



Distributions of oxygen and carbon stable isotopes and CFC-12 in the water masses of the Southern Ocean at 30°E from South Africa to Antarctica: results of the CIVA1 cruise

Anne-Sophie Archambeau, Catherine Pierre, Alain Poisson, Bernard Schauer

► To cite this version:

Anne-Sophie Archambeau, Catherine Pierre, Alain Poisson, Bernard Schauer. Distributions of oxygen and carbon stable isotopes and CFC-12 in the water masses of the Southern Ocean at 30°E from South Africa to Antarctica: results of the CIVA1 cruise. *Journal of Marine Systems*, 1998, 17 (1-4), pp.25-38. 10.1016/S0924-7963(98)00027-X . ird-00145342

HAL Id: ird-00145342

<https://ird.hal.science/ird-00145342>

Submitted on 9 May 2007

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Distributions of oxygen and carbon stable isotopes and CFC-12 in the water masses of the Southern Ocean at 30°E from South Africa to Antarctica: results of the CIVA1 cruise

Anne-Sophie ARCHAMBEAU ¹, (tel : +33-1-4427-4868; Fax: +33-1-4427-4993; E-mail: archambo@ccr.jussieu.fr)

Catherine PIERRE ² (email : cat@lodyc.jussieu.fr)

Alain POISSON ¹ (email: apoisson@ccr.jussieu.fr)

Bernard SCHAUER ¹ (email: schauer@ccr.jussieu.fr)

- 1 Laboratoire de Physique et Chimie Marines (LPCM), (URA 2076: CNRS/UPMC), Université Pierre et Marie Curie, tours 24-25, 4 place Jussieu, F 75252 Paris Cedex 05, France
- 2 Laboratoire d'Océanographie Dynamique et de climatologie (LODyC), (UMR: CNRS/ORSTOM/UPMC), Université Pierre et Marie Curie, tour 26, 4 place Jussieu, F 75252 Paris Cedex 05, France

Abstract

This study presents oceanic distributions of stable isotopes ($\delta^{18}\text{O}$ of water and $\delta^{13}\text{C}$ of ΣCO_2) and CFC-12 from samples collected during the CIVA1 cruise (February/March 1993), across the Southern Ocean, along a meridian section at 30°E, from South Africa (44°S) to Antarctica (70°S). The isotopic measurements show important variations between the subantarctic surface waters with low $\delta^{18}\text{O}$ –high $\delta^{13}\text{C}$ values and the antarctic surface waters with very low $\delta^{18}\text{O}$ –low $\delta^{13}\text{C}$ values. The surface distributions of $\delta^{13}\text{C}$ values follow the major frontal oceanic structures; the vertical distribution shows the progressive upwelling from the subantarctic zone to the antarctic divergence of ^{13}C -depleted CO_2 derived from remineralization of organic matter. Along the Antarctic continental shelf, between 2500 and 4000 m, a core of water with $\delta^{18}\text{O}$ values close to -0.1‰ is associated with a relative maximum in CFC-12 concentration, although this core is not detected by its temperature and salinity parameters. This water mass, which corresponds to recently formed deep water, may originate from the eastward extension of the Weddell gyre or from bottom waters coming from the East and formed near Prydz Bay.

Résumé

Cette étude présente les distributions dans l'océan Austral des isotopes stables ($\delta^{18}\text{O}$ de l'eau et $\delta^{13}\text{C}$ de ΣCO_2) et des CFC-12 provenant d'échantillons prélevés durant la mission CIVA1 (Février/Mars 1993), le long d'une section méridienne à 30°E, entre l'Afrique du Sud (44°S) et l'Antarctique (70°S). Les compositions isotopiques montrent des variations importantes entre les eaux de surface subantarctiques caractérisées par des valeurs basses de $\delta^{18}\text{O}$ et élevées en $\delta^{13}\text{C}$ et les eaux de surface antarctiques très appauvries en ^{18}O et pauvres en ^{13}C . En surface, les valeurs de $\delta^{13}\text{C}$ varient en suivant les structures frontales océaniques; la distribution verticale de ce paramètre montre l'upwelling progressif de la zone subantarctique vers la divergence antarctique, de CO_2 appauvri en ^{13}C provenant de la reminéralisation de la matière organique. Le long de la marge continentale Antarctique, entre 2500 et 4000 m, des maxima relatifs des valeurs de $\delta^{18}\text{O}$ et de CFC-12 sont observés. Ils indiquent la présence d'une masse d'eau bien ventilée et récemment formée, bien qu'aucun indice ne soit révélé sur les profils de température et de salinité. Cette masse d'eau pourrait provenir soit de l'extension Est de

la gyre de Weddell soit des eaux profondes venant de l'est et formées dans la région de Prydz Bay.

Keywords: Stable isotopes; ^{18}O ; ^{13}C ; CFC-12

Introduction

The Southern Ocean is a special area from where the main bottom and intermediate water masses of the world ocean originate. Many studies using dynamic and geochemical observations in the Weddell Sea and near Prydz Bay provide information on the formation of deep and bottom waters and on their circulation in this area; however numerous questions remain regarding the eastern extension of the Weddell gyre and the recirculation of the bottom water masses (Gordon, 1971; Foster and Carmack, 1976; Jacobs and Georgi, 1977; Carmack, 1977; Deacon, 1979; Weiss et al., 1979; Smith et al., 1984; Jean-Baptiste et al., 1991; Mantsi et al., 1991; Schlosser et al., 1991; Gouretski and Danilov, 1993, Orsi et al., 1993). Measuring distributions of $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ values, which are natural tracers of oceanic water masses, is of great interest because they add specific information to help understand the mechanisms responsible for internal variability of water masses in their source region and at depth, and allow one to follow their circulation far away from their source.

The data presented here were collected during the WOCE/CIVA1 cruise on the Marion Dufresne during February/March 1993 in the southwest Indian sector of the Southern Ocean, along a section at 30°E , from South Africa (44°S) to Antarctica (70°S) (Fig. 1). This study presents a novel set of isotopic data ($\delta^{18}\text{O}$ of water and $\delta^{13}\text{C}$ of ΣCO_2) at 10 vertical stations throughout the water column. In addition to the stable isotope measurements, CFC-12 measurements from the CIVA1 cruise are also presented as transient tracers of the origin and age of the water masses (Bullister, 1984; Mantsi et al., 1991).

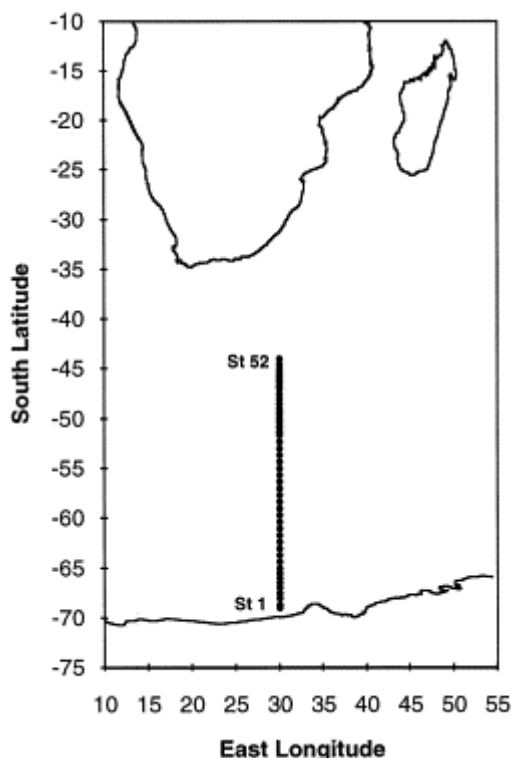


Fig. 1. Location of the hydrological stations sampled during the CIVA1 cruise.

