sub>
Ca	0.24	<b>0.44</b>	0.37	0.20	0.23	0.14	0.04	<b>0.47</b>
Mg		<b>0.55</b>	<b>0.54</b>	<b>0.74</b>	<b>0.44</b>	0.06	0.27	0.34
Na			<b>0.80</b>	<b>0.40</b>	<b>0.78</b>	0.01	<b>0.52</b>	<b>0.44</b>
K				0.23	<b>0.73</b>	0.05	<b>0.70</b>	0.22
Si					0.23	0.14	0.04	<b>0.47</b>
Cl						0.02	<b>0.75</b>	0.28
NO <sub>3</sub>							0.06	-0.22
SO <sub>4</sub>								0.16

Table 7 : Mean values of major elements exportation in the dissolved (QD\_el) and suspended (QS\_el) phases (in  $t \text{ km}^{-2} \text{ y}^{-1}$ ).  $STI = QD/(QD + QS)$

	Maroni R.			Oyapock R.		
	QS_el	QD_el	STI	QS_el	QD_el	STI
Al	2,34	0,10	4%	2,06	0,16	7%
Fe	1,15	0,11	8%	1,01	0,10	9%
Ca	0,06	0,94	94%	0,06	0,93	94%
Mg	0,04	0,53	93%	0,03	0,50	94%
Na	0,05	1,70	97%	0,03	2,00	99%
K	0,11	0,69	86%	0,07	0,64	90%

Table 8 : X/Cl rain data for the Maroni catchment (Négrel & Lachassagne, 2000) and for Cayenne (Négrel et al., 1997). Mean values are arithmetic mean (n.d. : non-determined value).

		Ca/Cl	Mg/Cl	Na/Cl	K/Cl	NO <sub>3</sub> /Cl	SO <sub>4</sub> /Cl
<b>Maroni basin</b>							
Location	Apagui	0.065	n.d.	0.814	0.215	0.021	0,014
	Manbari	0.144	0.08	1.168	0.222	0.078	0,09
	Poligodou Soula	0.099	0.041	1.2	0.354	0.016	0,041
	Grand Santi	n.d.	n.d.	1.15	n.d.	n.d.	n.d.
	Antecume Pata	0.095	n.d.	1.267	0.553	0.02	0,027
	Elae	0.148	n.d.	1.284	0.609	n.d.	n.d.
<b>Cayenne</b>							
Month	January	0.213	0.164	1.006	0.067	0	0,071
	February	0.042	0.143	0.872	0.149	0.016	0,009
	March	0.049	0.119	0.837	0.058	0.024	0,052
	April	0.081	0.083	0.547	0.091	0.017	0,056
	May	0.018	0.035	0.543	0.077	0.022	0,075
	June	0.042	0.069	0.495	0.051	0.119	0,091
	July	0.105	n.d.	0.421	0.337	0.124	0,074
	October	0.02	0.04	0.569	0.247	0.02	0,059
	November	0.085	0.093	0.648	0.276	0.044	0,042
	December	0.108	0.013	0.583	0.508	n.d.	n.d.
	Mean	0.082	0.073	0.838	0.254	0.04	0.054
	RSD	0.056	0.051	0.304	0.186	0.041	0.027

Table 9: Relative atmospheric inputs on the Maroni (LT) and Opayock (SM) rivers. The RSD monthly values are the RSD calculated on the relative monthly contribution of rain inputs. The rain composition RSD is the integration of the variability of the X/Cl in the calculation of the atmospheric contribution.

		Ca	Mg	Na	K	NO <sub>3</sub>	SO <sub>4</sub>
LT	mean	15%	15%	48%	61%	35%	64%
	RSD of monthly values	3%	5%	6%	11%	18%	21%
	rain composition RSD	13%	14%	20%	53%	54%	47%
SM	mean	27%	23%	63%	91%	53%	81%
	RSD of monthly values	6%	4%	8%	11%	24%	19%
	rain composition RSD	22%	18%	26%	89%	77%	69%

