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► **To cite this version:**

Christoph Rüdiger, Jean-Christophe Calvet, Claire Gruhier, R.H. Holmes Thomas, A.M. de Jeu Richard, et al.. An Intercomparison of ERS-Scat and AMSR-E Soil Moisture Observations with Model Simulations over France. *Journal of Hydrometeorology*, 2009, 10, pp.431-448. 10.1175/2008JHM997.1 . ird-00394903

HAL Id: ird-00394903

<https://ird.hal.science/ird-00394903>

Submitted on 23 Jul 2009

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An Intercomparison of ERS-Scat and AMSR-E Soil Moisture Observations with Model Simulations over France

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Submitted to:
Journal of Hydrometeorology

Date of first submission
19 November 2007

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1 **Abstract**

2 This paper presents a study undertaken in preparation of the work leading up to the
3 assimilation of SMOS observations into the land surface model (LSM) ISBA at Météo
4 France. This study consists of an inter-comparison experiment of different space-borne
5 platforms providing surface soil moisture information (AMSR-E and ERS-Scat) with the
6 reanalysis soil moisture predictions over France from the model suite SIM (SAFRAN-ISBA-
7 MODCOU) of Météo France for the years 2003 to 2005. Both modelled and remotely sensed
8 data are initially validated against in-situ observations obtained at the experimental soil
9 moisture monitoring site SMOSREX in south-western France. Two different AMSR-E soil
10 moisture products are compared in the course of this study (the official AMSR-E product
11 from the National Snow and Ice Data Centre (NSIDC) and a new product developed at the
12 Vrije Universiteit Amsterdam and NASA (VUA-NASA)), which were obtained using two
13 different retrieval algorithms. This allows an additional assessment of the different
14 algorithms, while using identical brightness temperature data sets. This study shows that a
15 good correlation exists between AMSR-E (VUA-NASA), ERS-Scat, and SIM, generally for
16 low altitudes and low-to-moderate vegetation covers (1.5 to 3kg m² vegetation water content),
17 with a reduction in the correlation in mountainous regions. It is also shown that the AMSR-E
18 (NSIDC) soil moisture product has significant differences, when compared to the other data
19 sets.

20

21

22

22 1. Introduction

23 Soil moisture is the governing variable for modelling soil surface-to-atmosphere energy
24 exchanges and land surface model (LSM) initialisation, as it controls both evaporation and
25 transpiration from bare soil surfaces and vegetation covers. Consequently, a significant
26 amount of studies have been and are currently being conducted to obtain soil moisture
27 estimates through land surface modelling (e.g. Dirmeyer et al. 1999; Georgakakos and
28 Carpenter 2006) and remotely sensed surface soil moisture observations (e.g. Wagner et al.
29 1999ab; Kerr et al. 2001; Njoku et al. 2003).

30 For the purpose of soil moisture remote sensing, observations in the microwave bands
31 have been found to produce the best results. The optimal wavelength lies within the L-band
32 range (~1-2GHz), as interference through vegetation water content at this frequency range is
33 lower than at higher frequencies.. However, instruments have in the past been and are
34 currently operated at higher frequencies (above 5GHz), mainly because none of these
35 missions were dedicated soil moisture missions. The first such dedicated soil moisture
36 mission will be the Soil Moisture and Ocean Salinity mission (SMOS), to be launched in
37 2009. The first microwave instrument operated for an extensive time and within adequate
38 wavelengths was the Scanning Multichannel Microwave Radiometer (SMMR) on Nimbus-7
39 (operational from 1978 to 1987), which operated at bands at and above 6.6GHz. SMMR was
40 followed by the Special Sensor Microwave/Imager (SSM/I; since 1987) and the similar
41 Tropical Rainfall Measuring Mission (TRMM) Microwave Imager (TMI; since 1997), which
42 operate at frequencies above 10GHz. Instruments which are currently operational at
43 frequencies similar to SMMR (and therefore closer to L-band), are the Advanced Microwave
44 Scanning Radiometer for the Earth Observing System (AMSR-E) on board NASA's Aqua
45 satellite, WindSat on board the American Navy's Coriolis satellite, and the scatterometers on
46 board the European Remote Sensing satellites (ERS-1 & -2). Finally, a new scatterometer

47 (ASCAT) was launched on board ESA's MetOp satellite in 2006 and its data will soon be
48 available operationally (Bartalis et al. 2007).

49 Despite almost 30 years of experience with these microwave remote sensing instruments,
50 it is still necessary to validate the soil moisture products obtained from these instruments
51 through in-situ soil moisture observations. However, such ground-truthing has only been
52 achieved over small temporal and spatial scales (eg. the Soil Moisture Experiments (SMEX)
53 or the Campaign for validating the Operation of SMOS (CoSMOS)), as it is economically and
54 practically infeasible to observe soil moisture at high spatial and temporal resolution over
55 large scales using in-situ observations, mainly because of its high spatial variability. Only in
56 the present decade there have been attempts to establish long-term and large scale soil
57 moisture observation networks or data banks such as the Global Soil Moisture Data Bank
58 (Robock et al. 2000), the Goulburn River experimental catchment in Australia (Rüdiger et al.
59 2007), or SMOSMANIA in south-western France (Calvet et al. 2005). However, these data
60 sets only represent single points in space. This lack of spatial extent limits the usefulness of
61 such data sets for assimilation into large scale land surface models and also for disaggregation
62 studies, as the large scale, but also subpixel variability is not captured with single point
63 measurements. Moreover, satellite products are generally available at scales of about 0.25° or
64 25km, which leads to problems in their validation process, due to the different spatial scales
65 (spatially averaged satellite products are compared to point measurements). Consequently,
66 new validation methods complementing the existing soil moisture networks have to be
67 conceived (Wagner et al. 2007). Under the assumption that LSMs, forced with high quality
68 atmospheric forcing data, adequately represent the surface soil moisture dynamics, the scale
69 issues can be reduced. This assumption in turn will then allow the large-scale and long-term
70 evaluation of the satellite products in terms of their temporal dynamics, as the products
71 considered are essentially independent models.

72 In this paper, both the need for large scale ground-truthing and understanding of the
73 subpixel heterogeneity of soil moisture are addressed. First, the temporal correlation of
74 satellite products at a large scale with a synthetic high-resolution surface soil moisture data
75 base is presented. The high-resolution meteorological observation network throughout France
76 (more than 1000 surface meteorological stations and more than 3500 daily rain gauges) has
77 resulted in a high-quality atmospheric forcing data base (Quintana-Seguí et al. 2008) for the
78 operational land surface model ISBA of Météo-France, within the modelling system SIM
79 (SAFRAN-ISBA-MODCOU; Habets et al., 2008). The SIM model simulates the soil moisture
80 dynamics.

81 The satellite products used for this study were obtained from AMSR-E and ERS-2.
82 Furthermore, the recent development of a new retrieval algorithm for AMSR-E (Owe et al.
83 2007) allowed to compare the official AMSR-E product (Njoku et al. 2003) with this new
84 data base. In the first part of this study, the different data sets used are discussed, followed by
85 a brief comparison of those remotely sensed data sets and SIM with in-situ observations of the
86 SMOSREX experimental site near Toulouse, France (de Rosnay et al. 2006), to determine
87 their capability to represent the temporal soil moisture dynamics of a point or pixel. The good
88 results of this analysis between the land surface model, in-situ observations, and satellite data
89 also shows that previous results obtained over Spain (Wagner et al. 2007; one single satellite
90 pixel), or over Australia (Draper et al. 2007; several in-situ observations for a number of
91 pixels) can be extrapolated to a national or even continental scale, as they show the same
92 tendencies. The differences between the various soil type data bases used in the satellite
93 retrieval schemes and the model data base, make it difficult to compare absolute values.
94 Consequently, the discussion of this paper will focus on the normalised data sets. In the
95 second part, the inter-comparison study then presents the correlations and mean differences
96 between all data sets (ie. also between the different satellite products).

97

98 **2. Data Sets**

99 Due to the limitation in the spatial extent of SIM, this study is limited to watersheds of
100 mainland France. Nevertheless, the surface and climatic conditions throughout the country are
101 sufficiently variable (ranging from sub-humid to alpine), to give a statistically sound data
102 basis for a representative analysis. The years 2003-2005 were chosen for this study, as data
103 exists for all sources (SIM – 1970 to 2006; ERS-Scat – 1992 to 2006; AMSR-E – 2003 to
104 date; and SMOSREX – 2001 to date). Moreover, this 3-year period includes both very dry
105 and very wet climatic conditions, which are necessary to determine the dynamic range of the
106 soil moisture observations within each pixel. The following sections briefly outline the
107 various data sets, used for this study.

108

109 *a. SAFRAN-ISBA-MODCOU (SIM)*

110 The modelled surface soil moisture data base was obtained from the modelling system
111 SAFRAN-ISBA-MODCOU (SAFRAN – atmospheric forcing data base; ISBA – land surface
112 model; MODCOU – hydrological routing model). Of these three model chain segments, only
113 SAFRAN and ISBA were of importance for the present study.