Methods

Seawater samples were collected by a rosette system fitted with twelve 12-l Niskin bottles and a Neil-Brown CTD. Onboard, seawater samples for ^{13}C measurements were siphoned from Niskin bottles into 125 ml glass bottles, then poisoned with 1 ml of a saturated solution of HgCl_2 . The ΣCO_2 was extracted from seawater by acidification with phosphoric acid as described by Kroopnick (1974). The method used for $\delta^{18}\text{O}$ measurement was the classical water- CO_2 equilibration at 25°C as described by Epstein and Mayeda (1953). The isotopic compositions are expressed in the δ conventional units defined by:

$$\delta = [(R/R^*) - 1] \times 1000$$

where $R = ^{18}\text{O}/^{16}\text{O}$ or $^{13}\text{C}/^{12}\text{C}$ and R^* refers to the reference: SMOW for $^{18}\text{O}/^{16}\text{O}$ ratio (Craig, 1961) and PDB for $^{13}\text{C}/^{12}\text{C}$ ratio (Craig, 1957). Reproducibility of these measurements is about $\pm 0.05\text{‰}$ and the standard deviation is about $\pm 0.01\text{‰}$.

The method for CFC measurement follows the technique of Bullister and Weiss (1988) and was described in detail by Mantis et al. (1991). For CFC-12, the detection limits of the system were 0.01 pmol/kg and the standard deviation in the measurements is about 2%.

Hydrography

The CIVA1 cruise was conducted during the austral summer 1993 along a meridian section at 30°E which crosses the main hydrological fronts of the Southern Ocean (Lutjeharms and Valentine, 1984; Park et al., 1993). These fronts are identified by abrupt changes affecting the isohalines (Fig. 2a) and isotherms (Fig. 2b). Along this section, the different fronts are located as follows: the SubAntarctic Front (SAF) at 48°S , the Polar front (PF) at 52°S and the Antarctic Divergence (AD) at 66°S . The different water masses, identified by their θ - S parameters, are those described previously for the Antarctic Ocean by Jacobs and Georgi (1977) and Whitworth and Nowlin (1987).

In the Antarctic zone, the Antarctic Surface Water (AASW) shows strong variability in temperature and salinity ($-1^\circ\text{C} < \theta < 4^\circ\text{C}$; $33.7 < S < 34$) and is easily identified by a subsurface temperature minimum ($\theta \approx -1.7^\circ\text{C}$) which characterizes the remnant signal of the Winter Water (WW). Beneath this layer of cold water lies the Circumpolar Deep water (CDW) which is characterized by temperature and salinity maxima ($\theta \approx 1.4^\circ\text{C}$; $34.67 < S < 34.72$) between 1000 and 3000 m. For depths greater than 3500 m, the Antarctic Bottom Water (AABW) is observed with $S \approx 34.66$ and $\theta \approx -0.7^\circ\text{C}$.

In the subantarctic zone, four main water masses are superimposed. The Subantarctic Surface Waters (SASW) are relatively warm with a large range of temperatures ($8^\circ\text{C} < \theta < 14^\circ\text{C}$); the Antarctic Intermediate Water (AAIW) which is characterized by a salinity minimum ($S = 34.25$), sinks at the SAF and lies at immersion depths from 200 m to 800 m; the CDW, which contains a major contribution of North Atlantic Deep Water (NADW) in this area, shows a temperature maximum ($\theta = 2.6^\circ\text{C}$) near 1600 m and a salinity maximum ($S = 34.76$) near 2500 m; the AABW is still observed.