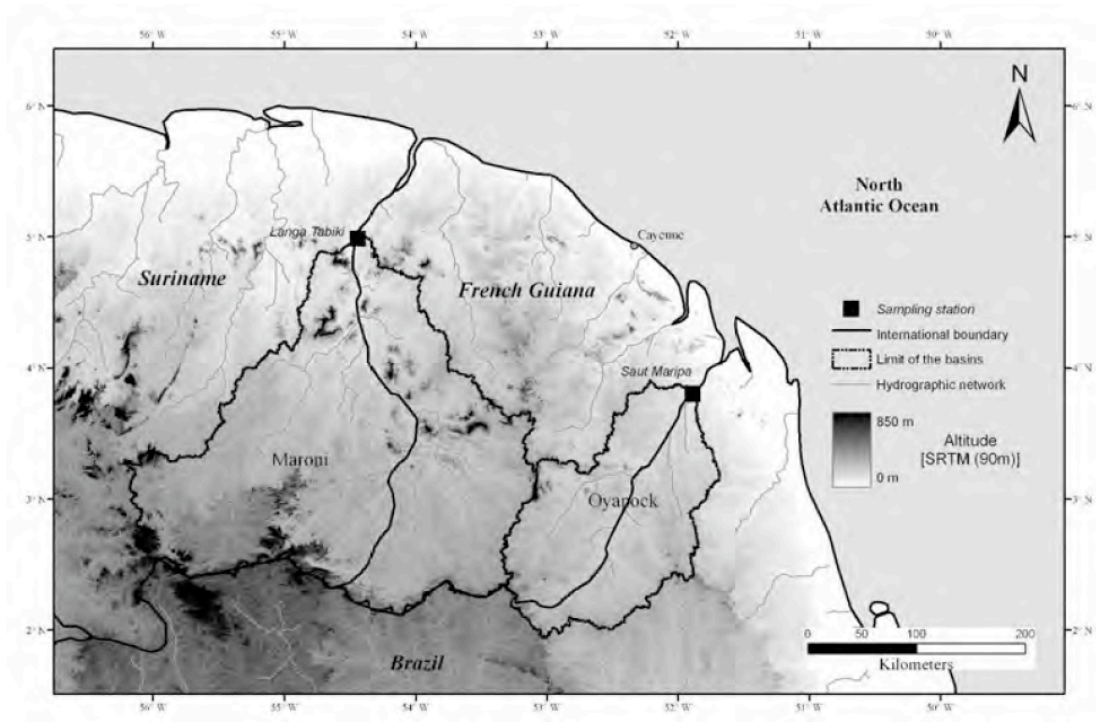


Fig. 1. Maroni and Oyapock rivers drainage basins (French Guiana)

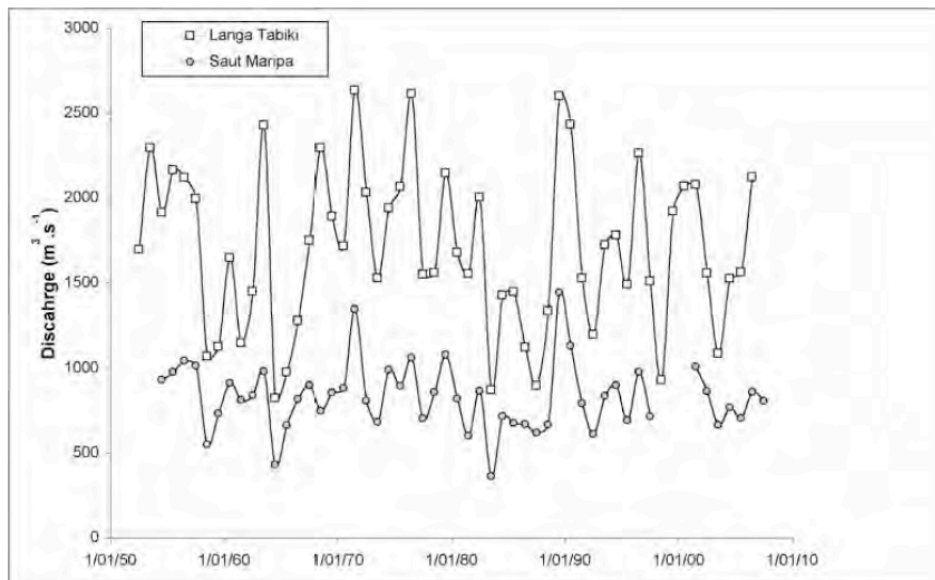


Fig. 2. Mean annual discharge (1951/1953-2007) Maroni River at Langa Tabiki and Oyapock River at Saut Maripa