114 SAFRAN (Système d'analyse fournissant des renseignements atmosphériques pour la
115 nivologie) is a reanalysis forcing data base, initially developed to improve snowfall and
116 avalanche forecasting. Within SAFRAN, the main atmospheric forcing parameters are
117 analysed. Each atmospheric parameter is analysed individually using an optimal interpolation
118 method. The final size of a grid cell within SAFRAN is 8x8 km². For precipitation, however,
119 the actual pixel sizes of SAFRAN vary, as they represent zones of climatic conditions rather
120 than regularly gridded areas. Each climatic zone covers about 1000km², resulting in about 600

121 such pixels over France. The forcing parameters are in principle assumed to be homogeneous
122 within one such pixel, however, they vary on a sub-pixel scale with topography. Apart from
123 precipitation the SAFRAN forcing data is available at 6-hourly intervals. Precipitation is
124 obtained using daily observations at the rain gauges and then interpolated into hourly time
125 steps as a function of the relative humidity during the day. The SAFRAN data base has
126 recently been validated against in-situ observations and found to be well correlated (Quintana-
127 Seguí et al. 2008).

128 The land surface model used in SIM is ISBA (Interactions of the Soil, Biosphere and
129 Atmosphere; Noilhan and Planton 1989; Mahfouf and Noilhan 1996), which is used
130 operationally as the land surface scheme within the numerical weather prediction system at
131 Météo-France. The soil layer and soil moisture dynamics are modelled within a 3-soil-layer
132 model (Boone et al. 1999), which is based on the force-restore approach, where the three soil
133 layers are a surface layer of 1cm depth, forming part of a root zone layer above the third, deep
134 layer.

135 There is no previous study presenting a verification of the SIM surface soil moisture
136 product. On the other hand, the ISBA model has been extensively validated for various
137 biomes. In particular, a number of studies exist comparing a point-specific calibrated ISBA
138 version to actual in-situ soil moisture observations in France (Calvet et al. 1998a,b, Boone et
139 al. 1999, Calvet and Noilhan 2000, and Sabater et al. 2007). The latter case corresponds to
140 SMOSREX, and the former to a previous experiment (MUREX) in the same region. The
141 RMSE for those cases is, respectively, 0.06 and $0.07 \text{ m}^3 \text{ m}^{-3}$, with a mean difference in the
142 order of 0.01 and $0.03 \text{ m}^3 \text{ m}^{-3}$. In both cases the Nash efficiency was calculated with 0.65/0.59.
143 Based on the results of Prigent et al. (2005), this level of error between in-situ observations
144 and model predictions is expected, while maintaining a good correlation despite the different
145 observation and model layer depths. The land surface parameters for ISBA are obtained from

146 ECOCLIMAP (Masson et al. 2003). The parameters provided by ECOCLIMAP are originally
147 provided at 1km resolution and are aggregated to the model resolution of 8km.

148

149 *b. AMSR-E*

150 AMSR-E is a passive microwave scanning radiometer, operating at six wavelengths
151 within the microwave spectrum (6.925, 10.65, 18.7, 23.8, 36.5, and 89GHz) in horizontal and
152 vertical polarisations, flown on NASA's Aqua satellite. The total swath width during an
153 overpass is approximately 1445km, with footprint resolutions ranging from 56km (6.925GHz)
154 to 5km (89GHz). Aqua is a sun-synchronous satellite orbiting Earth approximately 14 times
155 each day, with morning/descending and afternoon/ascending overpasses, at around
156 1.30am/pm. This configuration results in a repeat coverage of approximately every three days
157 in the equatorial latitudes and more frequent coverage in higher latitudes. For the particular
158 case of France, Aqua overpasses take place at 4 out of 5 days for both ascending and
159 descending orbits.

160 Currently, two different data products are freely available. The official product can be
161 obtained through the National Snow and Ice Data Center (NSIDC, hereafter AMSR-E
162 (NSIDC)), while a new product has recently been made available through the Vrije
163 Universiteit Amsterdam in collaboration with NASA (hereafter AMSR-E (VUA-NASA)).
164 Both products are briefly described in the following sections.

165 *i. AMSR-E (NSIDC)*

166 The AMSR-E (NSIDC) data used for this study were obtained from the operational Level
167 3 B03 AMSR-E data set (Njoku 2006). While the original resolution at 10.65GHz is ~38km,
168 the data is binned into regular 0.25°x0.25° pixels, through oversampling at 10km intervals.

169 The NSIDC method uses two low frequency dual polarized channels to optimize the three
170 parameters (soil moisture, vegetation optical depth and the effective soil temperature)
171 simultaneously. Originally, the method was developed and tested for the C- and X- band
172 channels. Unfortunately, severe radio-frequency interference (RFI) was discovered within C-
173 band (6.925GHz) over the USA and Japan and X-band over Italy and Great Britain (Li et al
174 2004, 2006). For this reason, the retrieval algorithm was applied to the X-band (10.65GHz)
175 and Ku-band (18.7 GHz) brightness temperatures. This has some important disadvantages: 1)
176 the 18 GHz channel introduces atmospheric influences and, 2) the observation depth of the
177 soil moisture product is reduced to 5-10mm, which is approximately half the potential range
178 of C-band and 3) vegetation attenuation effects are more significant than at lower frequencies.

179 ii. AMSR-E (VUA-NASA)

180 The VUA-NASA retrieval products from AMSR-E are derived according to the Land
181 Surface Parameter Model (LPRM) (Owe et al. 2007). The LPRM is a three-parameter
182 retrieval model for passive microwave data, using one dual polarized channel (either 6.925 or
183 10.65GHz) for the retrieval of both surface soil moisture and vegetation water content
184 (VWC). The land surface temperature is derived separately from the vertically polarized
185 36.5GHz channel.

186 The forward radiative transfer model in LPRM is based on one vegetation layer (τ - ω
187 approach) and the vegetation optical depth is parameterized as a function of the Microwave
188 Polarization Difference Index (MPDI) and soil moisture according to Meesters et al. (2005).
189 This method is applied globally, and requires no regional calibration or fitting parameters to
190 aid the retrieval process.

191 The main differences with the AMSR-E (NSIDC) soil moisture product lies in the use of
192 a higher frequency band for the retrieval of the land surface temperature (LST), and the

193 parameterization of the vegetation optical depth, leaving only the soil moisture to be
194 optimized.

195

196 *c. ERS-Scat*

197 The ERS-Scat data is obtained through active microwave remote sensing, ie. an energy
198 pulse is sent to the surface and the intensity of the returned signal is then used within the
199 retrieval algorithm to derive a relative soil moisture state. ERS-Scat is operated at 5.3GHz (C-
200 band), observing only the vertically polarised backscatter within this band, thus resulting in a
201 similar observation depth as AMSR-E. RFI has been found to have little impact on active
202 microwave remote sensing at this frequency. ERS-Scat has a morning/descending and
203 evening/ascending orbit at 10.30am/pm, with a varying repeat coverage of about 2 to 8 days.
204 The spatial resolution of an ERS-Scat footprint is in the order of 50km, while the soil
205 moisture product is binned into pixels of 0.25° (north-south extent) and 25km (west-east
206 extent).

207 The soil moisture product is provided in relative values, ranging from 0 to 100%. The
208 normalisation of the backscatter signal is done, using the minimum and maximum observed
209 backscatter from the 1992-2000 period, as dry and wet references. The retrieval algorithm is
210 described in detail in Wagner et al. (1999, 2003).

211 *d. SMOSREX*

212 SMOSREX is an experimental field site for in-situ and remotely sensed soil moisture
213 observations jointly operated by various research institutes in France and located to the south
214 of Toulouse (43°23'N, 1°17'E) in south-western France (De Rosnay et al. 2006). The overall
215 size of SMOSREX is approximately 6000m² separated into two areas with either bare soil or
216 fallow. The climate is temperate with monthly mean maximum temperatures of 5°C in winter

217 and 24°C in summer and an average annual cumulative precipitation of about 650mm. The
218 surface soil consists of a sandy loam, with 16% clay, 47% silt, and 37% sand.

219 Most instruments installed at the site have been in operation since 2001. The main feature
220 of this site is a tower-mounted L-band radiometer for the production of multi-angle brightness
221 temperatures. Other instrumentation include a weather station, and soil temperature and
222 moisture sensors, installed at various points and depths. The soil moisture sensors (Theta
223 Probes ©) used for this study are located at four points within the fallow section of the site
224 (most representative for the overall region and therefore the model simulations), with a
225 spacing of only a few metres. The sensors are vertically inserted at the surface, therefore
226 integrating the soil moisture content from 0 to 6cm, and a temporally averaged soil moisture
227 content is stored every 30 minutes individually for all sensors. The calibration of the sensors
228 is presented in (De Rosnay et al. 2006).