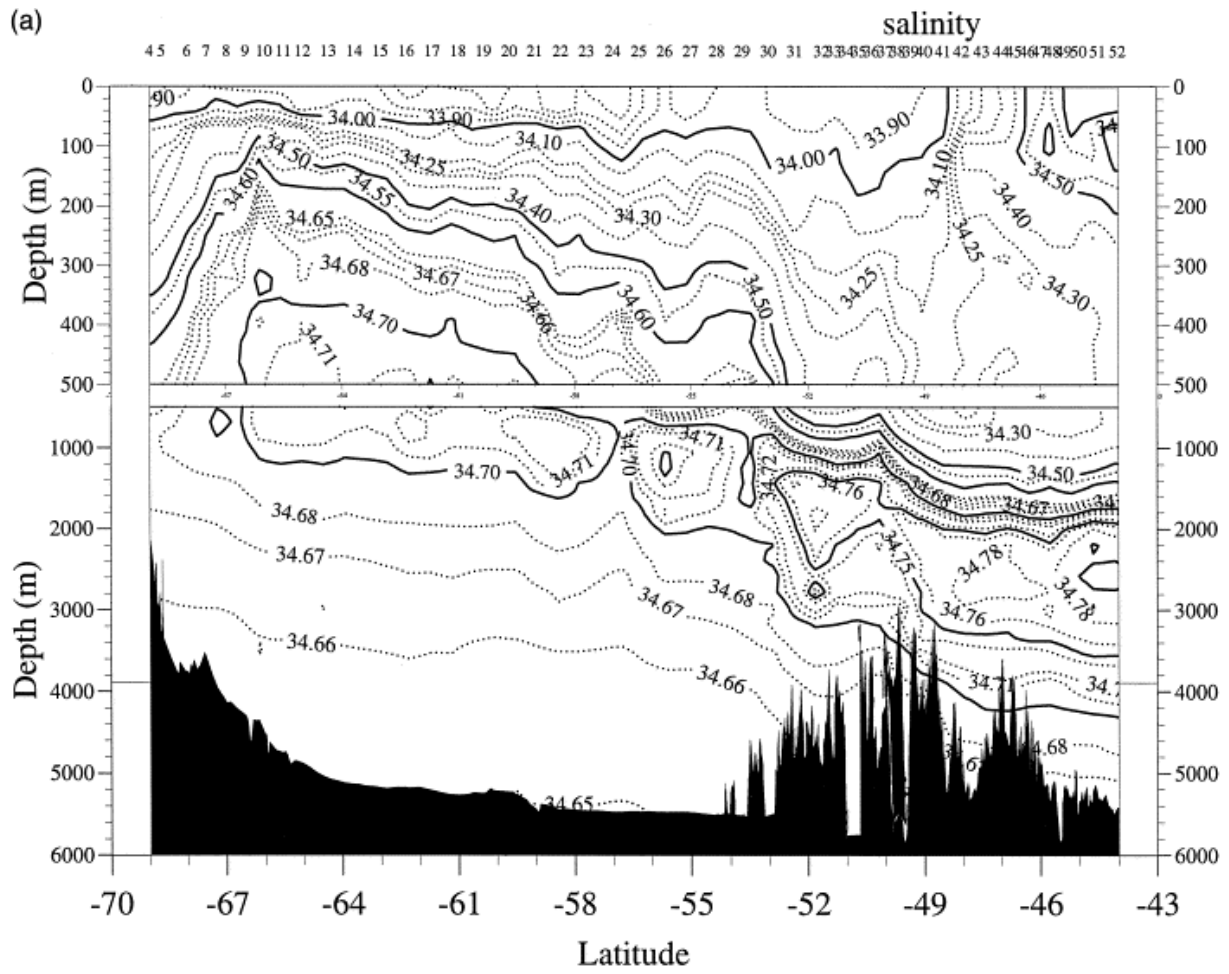


Fig. 2. (a) Latitudinal distribution of salinity

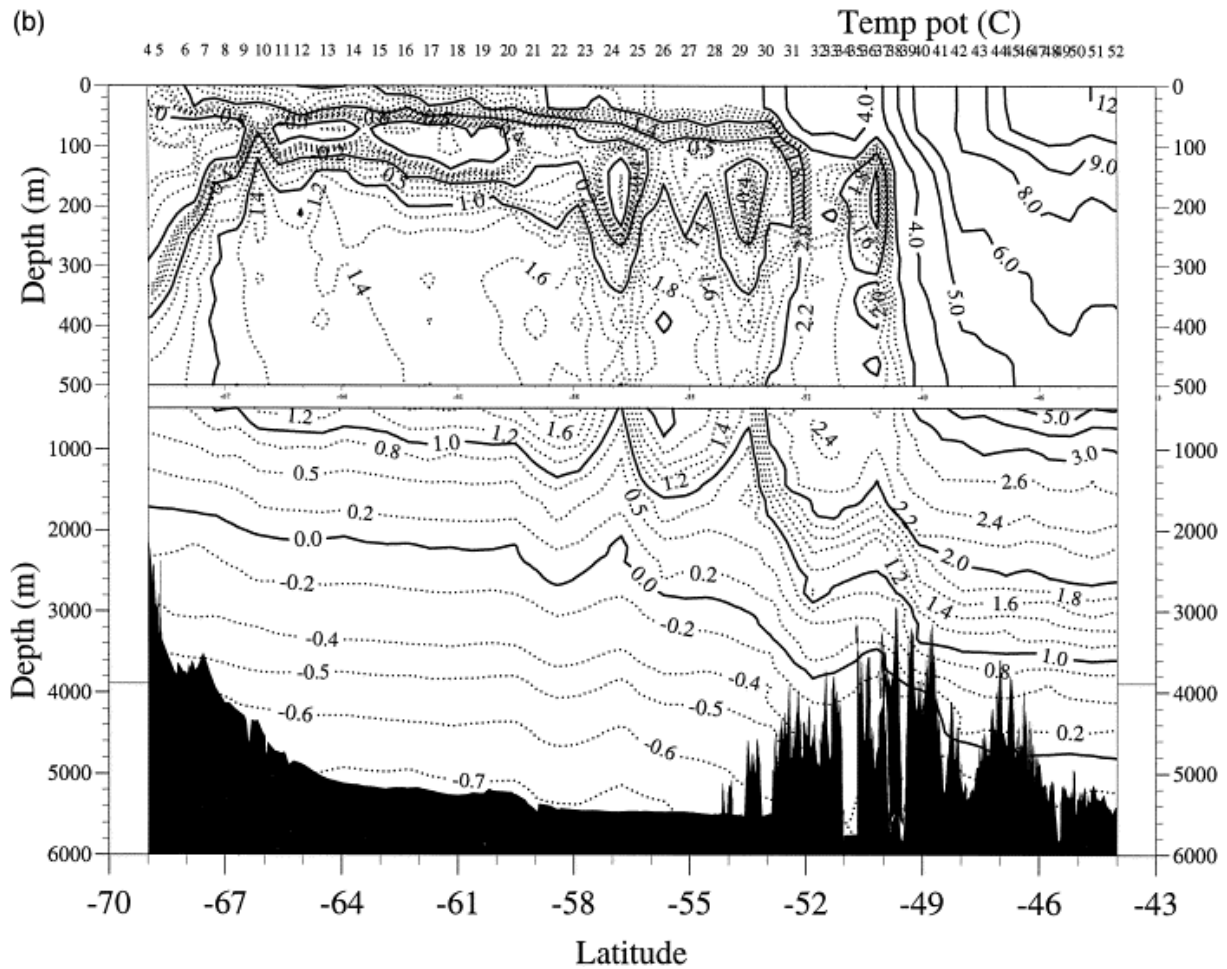


Fig. 2. (b) Latitudinal distribution of potential temperature (°C).

The chlorofluorocarbons (CFC) distribution in the water masses

The vertical distribution of CFC-12 along the section at 30°E (Fig. 3) gives information on the penetration of the transient tracers in this area of the Southern Ocean as well as on the global circulation at depth. The penetration of water with high CFC values in the subantarctic zone is observed down to 1000 m; this characterizes the sinking of newly formed AAIW (CFC-12=1.0 pmol/kg) at the SAF from 200 m to intermediate depths (800 m). The CDW, which largely extends in this area, is characterized by low CFC concentrations ($0.05 \text{ pmol/kg} < \text{CFC-12} < 0.10 \text{ pmol/kg}$); its upwelling at the AD stops the penetration of atmospheric anthropogenic CFCs in the antarctic zone.