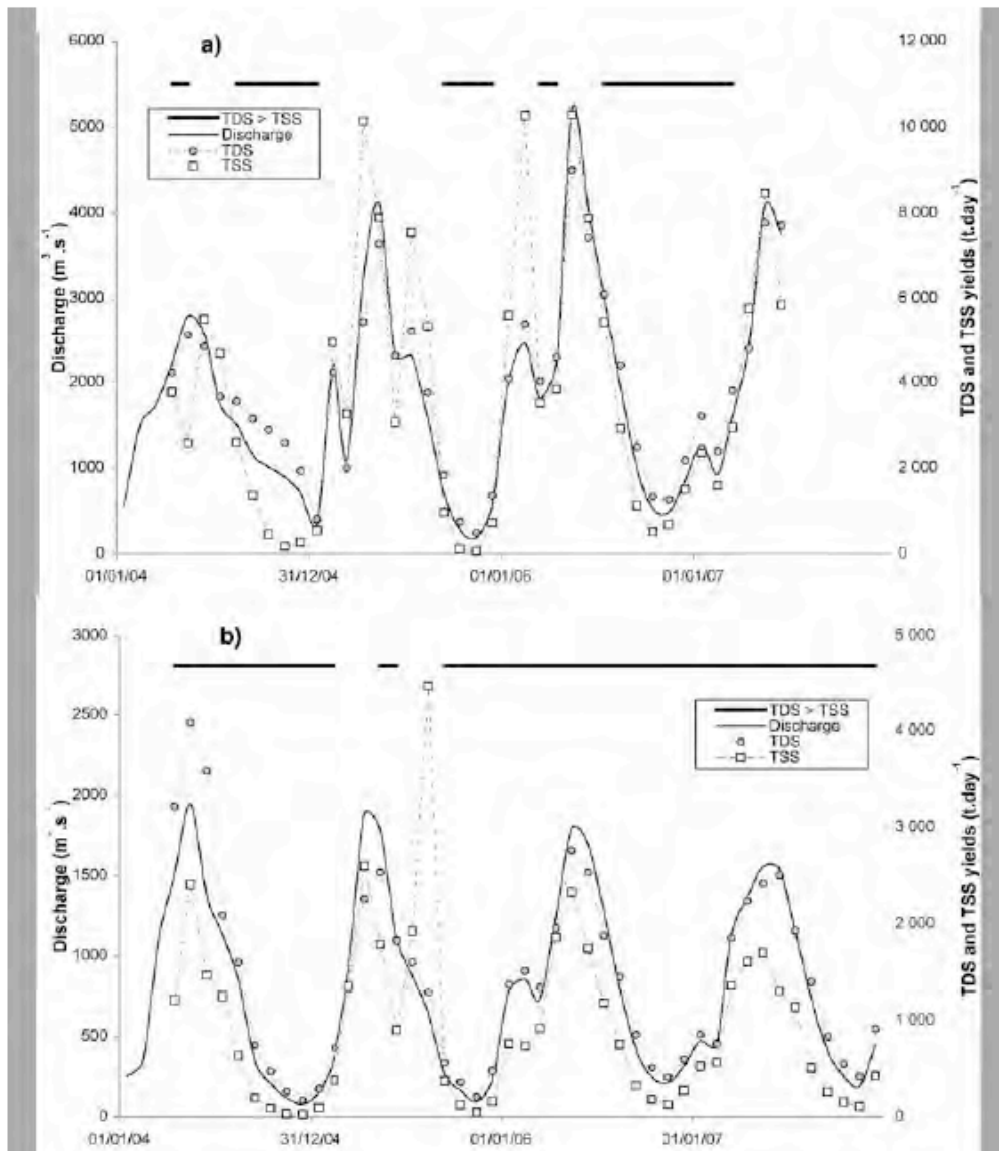


Fig. 3. Monthly discharge, dissolved and suspended sediment yields for (a) the Maroni River at Langa Tabiki (2004-2008) and (b) the Oyapock River at Saut Maripa (2004-2008). It should be noted that TDS is > to TSS nearly 95% of time for Oyapock and about 50% for Maroni.