229 For the purpose of this study, the in-situ observations were aggregated into daily averages
230 and compared to the respective data sets obtained through the model and remote sensing.
231 Spatially averaging the observations of those four probes reduces the effect of spatial
232 variability within and increases the representativity of the soil moisture observations, and also
233 reduces the individual observations to one point in space.

234

235 *e. Data Preparation*

236 The results presented in this paper are based solely on the data sets from descending
237 orbits (nighttime) to avoid overly solar effects in the satellite data, due to sun glint and strong
238 temperature gradients between the vegetation and the surface, and also within the surface
239 layer, but also due to Faraday rotation and temperature gradients within the sensor which are
240 more pronounced during daytime overpasses (Kerr and Njoku 1990). Other effects such as

241 quick dry-down or the lack thereof due to local changes in solar radiation, which can not be
242 adequately represented in an LSM and in reality may be affected by cloud coverage and wind,
243 among other factors play a significant role in the daytime evolution of surface soil moisture.
244 While the in-situ observations were spatially and temporally averaged, the soil moisture
245 simulations were extracted for the time steps close to the overpass times of the satellites. A
246 comparison of the differences between the individual measurements of the soil moisture
247 probes and their spatial average at 6am and also between the daily average with the spatial
248 average at 6am resulted in an RMSE of $0.036 \text{ m}^3 \text{ m}^{-3}$ in both cases. This shows that spatial and
249 diurnal variabilities contribute to the same extent to the uncertainty in the in-situ observations.
250 The use of a spatially and temporally daily average is therefore justifiable.

251 All data have been reprojected from their original coordinate systems onto a regular
252 $0.25^\circ \times 0.25^\circ$ grid using a nearest neighbour approach. As the overall footprints of AMSR-E
253 and ERS-Scat are in the order of 50km with a spacing of about 10km between the centre
254 points, and the gridded products used in this study are binned at 25km or 0.25° , respectively, a
255 spatial shift in the data due to the reprojection process (a maximum of 12km) is not expected
256 to add any additional noise to the data or affect the data quality, as a footprint with its centre
257 12km from the pixel centre would still include information from more than half of the land
258 surface corresponding to the pixel area due to its size. To obtain an average pixel value within
259 the reprojected pixels, all original pixels with their centre falling into one reprojected pixel
260 were averaged to one single value. This average value was then assumed to be the
261 representative soil moisture of the reprojected pixel. In the case of the satellite observations,
262 only one original pixel would generally fall into a reprojected pixel, due to the similarity in
263 size, so that no errors are introduced due to the averaging of two satellite pixels. For all data
264 sets, the same general rule applied for the reprojection process, to avoid inconsistencies
265 between the data sets introduced through the reprojection and aggregation process.

266 In a brief study it was examined whether the variability between the soil moisture of high
 267 resolution SIM pixels with their averaged low resolution equivalent resulted in any errors
 268 within the analysis. However, no relationship between the this subpixel heterogeneity and the
 269 spatial distribution of the correlation coefficients between the different soil moisture products
 270 presented in the following section was found.

271 The soil moisture data from the satellites and SMOSREX were normalised following the
 272 approach presented by Pellarin et al. (2006), where the maximum and minimum of the soil
 273 moisture range was not determined by the soil type, but rather by the observed dynamic range
 274 within each individual pixel within the full study period (2003-2005). To exclude any
 275 abnormal outliers due to observational errors or instrument noise, the 90% confidence interval
 276 was chosen to define the upper and lower soil moisture content, respectively, using (1) and
 277 (2).

$$278 \quad int^+(SM) = \mu(SM) + 1.64 * \sigma(SM) \quad (1)$$

279 and

$$280 \quad int^-(SM) = \mu(SM) - 1.64 * \sigma(SM) \quad (2)$$

281 where int^+ and int^- are the upper and lower confidence limits; $\mu(SM)$ is the average soil
 282 moisture content for the pixel; and $\sigma(SM)$ the standard deviation of the soil moisture content
 283 for each pixel. With the knowledge of the upper and lower soil moisture content the absolute
 284 soil moisture value is then normalised using (3):

$$285 \quad \theta_n = \frac{SM_{obs} - int^-}{int^+ - int^-} \quad (3)$$

286 where SM_{obs} is the individual soil moisture observation and θ_n is its normalised soil moisture
 287 value. As a simplification it is assumed that the data are normally distributed, so that 90% of
 288 the data lie by definition within a range of $\mu \pm 1.64\sigma$. All data outside of this range were

289 discarded. Also, pixel values were excluded from the overall analysis, where SIM predicted
290 frozen soil water. As model simulations as such have no outliers due to instrumentation
291 errors, no screening of extreme values is required. The soil moisture from SIM is therefore
292 normalised using the modelled maxima and minima of each individual pixel, instead of int^+
293 and int^- .

294 Pixels located over major urban agglomerations (ie. Lille, Paris, Lyon, Bordeaux,
295 Toulouse, and Marseille) were not excluded. However, the correct representation of the soil
296 moisture is doubtful, as SIM is not capable to give realistic soil moisture conditions over
297 urban (and consequently sealed) areas, and moreover, the possibility of pixels subjected to
298 potential radio-frequency interference (Li et al. 2004) is higher in these areas. Nevertheless,
299 the number of these pixels is small (<0.5% of the total), compared to the total over France and
300 their overall effect on the statistical analyses was found to be negligible.

301

302 **3. Comparison of the soil moisture products with in-situ observations**

303 An evaluation of the surface soil moisture products obtained from SIM and the satellites
304 was undertaken, using the same three years of in-situ soil moisture observations as for the
305 remainder of this study (2003-2005). The in-situ data were obtained from the observations at
306 the experimental site SMOSREX. The data from the four surface soil moisture sensors
307 installed at SMOSREX, were averaged both spatially and over time, so that one daily
308 averaged observation was obtained for each day. This approach reduced the existing noise
309 levels in the in-situ observations, as discussed in the previous section. The model and satellite
310 data used here are the binned and reprojected data as for the large scale study in section 4, as
311 described above. SIM was not especially calibrated to the conditions at SMOSREX. For this
312 evaluation study various statistical parameters were calculated: the root mean square error

313 (RMSE), the mean difference or bias between two data sets, the correlation coefficient (r)
314 between two data sets, and the Nash efficiency coefficient (N). All statistics presented in the
315 following sections were calculated for the normalised soil moisture values and are therefore
316 dimensionless.

317 In a first step, the absolute values of the soil moisture products were compared with the
318 in-situ data. For this purpose, the already normalised ERS-Scat data were transferred into
319 absolute values, using the known maximum and minimum surface soil moisture observations
320 at SMOSREX. While a good correlation exists between SIM and SMOSREX data sets, a
321 severe lack of soil moisture dynamics is observed for the AMSR-E (NSIDC) data set (not
322 shown). However, the AMSR-E (VUA-NASA) data is well correlated despite an apparent wet
323 bias. Finally, the ERS-Scat observations are also well correlated in terms of their temporal
324 dynamics. In contrast to the AMSR-E (VUA-NASA) data, the ERS-Scat data exhibits a dry
325 bias. Due to the different soil moisture dynamics and biases, it is difficult to compare the
326 various data sets in detail, consequently, all comparison in the remainder of this paper will be
327 undertaken with normalised data (Fig. 1).

328 The comparison of the normalised SIM and SMOSREX data sets shows a good temporal
329 correlation ($r = 0.755$; $N=0.478$), with a bias (-0.083) towards the in-situ observations (ie. the
330 in-situ observations tend to be drier), with the exceptions of very dry conditions, when the
331 model has the tendency to overestimate the soil moisture at this site (Fig. 2). Throughout the
332 years, a higher level of surface soil moisture dynamics is observed within the model data (Fig.
333 1), which results in a root mean square error (RMSE) of 0.198. This phenomenon is explained
334 by inaccuracies in the forcing data due to the spatial interpolation process within SIM and the
335 differences in the thickness of the observed soil layers (1cm for SIM against 0-6cm for the
336 ThetaProbes). However, there are only few data points causing this noise and this is
337 consequently deemed acceptable.