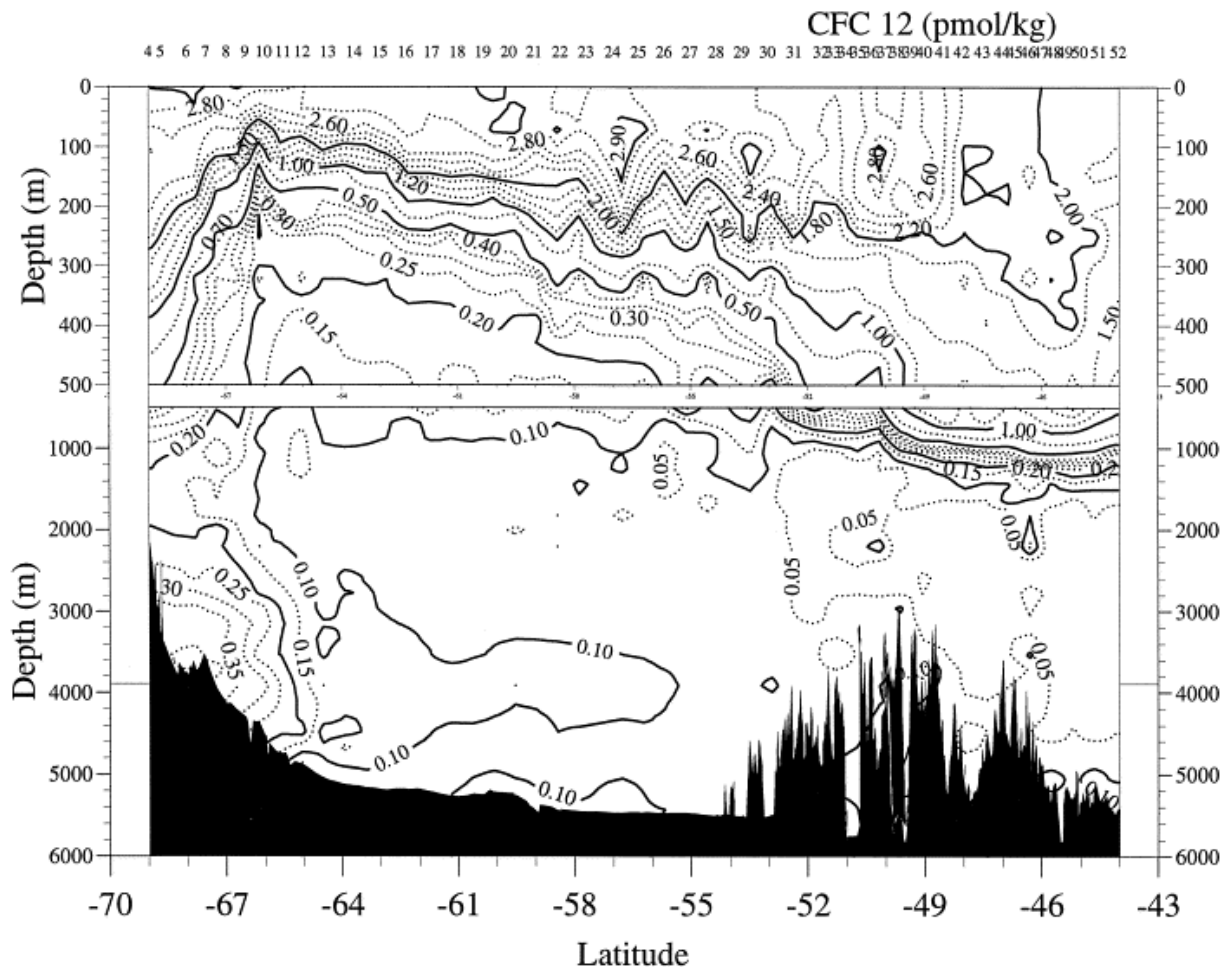


Fig. 3. Latitudinal distribution of CFC-12 (pmol/kg).

Along the Antarctic continental shelf, a relative maximum of CFC concentration is observed between 2000 and 4500 m ($0.2 \text{ pmol/kg} < \text{CFC-12} < 0.4 \text{ pmol/kg}$), indicating that these waters were recently ventilated. Between the AD and 55°S , at an immersion depth of about 4000 m, a tongue of CFC rich water ($0.1 \text{ pmol/kg} < \text{CFC-12} < 0.15 \text{ pmol/kg}$) indicates the existence of a "young" deep water mass. Another relative maximum occurs near the bottom, at 5000 m ($0.1 \text{ pmol/kg} < \text{CFC-12} < 0.15 \text{ pmol/kg}$) which is correlated with a temperature minimum ($\theta = -0.7^\circ\text{C}$) and an oxygen maximum ($240 < \text{O}_2 \mu\text{m/kg} < 245$).

The stable isotope composition of the water masses

The oxygen and carbon stable isotope compositions of ocean waters provide specific information on both the water and carbon cycles, providing insight to better understand the mechanisms and magnitude of the fluxes between the oceanic and atmospheric reservoirs.

The $\delta^{18}\text{O}$ values of ocean waters are sensitive to variations in the evaporation-dilution ratio; the concentrations of heavier isotopes in the ocean surface waters are enriched by evaporation and depleted by dilution due to precipitation. These processes cause linear relationships between oxygen-18 and salinity whose slopes depend on climatic constraints (Craig and Gordon, 1965). At high latitudes, the effect of freezing at the ocean surface increases the salinity, however it causes only a weak depletion of the

$\delta^{18}\text{O}$ values of the remaining liquid seawater. Leaving the surface, for intermediate or deeper depths, the water is no longer affected by air–sea exchange; the $\delta^{18}\text{O}$ parameter is thus conservative and its distribution follows the circulation pattern of the water masses (Craig and Gordon, 1965). The melting of sea-ice causes large changes in salinity and leaves the $\delta^{18}\text{O}$ almost unchanged (the resulting mixture may be slightly enriched in ^{18}O), while continental-ice melt causes large drops both in salinity and $\delta^{18}\text{O}$. Therefore, at high latitudes, the $\delta^{18}\text{O}$ values of waters may also help to differentiate dilution effects induced by the melting of continental ice as opposed to sea ice (Schlosser et al., 1990).

The variations of $\delta^{13}\text{C}$ of ΣCO_2 occur mostly in the upper ocean waters and depend on the ocean–atmosphere CO_2 exchange and on the photosynthesis–respiration processes. Carbon isotopes are fractionated during the production of organic matter which incorporates preferentially the carbon-12. The $\delta^{13}\text{C}$ values in ocean waters are maximum at the surface where photosynthesis occurs; they decrease at depth where remineralization of organic matter releases ^{13}C -depleted CO_2 (Kroopnick, 1985; see Table 1 Table 2).

Table 1: Stable isotope composition ($\delta^{18}\text{O}$ of water and $\delta^{13}\text{C}$ of ΣCO_2) and salinity data for the CIVA1 surface samples

Station	Latitude (°S)	Longitude (°E)	Salinity	θ (°C)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)
1	68.98	30.13	33.4524	−1.6376	−0.42	0.91
2	68.98	30.08	33.4097	−1.5796	−0.59	1
3	68.99	29.98	33.3987	−1.3862	−0.38	–
4	68.98	29.85	33.715	−0.7996	−0.56	–
5	68.67	29.96	33.7513	−1.0749	−0.46	1.05
6	68.01	30	33.9122	0.2009	−0.53	0.97
8	66.99	29.98	33.8657	0.646	−0.47	0.54
10	66.01	29.99	33.8424	0.8639	−0.61	0.81
11	65.51	29.98	33.8382	0.978	–	0.68
12	65.01	30	33.8243	1.004	−0.47	0.76
13	64.35	30	33.8003	1.106	–	0.81
14	63.68	30	33.8285	1.026	−0.46	0.96
15	63.01	29.99	33.7685	1.266	–	0.87
16	62.34	30	33.5923	1.6909	–	1
17	61.67	30	33.7085	1.551	–	0.92
18	61	30	33.6881	1.7759	−0.49	0.88
19	60.34	29.99	33.6306	1.779	–	0.89