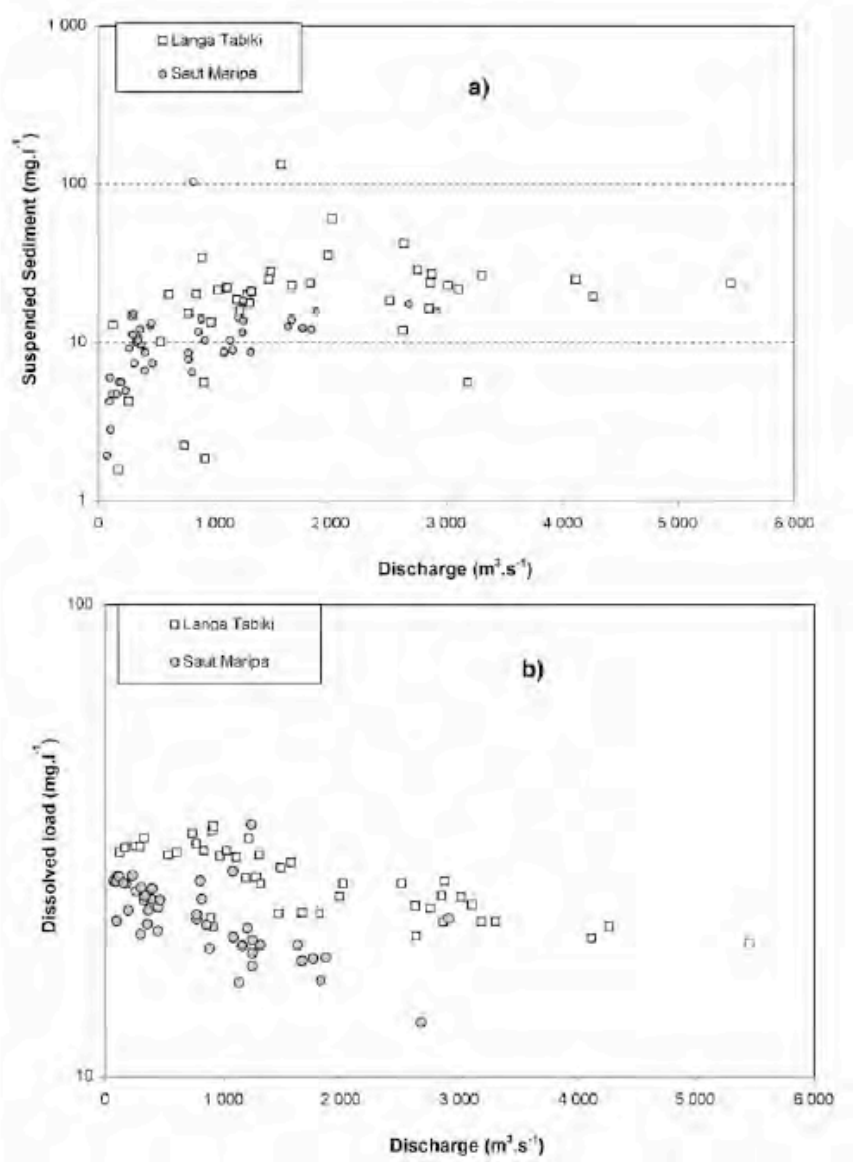


Fig. 4. Relation of the suspended sediment load with discharge (a) and of the dissolved load with discharge (b)