338 The normalised AMSR-E (NSIDC) data display a very high variation, with interchanging
339 peaks and troughs every three months (Fig. 1). Every year, minimal values are reached during
340 winter and their maximum in summer. This recurring negative correlation with the in-situ data
341 results in a high RMSE and low overall correlation ($r = 0.132$; $N = -0.734$; bias = 0.132; RMSE
342 = 0.356). In contrast to the comparison of the absolute data, the persistent wet bias in the
343 AMSR-E (VUA-NASA) data has been reduced due to the normalisation. Similar to the SIM
344 predictions, a strong correlation between in-situ and remotely sensed data is found in this case
345 with a wet bias towards the AMSR-E data ($r = 0.775$; $N = 0.471$; bias = 0.072; RMSE = 0.194).
346 Finally, the ERS-Scat observations are also well correlated over time with the in-situ data ($r =$
347 0.618 ; $N = 0.125$), however a dry bias (-0.085) results in a more significant RMSE of 0.244
348 than for the AMSR-E (VUA-NASA) data. As ERS-Scat data are only available from August
349 2003 onwards, the identical periods of data cover (ie. August 2003 – December 2005) for the
350 other surface soil moisture products was undertaken (not shown), in order to verify that the
351 first seven months did not introduce significant biases in the statistical analyses, which then
352 would not be seen in the ERS-Scat comparisons. The differences in correlation, RMSE, and
353 bias did not change significantly for any of the inter-comparisons covering either the full three
354 years or the period August 2003 – December 2005. Consequently, all comparisons shown in
355 the remainder of this paper are based on the full period. The Nash efficiency coefficient for
356 SIM and AMSR-E (VUA-NASA) are acceptable. They are also similar to each other
357 suggesting that the two data sets perform equally well compared to the in-situ observations,
358 while the low Nash efficiency of ERS-Scat is due to the relatively strong bias in the satellite
359 data. In the case of the AMSR-E (NSIDC) data, the negative Nash efficiency suggests by
360 definition that an average value of the in-situ observations would compare better with the
361 overall observations than the remotely sensed observations. This is an important finding as it
362 shows the extreme difference between the in-situ observations and the satellite product.

363 Four aspects have to be considered for the cause of the differences observed in this
364 evaluation: i) the scale difference (8km and 0.25° for the model and the satellite, respectively,
365 against a single point observation), as the comparison or validation of soil moisture products
366 at different spatial scales will remain difficult in most cases, unless a representative catchment
367 average soil moisture monitoring site (Grayson and Western 1998) can be identified; ii) the
368 soil data base, as the model soil information constitutes an average of the soil particle size
369 distribution within an 8km/0.25° pixel, which may result in significant differences compared
370 to the soil conditions at the point of observation (the particle size analysis for SMOSREX
371 yielded 16% clay, 47% silt, and 37% sand; the particle size distribution within ECOCLIMAP
372 is 25%/25%/50%), iii) the forcing data, as it is obtained by interpolation between observations
373 and atmospheric predictions, which may miss localised events, iv) the observation depth, with
374 the model layer of 1cm and approximately the same depth for the satellite observations
375 against the integrated soil moisture content at 0-6cm for the in-situ observations, may result in
376 different dynamics.

377 Considering the above four aspects, SIM, AMSR-E (VUA-NASA) and ERS-Scat perform
378 well when compared to the SMOSREX in-situ observations, and also show a good
379 representation of the dynamic behaviour of the soil moisture content. For SIM, an RMSE of
380 0.198 with a dynamic range of the surface soil moisture at the site of $\sim 0.3 \text{ m}^3 \text{ m}^{-3}$, can be
381 translated into an absolute error in the soil moisture of just under $0.06 \text{ m}^3 \text{ m}^{-3}$. This result is
382 particularly good, as SIM was not calibrated to the conditions at SMOSREX, but rather used
383 the vegetation and soil conditions obtained from ECOCLIMAP. Moreover, despite the
384 differences in scale these errors are identical to the performance of the site-specific calibrated
385 model.

386 Depending on the application, the calculated error may be considered large or acceptable.
387 For atmospheric studies, it is more important to obtain a good representation of the temporal

388 dynamics, while the absolute soil moisture state is less important. On the other hand, an error
389 of $0.06 \text{ m}^3\text{m}^{-3}$ exceeds the validation goals of future satellite missions (Kerr et al, 2001). In
390 the first case, the evaluation of satellite data against any benchmark is necessary, shown by
391 the lack of temporal dynamics in the AMSR-E (NSIDC) data. In the second case, two factors
392 influencing the RMSE have to be considered to qualify the above value of $0.06 \text{ m}^3\text{m}^{-3}$: i) the
393 mean difference or bias between SIM and SMOSREX and ii) the spatial uncertainty of the in-
394 situ observation. Biases, in the case of the normalised data 0.084 (or $0.025 \text{ m}^3\text{m}^{-3}$), may be
395 removed using various techniques (eg. Drusch et al., 2005), while the uncertainty in the
396 spatial averaging of the four in-situ observations is in the order of $0.036 \text{ m}^3\text{m}^{-3}$. In particular
397 the removal of the bias would lead to a significant decrease of the RMSE. Consequently, it is
398 concluded that SIM may be used with reasonable confidence for a large scale model
399 intercomparison study, assuming that ECOCLIMAP provides similarly good information for
400 all other model pixels.

401 While the correlations derived from Fig. 1 are relatively large for SIM, AMSR-E (VUA-
402 NASA) and ERS-Scat, much of the captured variability is seasonal (dry in summer, wet in
403 winter). In order to assess the coherence with the in-situ observations and to avoid seasonal
404 effects, monthly anomalies are calculated. The difference to the mean is calculated for a
405 sliding window of five weeks, and the difference is scaled to the standard deviation. Table 1
406 shows seasonal scores, including the Kendall statistics and p-value. All the products are
407 significantly correlated to the in-situ observations, except for satellite products at specific
408 periods of the year. While SIM presents significant correlations throughout the year, all the
409 satellite products are not significantly correlated to in-situ observations at wintertime (DJF).
410 This may be explained by the sensitivity of the microwave signal (either active or passive) to
411 soil freezing and by the reduced dynamics of the surface soil moisture at wintertime. Both
412 VUA-NASA and NSIDC products present high correlations of the anomalies for the other

413 seasons. On the other hand, ERS-Scat has significant correlations at springtime (MAM), only.
414 The lack of significance of ERS-Scat during the summer and autumn seasons (JJA and SON,
415 respectively), may be explained by the small number of observations over the SMOSREX site
416 (28 and 41, respectively), compared with AMSR-E (184 and 175, respectively, for the VUA-
417 NASA product).

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423 **4. Inter-Comparison (all Data Sets)**

424 *a. General Correlation (all data)*

425 Fig. 3 shows correlation maps of the different remote sensing and modelled data sets for
426 all surface conditions and all years. A good correlation exists between the three data sets
427 AMSR-E (VUA-NASA), ERS-Scat, and SIM, in particular for regions of herbaceous
428 vegetation over regions with little relief, with a range of the coefficient of correlation from 0.2
429 to 0.9. Areas with denser vegetation, such as the forest of Les Landes in the South-West along
430 the Atlantic coast show a lower level of correlation, which would have to be expected due to
431 the masking effect on the microwave emissions of the soil moisture through vegetation.
432 Similarly, low correlations are found in regions with strong relief such as the Massif Central
433 and the Alps. The good correlation of ERS-Scat with SIM in the Italian Alps should be
434 ignored, as only a few data points were available due to the overpass rate of ERS over the
435 region and the filtering of days with frozen soils or snow. Mountainous regions cause errors in
436 both the modelling of soil moisture and its retrieval from satellite observations. First, there

437 exists a high level of uncertainty in the soil depth and its variability in those regions, which
438 impacts on the predictions of the soil moisture dynamics in the SIM model. Secondly, relief
439 interferes with the retrieval of low resolution remotely sensed soil moisture observations and
440 may cause considerable levels of errors (Mätzler and Standley 2000). The AMSR-E (NSIDC)
441 product has virtually only low correlations with any of the other data sets, even producing
442 negative correlations overall (Table 2).

443 This analysis also shows that previous results obtained over Spain (Wagner et al. 2007;
444 one single satellite pixel), or over Australia (Draper et al. 2007; several in-situ observations
445 for a number of pixels) can be extrapolated to a national or even continental scale, as they
446 show the same tendencies. In particular, the lack of soil moisture dynamics within the AMSR-
447 E (NSIDC) data set are apparent and are shown in all studies.

448 The data used to derive the spatial plots of Fig. 3 are summarised in Table 2 as showing
449 the respective coefficient of correlation (r), root mean square error (RMSE) and bias between
450 the data sets. Compared to SIM, the ERS-Scat data set has the highest overall correlation ($r =$
451 0.728) and lowest RMSE (0.201), followed by AMSR-E (VUA-NASA) with an $r = 0.491$ and
452 an RMSE of 0.297 . As mentioned before, AMSR-E (NSIDC) has a negative correlation of $-$
453 0.014 with an RMSE of 0.370 . The RMSE presented here is the RMSE obtained from the
454 normalised results, ie. it represents the relative error of the soil moisture dynamical range.
455 Assuming an average dynamic range of $0.3 \text{ m}^3\text{m}^{-3}$ and that SIM gives accurate in-situ
456 observations, this would translate into an average error of $0.056 \text{ m}^3\text{m}^{-3}$ for ERS-Scat, which is
457 higher than the design accuracy of SMOS ($0.04 \text{ m}^3\text{m}^{-3}$).