Station	Latitude (°S)	Longitude (°E)	Salinity	Θ (°C)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)
20	59.68	29.99	33.717	1.8669	–	0.88
21	59.02	30	33.7054	1.8109	–	0.75
22	58.35	29.97	33.5976	2.1938	–0.46	–
23	57.68	29.99	33.7095	2.1798	–	0.78
26	55.68	30	33.9342	2.8325	–0.44	1.04
27	55.01	30	33.8169	2.3613	–	1.36
29	53.68	29.98	33.9538	2.7055	–	0.86
30	53.01	30	33.8815	3.1221	–0.51	–
31	52.33	30.02	33.8802	4.177	–	1.19
32	51.64	29.97	33.8812	4.6461	–0.28	1.32
34	51	29.97	33.8913	4.6703	–0.47	1.21
36	50.35	29.99	33.8451	3.1573	–0.27	–
38	49.68	30.02	33.8755	4.5363	–0.35	0.99
40	49	30.01	33.8919	6.3907	–	1.23
42	48.03	30	34.1661	8.3457	–0.24	1.74
43	47.51	29.98	34.2679	8.8479	–	1.18
44	47.01	30	34.3512	10.1607	–0.1	1.14
45	46.67	30.02	34.2965	9.93	–	1.32
46	46.34	29.98	34.5306	10.8762	–	1.61
47	46	29.99	34.4394	10.8862	–	1.26
48	45.68	29.99	34.6698	11.5325	–0.03	1.33
49	45.38	30	34.4968	10.8949	–	1.47
50	45.01	29.98	34.4389	10.7013	–	1.18
51	44.52	29.98	34.4487	12.0471	–	1.6
52	44.01	29.96	34.4794	12.315	0.04	1.56

Table 2 : CFC-12, oxygen-18 and carbon-13 data for CIVA1 seawater samples. Each sample is listed by station number and depth, together with its potential temperature (θ) and salinity (S)

Depth (m)	θ ($^{\circ}\text{C}$)	S	F12 (pmol/kg)	$\delta^{18}\text{O}$ (‰)	$\delta^{13}\text{C}$ (‰)
<i>Station 5 (68.67°S 29.96°E)</i>					
10.1	-1.0749	33.7513	2.94	-0.24	-1.05
20.6	-0.1286	33.8975	2.84	-0.46	1.35
39.1	-0.3232	33.9091	2.82	-0.38	1.24
60.7	-1.6306	34.1131	2.60	-0.48	0.91
80.1	-1.7271	34.1751	2.47	-0.49	–
100.3	-1.7986	34.2010	2.53	-0.53	0.83
198.4	-1.5708	34.2888	2.38	-0.30	0.71
325.5	-0.8186	34.4444	1.67	-0.31	0.57
399.3	0.1062	34.5601	0.99	-0.31	0.38
498.5	0.6585	34.6434	0.55	-0.17	0.35
697.1	0.6859	34.6785	0.34	-0.25	0.31
998.8	0.4346	34.6827	0.21	–	0.31
1345.8	0.1750	34.6735	0.18	-0.10	0.32
1748.2	-0.0280	34.6678	0.18	-0.15	0.25
1996.0	-0.1251	34.6670	0.20	–	0.29
2496.4	-0.2797	34.6622	0.28	-0.23	–
3021.7	-0.4232	34.6590	0.38	-0.08	0.29
<i>Station 6 (68.01°S 30.00°E)</i>					
10.2	0.2009	33.9122	2.92	-0.21	0.97
19.9	0.1145	33.9158	2.93	-0.53	1.21
40.0	-0.1182	33.9307	2.86	-0.44	1.27
60.7	-1.7257	34.2924	2.50	-0.51	0.86
81.0	-1.7582	34.3304	2.50	-0.11	0.71
100.6	-1.7636	34.3487	2.47	-0.16	0.80
199.7	-0.5536	34.4732	1.52	0.07	0.60

324.4	0.7518	34.6407	0.46	0.04	0.17
403.2	0.8128	34.6642	0.42	-0.03	-0.02
698.8	0.7136	34.6916	0.23	0.08	0.26
998.8	0.4484	34.6879	0.18	-0.15	0.08
1346.3	0.2066	34.6780	0.13	-0.13	0.34
1747.9	-0.0171	34.6713	0.15	-0.15	0.37
1996.4	-0.1300	34.6672	0.19	-	0.33
2497.6	-0.2729	34.6630	0.24	-0.25	0.38
2997.6	-0.3780	34.6601	0.36	-0.07	0.41
3495.9	-0.4769	34.6562	0.41	-0.06	0.41
3681.3	-0.5667	34.6538	0.40	-0.01	-
<i>Station 8 (66.99°S 29.98°E)</i>					
10.6	0.6460	33.8657	2.89	0.19	-
20.2	0.6406	33.8655	2.84	-0.47	-
40.0	-1.505	34.2558	2.46	-0.52	-
61.0	-1.6737	34.3407	2.35	-0.45	-
80.8	-1.7041	34.3543	2.36	-0.55	-
100.0	-1.6465	34.3737	2.24	-0.40	-
201.0	0.7616	34.6069	0.61	-0.18	-
322.1	1.1479	34.6803	0.29	-0.27	-
397.5	1.1609	34.6970	0.22	-0.22	-
493.7	1.1169	34.7074	0.16	0.08	-
699.7	0.8876	34.7082	0.15	-0.12	-
997.1	0.5854	34.6986	0.09	-0.24	-
1300.3	0.2866	34.6845	0.16	-0.23	-
1299.9	0.2877	34.6836	0.15	-0.37	-
1595.7	0.1155	34.6753	0.13	-0.40	-
1993.8	-0.0865	34.6692	0.25	-0.24	-
2490.9	-0.2596	34.6639	0.29	0.00	-
3193.9	-0.4247	34.6577	0.39	-0.37	-
3697.2	-0.5206	34.6547	0.42	-0.44	-

4065.9	-0.6265	34.6518	0.36	-0.48	–
4065.7	-0.6255	34.6508	0.36	-0.56	–
<i>Station 10 (66.01°S 29.99°E)</i>					
13.2	0.8639	33.8424	2.86	-0.56	0.81
23.3	0.8255	33.8495	2.84	-0.61	–
82.4	0.4151	34.5224	1.09	0.10	–
102.0	0.8490	34.5760	0.65	-0.21	–
202.8	1.5142	34.6852	0.20	-0.12	–
324.3	1.4360	34.7080	0.14	-0.16	–
399.1	1.3432	34.7120	0.17	-0.34	–
500.3	1.2390	34.7171	0.11	-0.10	–
698.2	1.0239	34.7164	0.11	-0.21	–
998.1	0.6925	34.7047	0.11	-0.34	–
1350.8	0.3953	34.6916	0.11	-0.32	–
1350.3	0.3963	34.6908	0.09	-0.32	–
1495.7	0.3028	34.6841	0.07	-0.33	–
1995.4	0.0335	34.6748	0.11	-0.22	–
2495.4	-0.1756	34.6665	0.17	-0.39	–
2994.9	-0.3338	34.6613	0.27	-0.39	–
3495.1	-0.4408	34.6601	0.22	-0.34	–
3998.4	-0.5312	34.6553	0.37	-0.37	–
3998.0	-0.5321	34.6552	0.30	-0.58	–
4477.0	-0.6604	34.6524	0.23	-0.44	–
<i>Station 14 (63.68°S 30.00°E)</i>					
11.0	1.0260	33.8285	2.87	–	0.96
20.8	0.7816	33.8316	2.89	-0.46	1.10
40.5	-0.8856	34.0561	2.76	-0.55	1.07
60.2	-1.5886	34.1637	2.69	-0.41	0.01
79.9	-1.6670	34.1966	2.62	-0.57	1.03
99.5	-1.0753	34.2850	2.22	-0.32	0.79
202.0	1.3446	34.6453	0.35	-0.02	0.02