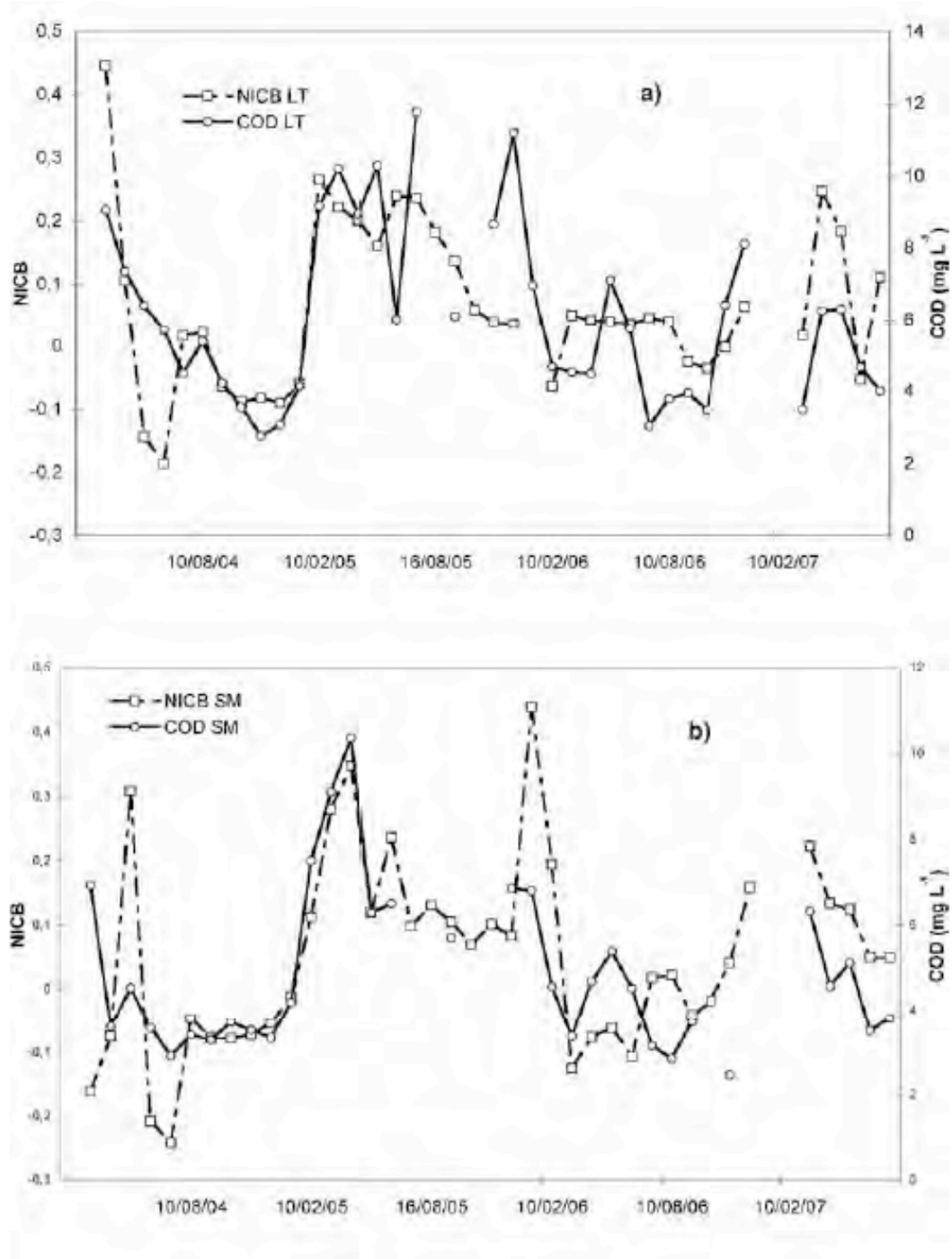


Fig. 5. NICB vs COD (mg L<sup>-1</sup>) content at Langa Tabiki (a) and Saut Maripa (b)