458 Like the bias between the SIM and SMOSREX data sets, the biases shown between SIM
459 and the three satellite products are all positive. This suggests that a consistent dry bias exists
460 within SIM. A first explanation for the bias between SIM and SMOSREX are the different
461 thickness of the observed soil layers (1cm in the model against 0-6cm in-situ), as the deeper

462 profile of the in-situ observations is likely to maintain a higher soil moisture content, as it is
463 less affected by evaporation than the thin surface layer in the model. Furthermore, other
464 aspects such as erroneous soil type information, biased forcing data, and biases in the soil
465 moisture retrieval for the satellites may result in consistent biases.

466 *b. Correlations Specific to Land Surface Cover*

467 A comparison of vegetation maps with the results of Fig. 3 suggested a connection
468 between the accuracy of the remotely sensed soil moisture information and the land cover.
469 Therefore, the dominant land surface cover within each satellite-type pixel was determined,
470 using the information from ECOCLIMAP, in order to identify vegetation specific correlations
471 for each data product. For this purpose, the different vegetation types within each 0.25° pixel
472 were aggregated into three dominant cover types: i) cultivated soils, ii) grasslands, and iii)
473 forests (Fig. 4). Relatively good correlations exist between SIM, ERS-Scat and AMSR-E
474 (VUA-NASA) for the two herbaceous vegetation covers (Fig. 5a & b). Like in the analysis of
475 the overall data set, ERS-Scat and SIM have the highest correlation coefficient and lowest
476 RMSE. Similarly, the pairs SIM/AMSR-E (VUA-NASA) and AMSR-E (VUA-NASA)/ERS-
477 Scat have slightly lower correlation coefficients and higher RMSEs, and AMSR-E (NSIDC)
478 having negative correlations throughout. These results (with the exception of the AMSR-E
479 (NSIDC) data) are not surprising given that remotely sensed soil moisture information should
480 theoretically be retrievable with a high level of accuracy over herbaceous vegetation types. In
481 herbaceous vegetation covers, active and passive methodologies are expected to show similar
482 performances, especially when using a similar frequency. The higher correlations of ERS-
483 Scat and SIM as compared to AMSR-E (VUA-NASA) and SIM, shows potential for
484 improvement of the AMSR-E (VUA-NASA) product. Part of the difference might be
485 explained by the limited range of moisture values in the optimization routine for the AMSR-E
486 (VUA-NASA) product (0-50%). For the retrieval of the current AMSR-E (VUA-NASA) data,

487 the soil moisture content is limited to a maximum of $0.5 \text{ m}^3\text{m}^{-3}$. However, it was found that
488 the surface soil moisture states often reached this point of saturation (Fig. 1). Consequently, if
489 this constraint were to be relaxed, and the retrieval process were allowed to produce higher
490 values, a quasi-normalised soil moisture product may be obtained (this aspect has been
491 considered for the next version of soil moisture data, which has recently been made
492 available). However, as a consequence of this constraint, the maximum soil moisture is
493 currently underestimated, which leads to an underestimation of the dynamic range, and
494 consequently a wet bias in the AMSR-E (VUA-NASA) data. The methodology behind the
495 ERS product avoids this caveat, by scaling between minimum and maximum observed signal
496 over the period 1992-2000.

497 The comparison of the various data sets for forested regions (Fig. 5c) overall shows lower
498 correlations and higher RMSEs. Again, ERS-Scat produces the best correlation with SIM,
499 followed by AMSR-E (VUA-NASA) and AMSR-E (NSIDC). Moreover, the ERS-Scat soil
500 moisture product appears to conserve its good correlation with SIM from the analysis of the
501 herbaceous vegetation types. Under the assumption that SIM is equally valid for forested
502 regions as for regions with low vegetation, it may be concluded that two effects may influence
503 the consistency of ERS-Scat for different vegetation types. Firstly, the retrieval process of
504 ERS-Scat implicitly takes into consideration the vegetation type by scaling the current signal
505 between the wet end dry ends of its long-term data base. This statement has significance for
506 other soil moisture missions in both active and passive microwave remote sensing, as the
507 approach taken for the retrieval of ERS-Scat soil moisture may be applied along with more
508 sophisticated radiative transfer models. Secondly, the ERS-Scat is well calibrated and has a
509 low radiometric noise of about 0.15 dB, which allows estimating soil moisture even in areas
510 where abundant forest cover reduces the effective sensitivity of backscatter to soil moisture.

511 An aspect of the data visible within the scatterplots of Fig. 5 is the apparent bi-modality
512 of the SIM data with data clouds forming for the lower and upper value ranges. Fig. 6 and 7
513 show histograms of the surface soil moisture from the four different low-resolution data
514 sources (SIM, AMSR-E (VUA-NASA), AMSR-E (NSIDC), and ERS-Scat) for SMOSREX
515 and for the whole of France, respectively. The histograms of the various data sources show
516 different patterns at the local scale (Fig. 6). While SIM and ERS-Scat show clear bi-
517 modalities, this is not the case for the two AMSR-E products, with AMSR-E (VUA-NASA)
518 having several peaks with a saturation at 1, and AMSR-E (NSIDC) data being almost
519 normally distributed, though all data sets, have a minima in the range of 0.4 to 0.6. The
520 histogram of the in-situ data at SMOSREX (Fig. 6e) also shows a bi-modality, although with
521 its maximum in the wet spectrum. This would suggest that preferred soil moisture states exist
522 at SMOSREX, but that the distribution is not correctly captured by the various models.

523 The non-normal distribution of the histograms have significance for the normalisation
524 process, as it was previously assumed that the soil moisture distribution was sufficiently
525 normal at each point. A violation of the assumption of normality would mean that the 90%
526 confidence interval could not be calculated with the equations (1) and (2). To assess this, the
527 distribution of the soil moisture states at the national scale was studied (Fig. 7).

528 An exception here is SIM with a clear peak in the dry spectrum (0.2) and AMSR-E
529 (VUA-NASA) being skewed towards the wet end (Fig. 7). The overall distributions show that
530 SIM retains its clear bi-modality with a peak in the dry spectrum, while the ERS-Scat and
531 AMSR-E data become more normally distributed. For the AMSR-E data sets, the distribution
532 of AMSR-E (NSIDC) data becomes almost Gaussian with a slight skew towards the wet end,
533 while the AMSR-E (VUA-NASA) data is more evenly distributed. As the normalisation
534 procedure of Pellarin et al. (2006) is only applied to the AMSR-E data, it is concluded that the
535 normalisation process is still applicable to the majority of the pixels throughout France.

536 The results shown here are in line with other studies. For example, Teuling et al. (2005)
537 showed that preferred soil moisture states may exist locally. However, they found that this
538 effect could not be observed at all sites studied and that it could not be linked to local soil
539 conditions and may therefore be a random effect. This conclusion is supported by Fig. 6 and
540 7, where the histograms for the data at SMOSREX suggest that local preferred wet and dry
541 states exist, while the distribution of all observations over France is not bimodal.

542 *c. Intra-seasonal Correlation*

543 The bi-modality presented in the previous section is unlikely to be caused by differences
544 in the soil types, as the soil moisture data were normalised, and SMOSREX also appears to
545 have this distribution (Fig. 6e). The bi-modality is related to the varying soil moisture states,
546 which are caused by either precipitation events or seasons. As the effect of precipitation
547 events on the soil moisture distribution is difficult to obtain, the results obtained for the
548 cultivated soils in Fig. 5 were separated according to the various seasons. This analysis (Fig.
549 8) clearly shows the different preferred soil moisture states in summer (dry) and winter (wet),
550 which are consequently the main reason for the creation of the data clouds in Fig. 5. Similar
551 results of preferred soil moisture states during the various seasons has been shown by Settin et
552 al. (2007), where they were largely attributed to the precipitation intervals and intensities
553 during the various seasons. Interestingly, the two AMSR-E products have nearly the same
554 correlations with SIM during springtime, which would suggest that the two radiative transfer
555 models work similarly well during this period.

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558 **5. Conclusion**

559 In this paper, an intercomparison study of several remotely sensed surface soil moisture
560 products with the re-analysis LSM predictions over France has been presented. First, the LSM
561 predictions, and the satellite observations were compared with a 3-year in-situ surface soil
562 moisture data set from an experimental site in south-western France (SMOSREX) to
563 determine their capability to represent the temporal dynamics of a point or pixel. A good
564 correlation was found between the model predictions and the in-situ data, despite a slight dry
565 bias within the model predictions. Based on this evaluation, it was then assumed that the land
566 surface model predictions over France may be used as a credible approximate estimate in the
567 absence of more direct surface soil moisture observations for the whole country.