324.4	1.3472	34.6891	0.20	-0.21	0.23
399.3	1.3442	34.7070	0.16	-0.33	0.19
498.7	1.1963	34.7095	0.15	-0.17	0.26
691.6	1.0014	34.7124	0.11	-0.07	0.38
998.0	0.6657	34.7026	0.09	-0.07	0.19
1347.9	0.3994	34.6928	0.09	-0.27	0.18
2495.2	-0.1687	34.6669	0.11	-0.38	0.25
3493.5	-0.4415	34.6596	0.14	-0.44	0.64
3994.2	-0.5385	34.6572	0.15	-0.41	0.39
3995.0	-0.5386	34.6568	0.14	-0.41	0.30
4491.9	-0.6271	34.6545	0.08	-0.37	–
5092.8	-0.7216	34.6514	0.08	-0.40	0.40
<i>Station 18 (61.00°S 30.00°E)</i>					
11.3	1.7759	33.6881	2.83	-0.36	0.88
20.9	1.6254	33.7183	2.85	-0.49	–
80.0	-1.6200	34.1229	2.74	-0.39	–
100.1	-1.6305	34.1397	2.66	-0.03	–
200.0	1.2090	34.5679	0.62	0.04	–
324.4	1.5217	34.6755	0.23	0.05	–
398.9	1.5356	34.7015	0.17	-0.14	–
498.8	1.4363	34.7135	0.10	-0.03	–
698.1	1.2098	34.7208	0.08	0.03	–
998.9	0.8452	34.7100	0.10	0.01	–
1346.3	0.5005	34.6957	0.09	0.01	–
1348.3	0.5003	34.6947	0.07	0.09	–
1495.7	0.3958	34.6901	0.07	-0.16	–
1994.7	0.1193	34.6790	0.08	-0.05	–
2494.0	-0.1077	34.6706	0.09	-0.22	–
2995.5	-0.2888	34.6626	0.10	0.05	–
3495.4	-0.4173	34.6600	0.08	-0.42	–
3993.6	-0.5189	34.6565	0.13	-0.46	–

5228.2	-0.7164	34.6515	0.10	-0.40	–
<i>Station 22 (58.35°S 29.97°E)</i>					
9.7	2.1938	33.5976	2.80	-0.47	1.41
19.9	2.1963	33.5974	2.78	-0.46	1.46
40.4	2.1712	33.6051	2.76	-0.39	1.43
62.4	-0.2850	33.9609	2.95	-0.52	1.24
82.4	-0.6427	34.0512	2.89	-0.43	0.78
101.5	-0.3753	34.1183	2.76	-0.13	0.83
201.7	0.6828	34.3620	1.57	-0.02	0.51
320.1	1.4822	34.5507	0.58	-0.20	0.30
396.6	1.6254	34.6269	0.34	0.13	0.22
497.7	1.8556	34.6953	0.13	-0.09	0.31
696.3	1.7796	34.7376	0.07	–	0.32
994.9	1.3741	34.7341	0.05	-0.10	0.30
1349.8	0.9444	34.7215	0.09	-0.09	–
1349.3	0.9464	34.7208	0.05	-0.05	0.27
1494.2	0.7931	34.7122	0.06	-0.18	0.27
1998.2	0.3760	34.6897	0.07	-0.18	0.17
2497.0	0.0873	34.6759	0.08	-0.28	–
3000.1	-0.1513	34.6696	0.07	-0.07	0.33
3496.3	-0.3267	34.6598	0.11	-0.22	0.32
3998.1	-0.4552	34.6574	0.11	-0.18	0.42
5371.4	-0.7105	34.6495	0.11	-0.42	0.36
<i>Station 26 (55.68°S 30.00°E)</i>					
10.2	2.8325	33.9342	2.67	-0.50	1.04
19.4	2.8249	33.9352	2.67	-0.44	1.48
38.7	2.8278	33.9362	2.67	-0.41	1.48
60.0	2.4319	33.9566	2.71	-0.48	1.33
79.8	0.8511	34.0314	2.75	-0.21	1.23
100.0	0.6233	34.0809	2.58	-0.02	1.10
199.9	1.4017	34.2946	1.27	0.04	0.69

322.6	1.9119	34.4610	0.57	0.03	0.43
392.3	2.0951	34.5489	0.29	0.16	0.30
503.4	2.1331	34.6271	0.15	-0.05	0.31
699.1	2.0807	34.7022	0.07	0.07	0.28
995.4	1.8621	34.7517	0.00	—	0.34
1343.8	1.4918	34.7541	0.04	0.12	0.33
1745.6	0.9319	34.7202	0.06	—	0.35
2244.9	0.4002	34.6911	0.07	0.09	0.26
2994.6	-0.0833	34.6703	0.08	—	0.28
3495.2	-0.2934	34.6633	0.09	0.12	0.31
3994.5	-0.4334	34.6585	0.11	—	0.36
5441.6	-0.7160	34.6503	0.10	-0.06	0.35
<i>Station 32 (51.64°S 29.97°E)</i>					
10.9	4.6461	33.8812	2.46	-0.27	1.32
20.2	4.6394	33.8824	2.38	-0.24	1.32
37.7	4.6391	33.8815	2.47	-0.42	1.59
60.9	4.6214	33.8822	2.37	-0.23	1.69
79.2	3.9949	33.9012	2.44	-0.20	1.51
99.3	2.9470	33.9513	2.58	-0.27	1.26
200.9	2.7721	34.1327	1.88	-0.12	0.99
325.9	2.4452	34.2137	1.31	-0.23	0.79
398.6	2.4088	34.2755	0.91	-0.07	0.67
499.8	2.3705	34.3515	0.57	-0.11	0.62
700.5	2.4456	34.5183	0.21	0.01	0.43
996.9	2.3759	34.6628	0.05	0.05	0.39
1345.0	2.2820	34.7661	0.04	-0.05	—
1349.1	2.2678	34.7643	0.03	0.05	0.43
1496.9	2.2403	34.7791	0.00	0.03	0.43
2012.0	1.6977	34.7807	0.00	0.03	0.51
2404.7	1.3461	34.7611	0.03	-0.04	0.42
2746.6	1.2286	34.7657	0.00	0.01	0.54