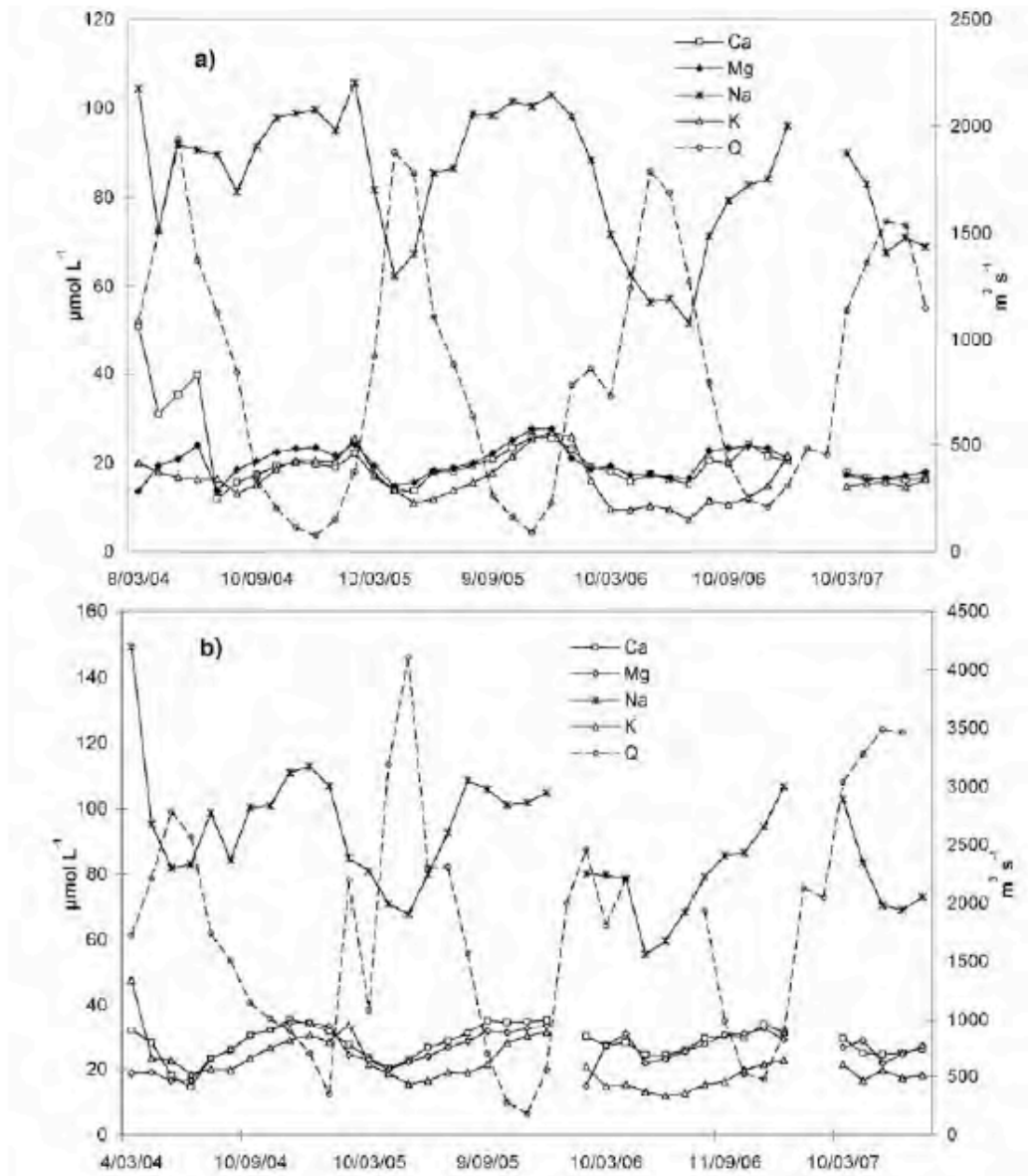


Fig. 6. Major cations content and water discharge at Saut Maripa (a) and Langa Tabiki (b)