568 The analyses of this study, have shown that two of the three satellite data sets (AMSR-E
569 (VUA-NASA) and ERS-Scat) have generally a good correlation with the model predictions,
570 while the AMSR-E (NSIDC) data set did not correlate well with any of the other data sets.
571 Generally, the AMSR-E (NSIDC) data showed a significant lack of seasonal soil moisture
572 dynamics, which was well captured by the other data sets. These results suggest that the
573 AMSR-E (NSIDC) data set is not correct, as three other independent models (a physically
574 based radiative transfer model, an empiric soil moisture retrieval scheme, and a land surface
575 model) show a good correlation with each other. This is further supported by the good
576 correlation between SIM, AMSR-E (VUA-NASA), ERS-Scat and the in-situ observations at
577 SMOSREX. It is possible that those three models are all wrong and coincidentally produce
578 the same results, though the comparison with SMOSREX suggest that this is not likely. The
579 results of the observations obtained from the scatterometer additionally highlights the
580 potential use of active microwave data sets, which will be continued by the MetOp ASCAT
581 observations.

582 The analysis of de-trended time series (monthly anomalies) of surface soil moisture over
583 the SMOSREX site shows that short term variations of SIM and all the satellite products

584 (included the NSIDC AMSR-E product) are meaningful. The significance is less for ERS-
585 Scat, which has a high sampling time.

586 For the moment it has to be acknowledged that there exists a good correlation between
587 some products for densely vegetated areas, but further studies are required to validate their
588 physical meaning or relevance. Given that we present only the temporal dynamics in this
589 paper, it is interesting to learn that some satellite products appear to represent those dynamics
590 better than others, even for forested areas.

591 While in-situ observations averaged to the land surface model or remotely sensed pixel
592 scale may be better suited for the evaluation of both land surface or radiative transfer models,
593 these observations are still sparse and difficult to obtain. This study presents an alternative to
594 the use of in-situ observations for such large scale evaluations through the inter-comparison
595 of independent and apparently similar soil moisture estimates from different models.

596

597 Finally, the good correlations between point observations and the low resolution model
598 predictions and satellite observations also show the importance of single point observations
599 for the verification of LSM and remotely sensed soil moisture products. They also support the
600 need of the installation of new and the maintenance of existing soil moisture monitoring
601 networks. This is particularly true for forested and mountainous regions, which in the past
602 have been neglected when new soil moisture monitoring sites were established. With the need
603 for the evaluation of land surface model performances and satellite validation campaigns, the
604 relatively few existing networks are not sufficient.

605

606

607 **Acknowledgments**

608 The authors acknowledge the NSIDC for the AMSR-E (NSIDC) data set. Eric Martin and
609 Jean-François Mahfouf from the CNRM are thanked for providing access to the SIM data
610 base and fruitful discussions. During the course of this study, Christoph Rüdiger was financed
611 through a post-doctoral fellowship with the Centre National d'Etudes Spatiales (CNES). The
612 SMOSREX project was co-funded by the «Programme National de Télédétection Spatiale»
613 and by the «Programme Terre Océan Surface Continentales et Atmosphère» (CNES), and by
614 participants to the experiment: CESBIO (CNES, CNRS, IRD, UPS), CNRM/GAME (Météo-
615 France, CNRS), INRA, and ONERA, all in the framework of the SMOS science preparatory
616 program.

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Figures

Figures

Figure 1. Time series plots (2003-2005) of the normalised values of the in-situ observations at SMOSREX (black lines) and the four surface soil moisture products, SIM, AMSR-E (NSIDC), AMSR-E (VUA), and ERS-Scat (+). The model predictions and satellite observations were obtained from respective low-resolution pixels covering SMOSREX.

Figure 2. Scatterplot of the normalised in-situ soil moisture observations at SMOSREX (vertical axis) with the four low-resolution data sets (SIM, AMSR-E (VUA), AMSR-E (NSIDC), and ERS-Scat) for the years 2003-2005. Darker regions show a higher density of data points.

Figure 3. Maps of the coefficient of correlation between the various soil moisture products (normalised values) over mainland France. The circles highlight the 6 major metropolitan areas of France.

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Figure 6. Histograms showing the relative frequency (vertical axis) of the various normalised soil moisture observations (horizontal axis) and predictions for the years 2003-2005 for the SMOSREX site: SIM model, AMSR-E product of VUA-NASA, AMSR-E product of NSIDC, ERS-Scat product of University of Vienna, in situ observations.

Figure 7. Histograms showing the relative frequency (vertical axis) of the various normalised soil moisture observations (horizontal axis) and predictions for the years 2003-2005 for whole of France: SIM model, AMSR-E product of VUA-NASA, AMSR-E product of NSIDC, ERS-Scat product of University of Vienna.

Figure 8. Scatterplots showing the comparison of the various soil moisture products for pixels with herbaceous vegetation only (cultivated soils and grasslands) for the four seasons a) spring, b) summer, c) autumn, and d) winter. The scatterplots and their corresponding statistics are located on opposite sides of each figure, ie. the scatterplot of the data pair SIM-AMSR (VUA) is in the top left hand corner, while the respective statistical values are found in the bottom right hand corner. Darker regions show a higher density of data points.

Tables

Table 1 – Comparison of monthly anomalies of surface soil moisture products (SIM, AMSR-E, ERS-Scat) with in-situ 0-6cm observations at the SMOSREX site, for three pooled annual cycles (2003 to 2005).

Product	Season	Number	Correlation	Bias	RMSE	Kendall τ	Kendall p-value
SIM	All	794	0.61	0.01	0.79	0.63	****
SIM	DJF	121	0.44	-0.03	0.93	0.53	**
SIM	MAM	219	0.74	0.01	0.65	0.73	****
SIM	JJA	255	0.58	0.03	0.78	0.58	****
SIM	SON	199	0.65	0.04	0.79	0.66	****
AMSR-E (NSIDC)	All	698	0.46	0.01	0.88	0.39	****
AMSR-E (NSIDC)	DJF	95	0.27	-0.20	0.99	0.17	NS
AMSR-E (NSIDC)	MAM	192	0.62	0.11	0.77	0.54	****
AMSR-E (NSIDC)	JJA	219	0.23	0.02	1.03	0.21	**
AMSR-E (NSIDC)	SON	192	0.54	0.01	0.88	0.48	****
AMSR-E (VUA-NASA)	All	606	0.38	0.01	0.97	0.38	****
AMSR-E (VUA-NASA)	DJF	75	0.12	-0.07	1.05	0.01	NS
AMSR-E (VUA-NASA)	MAM	172	0.49	0.09	0.90	0.53	****
AMSR-E (VUA-NASA)	JJA	184	0.28	0.01	1.00	0.28	***
AMSR-E (VUA-NASA)	SON	175	0.42	0.00	1.01	0.44	****
ERS-Scat	All	133	0.34	-0.08	0.85	0.30	**
ERS-Scat	DJF	32	0.57	-0.12	0.51	0.39	NS
ERS-Scat	MAM	32	0.55	-0.06	0.54	0.57	*
ERS-Scat	JJA	28	0.28	-0.10	0.64	0.30	NS
ERS-Scat	SON	41	0.19	0.02	0.81	0.07	NS

*The monthly anomaly is the difference to the mean divided by the standard deviation, for a period of 5 weeks. The Kendall τ is a non-parametric measure of correlation that assesses how well an arbitrary monotonic function could describe the relationship between two variables, without making any assumptions about the frequency distribution of the variables. It is used to measure the degree of correspondence between two rankings and assessing the significance of this correspondence. The p-value indicates the significance of the test, if it is small (below 0.05 at least), it means that the correlation is not a coincidence. The following thresholds on p-values are used: (i) NS (non significant) for p-value greater than 0.05, (ii) * between 0.05 and 0.01, (iii) ** between 0.01 and 0.001, (iv) *** between 0.001 and 0.0001 and (v) **** below a value of 0.0001.*

Table 2 – Statistics of the inter-comparison between the difference data sets (normalised surface soil moisture data). The values in each cell correspond to the coefficient of correlation, bias, and RMSE, respectively.