2994.3	0.8648	34.7290	0.00	-0.08	0.46
3495.4	0.3224	34.6895	0.06	0.02	0.36
4854.7	-0.3832	34.6611	0.08	-0.19	0.50
<i>Station 42 (48.03°S 30.00°E)</i>					
7.3	8.3457	34.1661	2.16	-0.28	1.74
18.2	8.3456	34.1642	2.17	-0.24	1.74
40.0	8.3484	34.1660	2.16	-0.40	1.74
58.3	8.3495	34.1671	2.16	-0.23	1.73
76.9	8.5443	34.3887	2.07	–	1.50
97.7	8.2066	34.4798	2.01	-0.11	–
198.5	6.4745	34.3083	2.07	-0.29	1.30
323.4	5.3758	34.2373	1.42	-0.27	1.18
399.0	5.1701	34.2882	–	-0.29	1.04
499.0	4.7635	34.3022	1.11	-0.35	1.19
698.3	3.2969	34.2365	1.01	-0.32	0.89
1097.6	2.8633	34.4643	0.22	–	0.56
1351.0	2.5481	34.5625	0.08	-0.21	0.44
1744.4	2.3917	34.6950	0.00	-0.20	–
1992.1	2.2862	34.7487	0.05	–	0.45
2244.3	2.1168	34.7731	–	-0.02	–
2745.4	1.7741	34.7923	0.04	0.01	–
2993.6	1.5461	34.7783	0.05	0.01	0.42
3498.3	0.9699	34.7333	0.00	0.00	0.33
3993.9	0.5506	34.7090	0.04	-0.09	–
5060.8	-0.2095	34.6668	0.09	–	0.32
<i>Station 52 (44.01°S 29.96°E)</i>					
9.8	12.3150	34.4794	1.82	0.03	1.56
19.4	12.2968	34.4783	1.87	0.04	1.59
39.7	12.0453	34.4526	1.95	0.02	1.56
61.0	11.8357	34.6558	1.86	-0.12	–
82.1	10.0045	34.6431	1.86	-0.13	–

100.5	9.2985	34.5982	1.89	0.09	–
197.4	7.9160	34.4974	1.75	0.09	0.96
323.3	6.3703	34.3396	1.62	–0.18	1.07
398.5	5.8962	34.3536	1.32	–0.02	1.04
499.0	5.1576	34.3321	1.09	–0.05	0.79
698.7	4.0819	34.3287	0.80	–0.18	0.79
1098.2	2.7701	34.4644	0.25	–0.17	0.49
1345.5	2.6602	34.5998	0.12	–0.07	–
1347.4	2.6561	34.6007	0.11	–	0.40
1748.2	2.4683	34.7234	0.00	–0.03	0.27
2244.9	2.2081	34.7934	0.00	0.08	0.58
2995.4	1.6311	34.7894	0.00	0.00	0.58
3496.4	1.0783	34.7521	0.00	–0.03	0.44
4497.9	0.1862	34.6903	0.00	–	0.37
5002.0	–0.1482	34.6705	0.09	–0.18	0.40
5440.7	–0.2290	34.6660	0.09	–0.15	–

The oxygen isotope composition

In the uppermost surface waters, the $\delta^{18}\text{O}$ values and salinities show latitudinal gradients (Fig. 4). North of the Subantarctic Front to the Polar Front, $\delta^{18}\text{O}$ values decrease from 0.0‰ to –0.4‰ and salinities drop from 34.7 to 33.9 due to dilution by atmospheric precipitations. In the antarctic zone, the $\delta^{18}\text{O}$ values of AASW remains around –0.45‰ while salinities are in the range 33.6–33.9 due to the melting of sea ice. At the AD, a slight drop in the $\delta^{18}\text{O}$ values to –0.6‰ associated with a very small salinity decrease is probably indicative of a very weak contribution from continental ice melting. South of 69°S, the decrease of $\delta^{18}\text{O}$ value which is associated with a salinity increase marks the effect of sea-ice freezing.

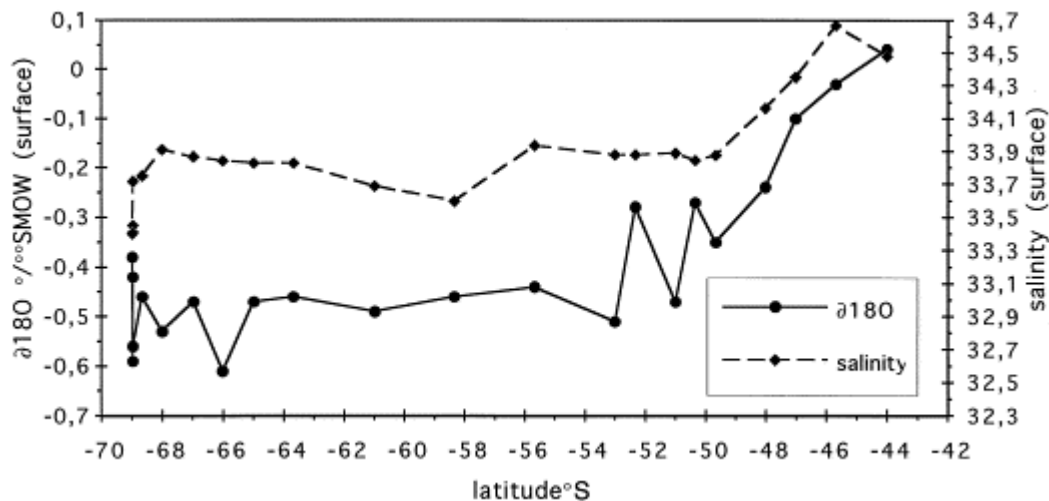


Fig. 4. Latitudinal variations of observed $\delta^{18}\text{O}$ and salinity values in the uppermost surface waters of the Southern Ocean.

On the $\delta^{18}\text{O}$ -S diagram (Fig. 5), the signal of Winter Waters (WW) is well defined by its salinity (34.15 to 34.30) and $\delta^{18}\text{O}$ value (-0.56‰ to -0.40‰). The other water masses are also differentiated by their $\delta^{18}\text{O}$ -S parameters: $-0.15\text{‰} < \delta^{18}\text{O} < -0.35\text{‰}$ and $34.22 < S < 34.34$ for AAIW; $-0.18\text{‰} < \delta^{18}\text{O} < 0.15\text{‰}$ and $34.70 < S < 34.80$ for CDW, the ^{18}O and salinity enrichments being due to the influence of North Atlantic Deep Water (NADW) in the northern part of the section; $-0.28\text{‰} < \delta^{18}\text{O} < -0.42\text{‰}$ and $34.64 < S < 34.68$ for AABW. This diagram shows a grouping of waters identified as CSBW (Continental Shelf Bottom Water) which have the same salinity as AABW ($S=34.66$) but higher $\delta^{18}\text{O}$ values (-0.2 to 0.0‰).