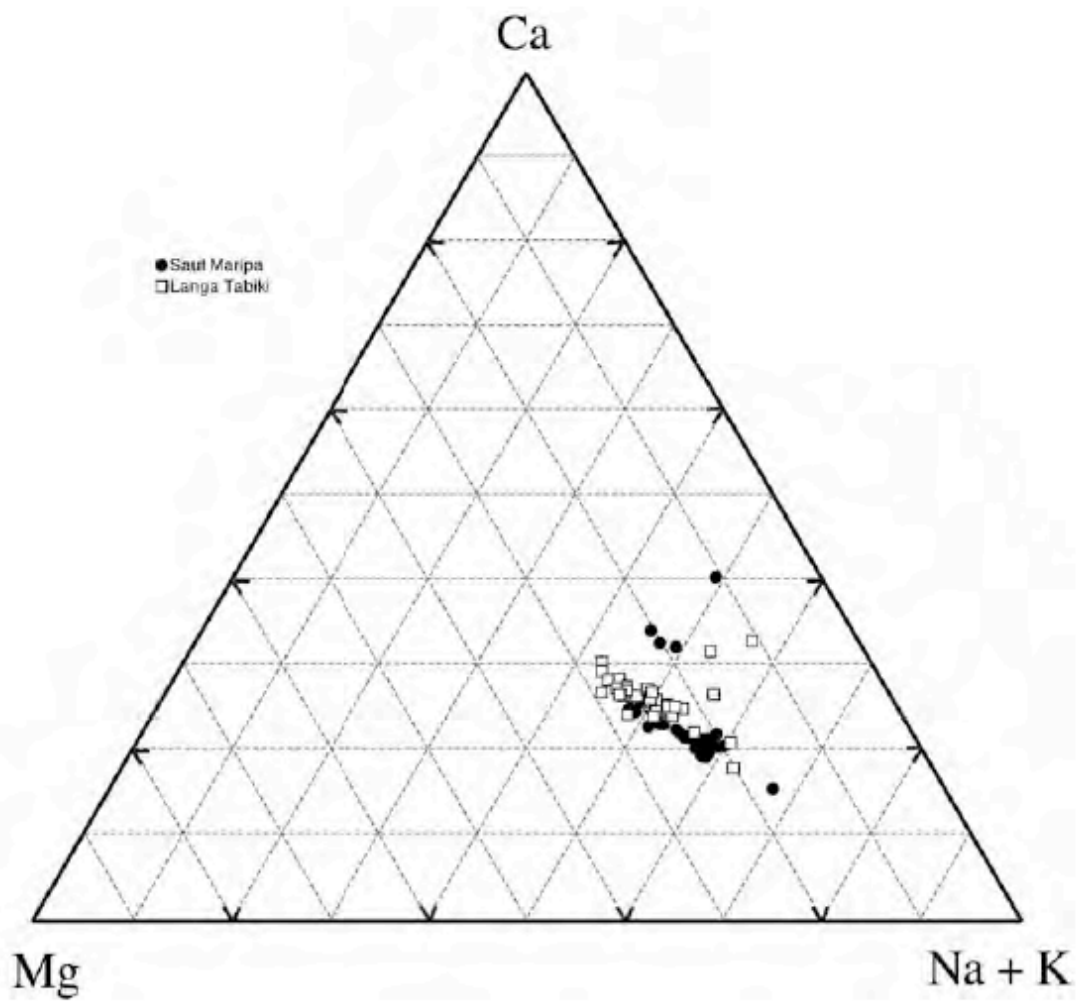


Fig. 7. Ca, Mg and Na+K ternary plot



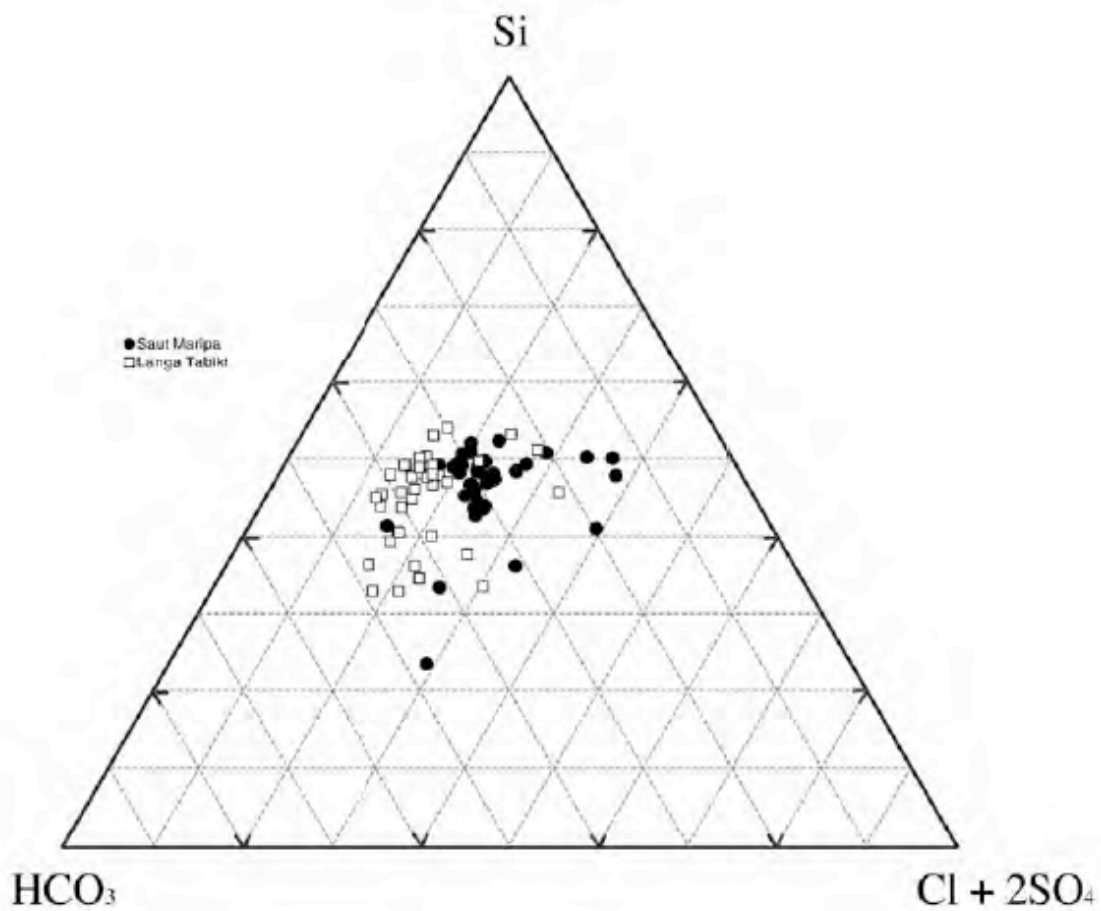


Fig. 8. Ternary plot of Si, alkalinity and Cl + 2SO4