		SIM	ERS-Scat	AMSR-E (VUA-NASA)
AMSR-E (NSIDC)	r bias RMSE	-0.014 0.215 0.370	-0.099 0.040 0.363	-0.115 0.043 0.361
AMSR-E (VUA-NASA)	r bias RMSE	0.491 0.177 0.297	0.397 0.099 0.296	
ERS-Scat	r bias RMSE	0.728 0.093 0.201		

Figures

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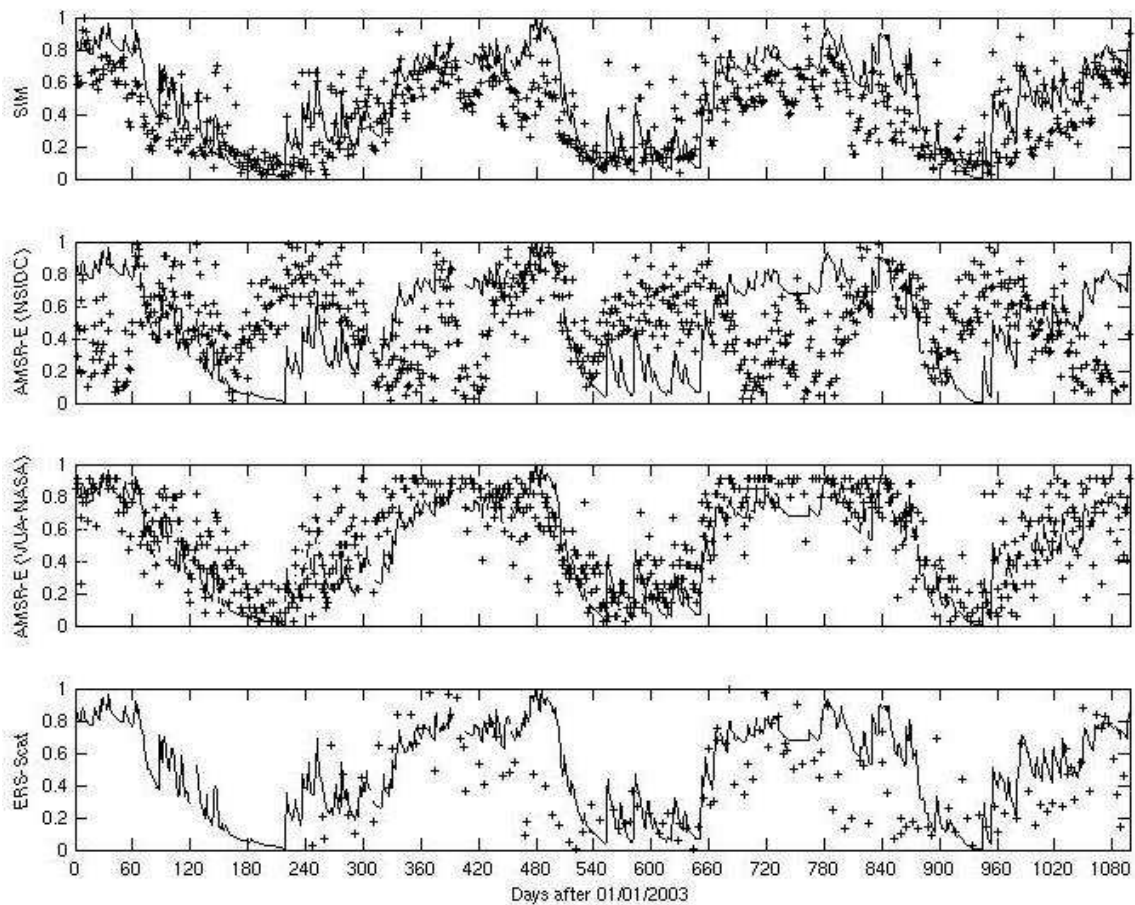


Fig. 2 Scatterplot of the normalised in-situ soil moisture observations at SMOSREX (vertical axis) with the four low-resolution data sets (SIM, AMSR-E (VUA), AMSR-E (NSIDC), and ERS-Scat) for the years 2003-2005. Darker regions show a higher density of data points.

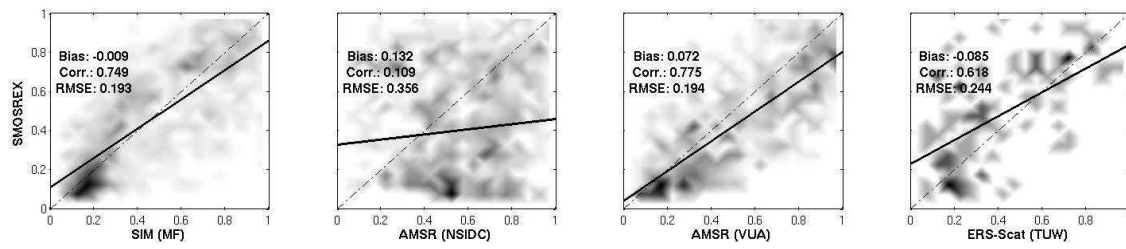


Fig. 3 Maps of the coefficient of correlation between the various soil moisture products (normalised values) over mainland France. The circles highlight the 6 major metropolitan areas of France.

Soil Moisture (Normalised) Correlation Maps for France, Years 2003-2005

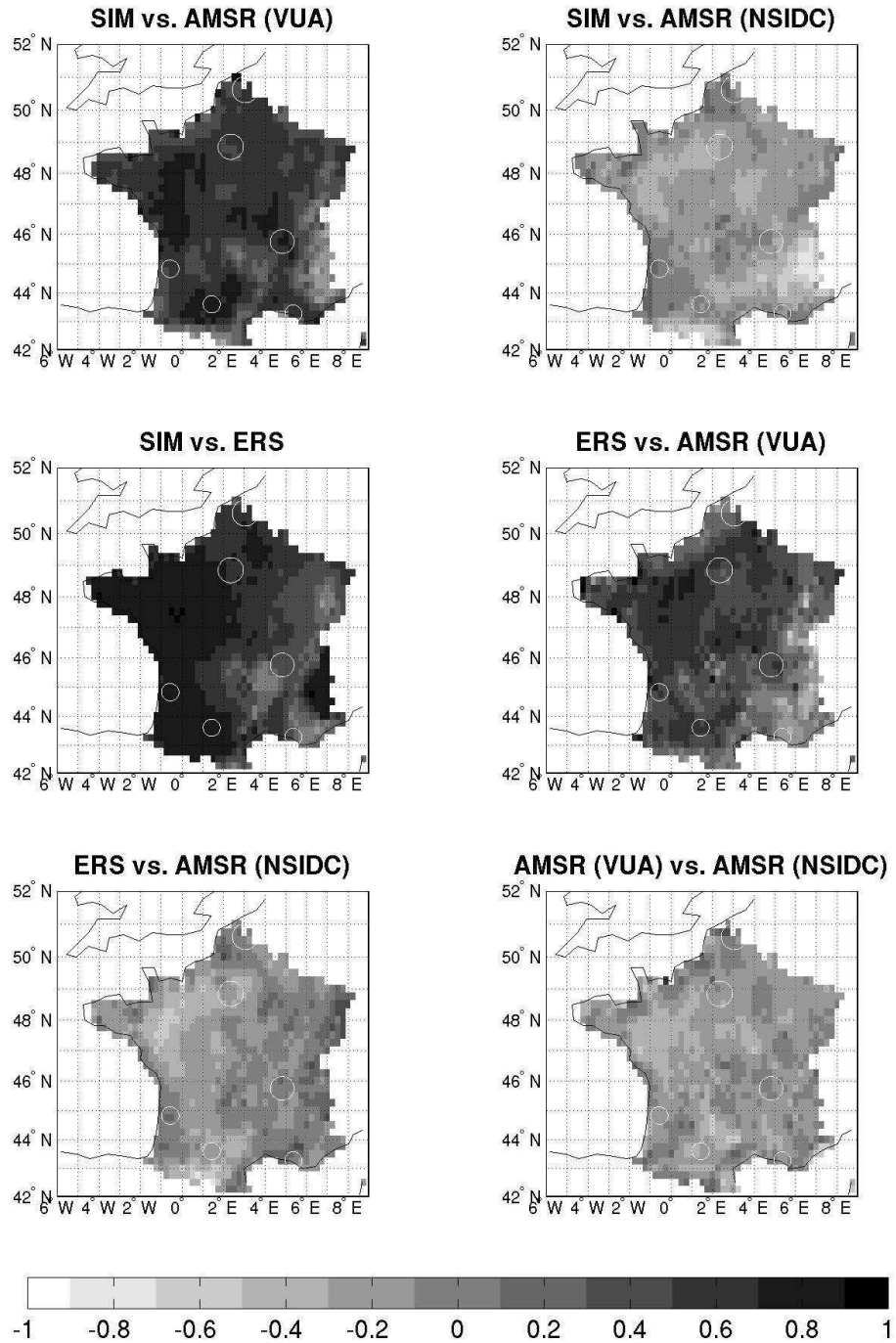


Fig. 4 Location of pixels with the different dominant land cover types (cultivated soils, grasslands, and forests), based on the fractional covers obtained from Ecoclimap and aggregated to 0.25° resolution.

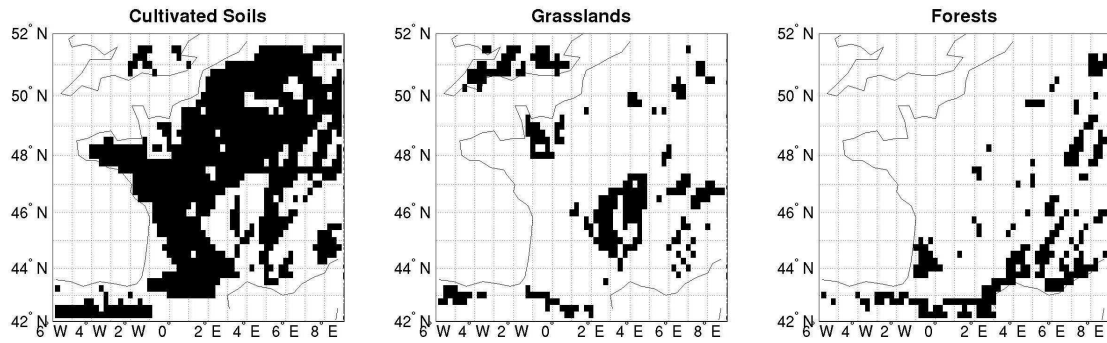


Fig. 5 Vegetation type specific comparison of the different soil moisture products for the three dominant vegetation types (a) cultivated soils, b) grasslands, c) forests), using the data from the period 2003-2005. The scatterplots and their corresponding statistics are located on opposite sides of each figure, ie. the scatterplot of the data pair SIM-AMSR (VUA) is in the top left hand corner, while the respective statistical values are found in the bottom right hand corner. Darker regions show a higher density of data points.

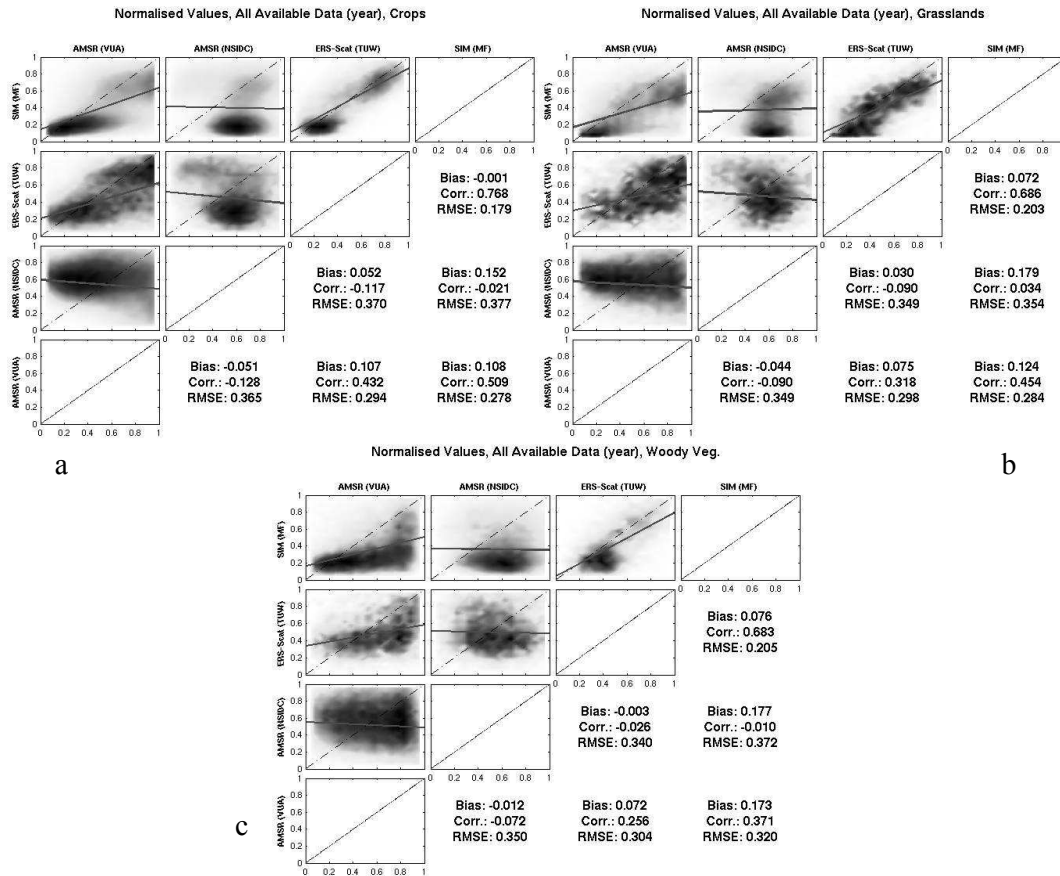


Fig 6. Histograms showing the relative frequency (vertical axis) of the various normalised soil moisture observations (horizontal axis) and predictions for the years 2003-2005 for the SMOSREX site: SIM model, AMSR-E product of VUA-NASA, AMSR-E product of NSIDC, ERS-Scat product of University of Vienna, in situ observations.

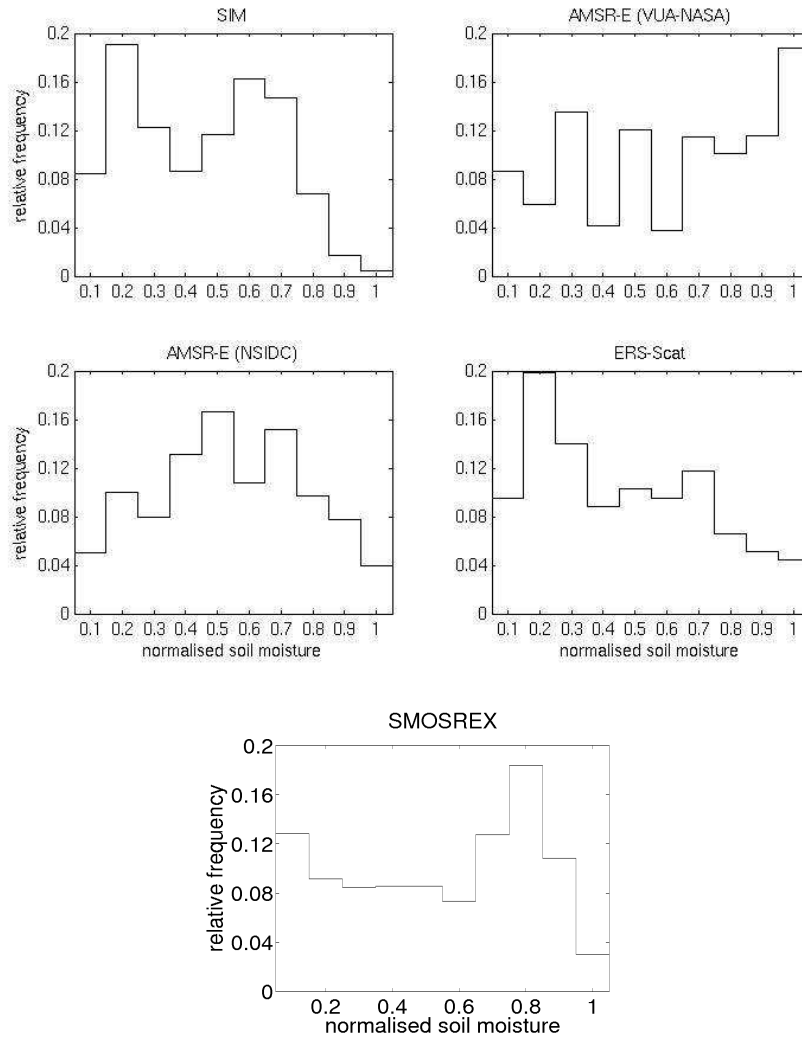


Fig 7. Histograms showing the relative frequency (vertical axis) of the various normalised soil moisture observations (horizontal axis) and predictions for the years 2003-2005 for whole of France: SIM model, AMSR-E product of VUA-NASA, AMSR-E product of NSIDC, ERS-Scat product of University of Vienna.

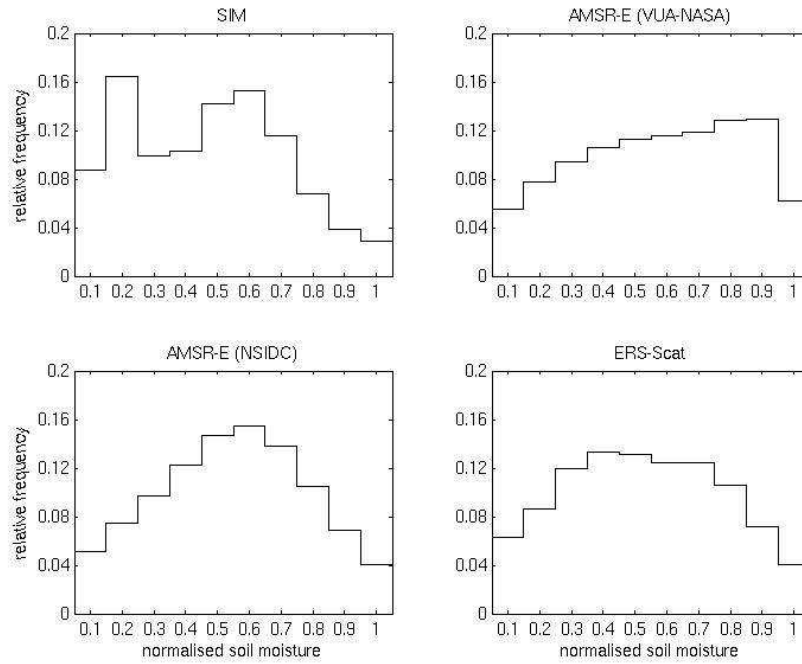


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