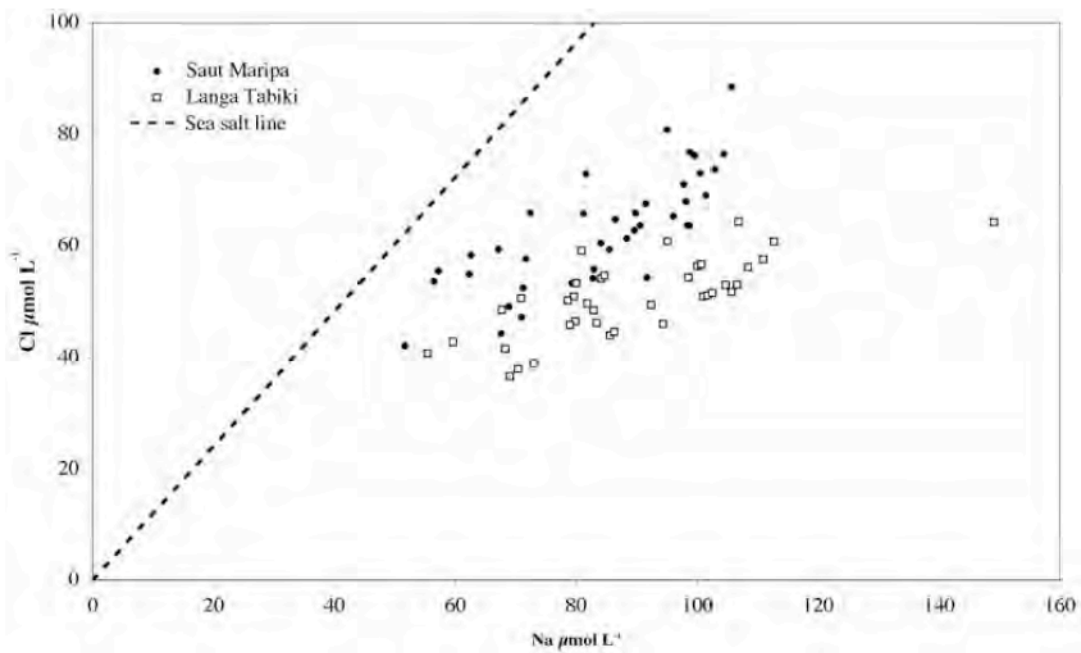


Fig. 9. Na / Cl ratio

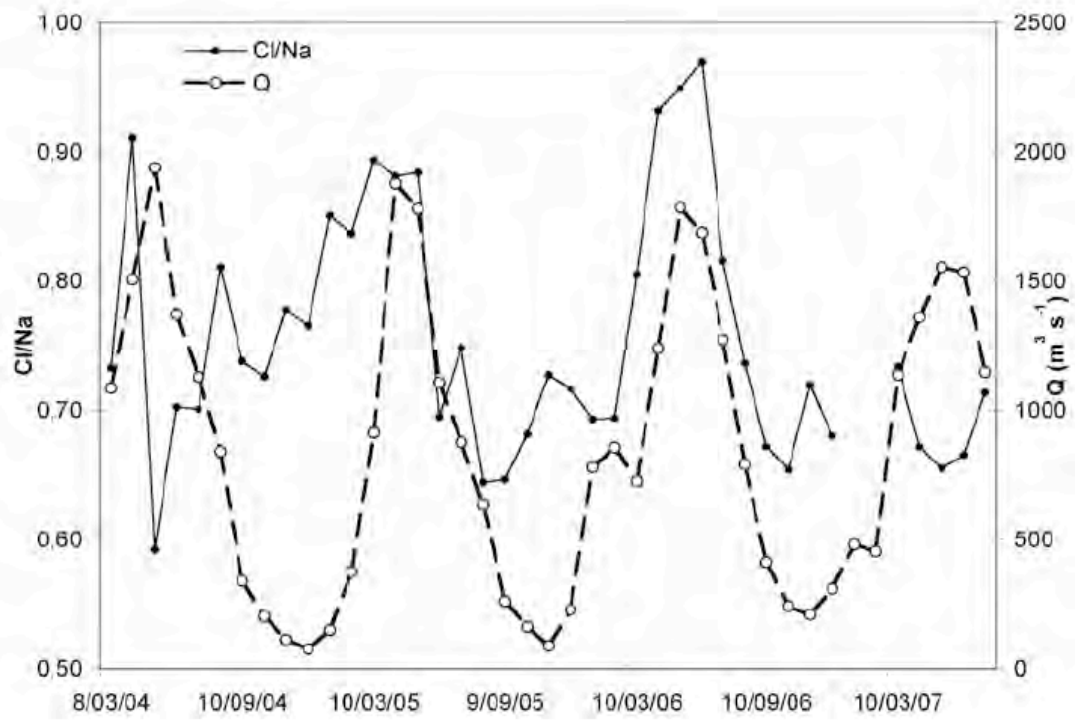


Fig. 10. Cl / Na ratio vs discharge at Saut Maripa