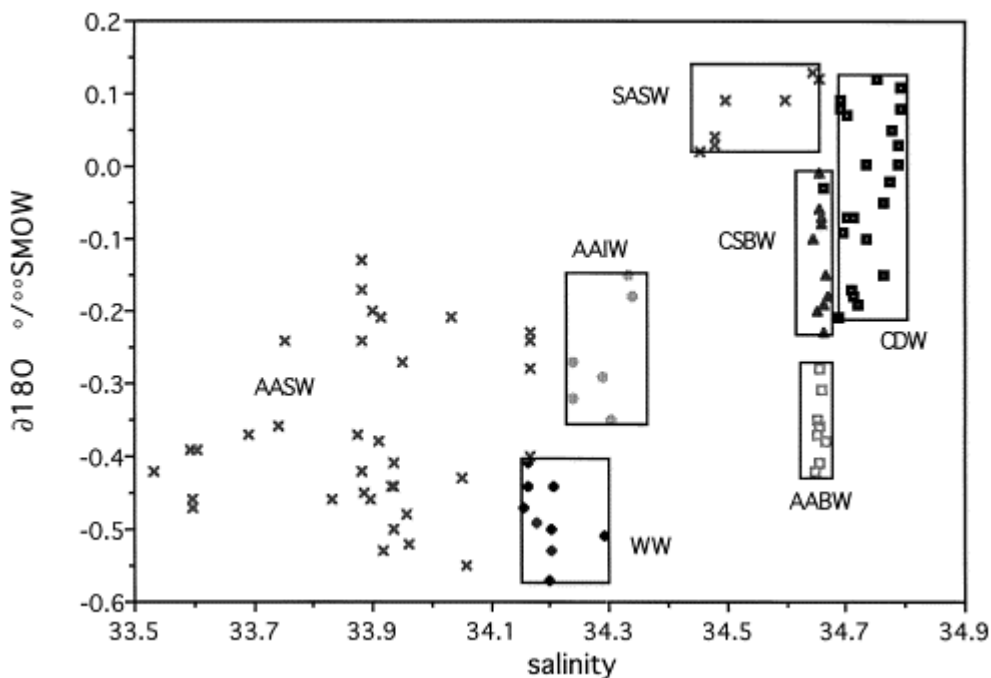
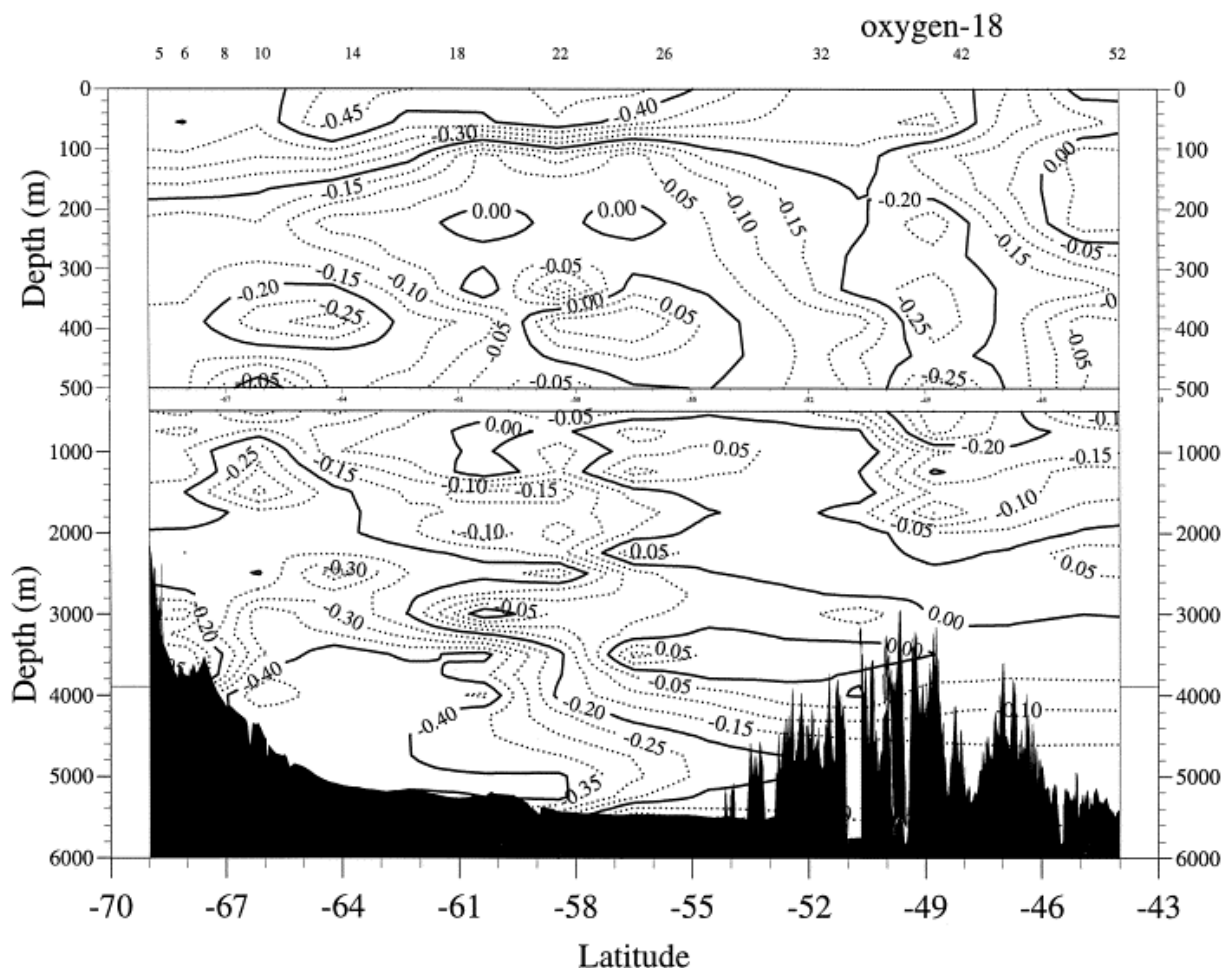


Fig. 5. The $\delta^{18}\text{O}$ /salinity relationship in the water masses of the Southern Ocean. The following water masses have been identified: AASW, Antarctic Surface Water; SASW, SubAntarctic Surface Water; WW, Winter Water; CDW, Circumpolar Deep Water; AABW, Antarctic Bottom Water; CSBW, Continental Shelf Bottom Water.

The vertical distributions of $\delta^{18}\text{O}$ values along the 30°E section (Fig. 6) follows rather well the general pathways of the main water masses. The AAIW, characterized by $\delta^{18}\text{O}$ values between -0.2 and -0.3‰ when it sinks at the Polar Front, is advected northward at a depth of about 800 m. The core of CDW is well recognized by $\delta^{18}\text{O}$ values close to 0.0‰ between 2000 and 3000 m; the upwelling of this water mass at the antarctic divergence is marked by the raising of the $\delta^{18}\text{O}=0\text{‰}$ isoline at immersion depths of about 1000 m. The AABW, characterized by $\delta^{18}\text{O}$ values around -0.4‰ , lies below 3500 m. Along the southern margin of Antarctica, a lense of bottom water with maximum values of $\delta^{18}\text{O}\approx 0.1\text{‰}$ is observed between 2500 m and 4000 m. This water mass corresponding to the CSBW, which is not differentiated by its temperature and salinity characteristics, has also been identified by a relative maximum in the CFC-12 concentrations. There is no evidence of local formation of deep water in this area so two hypothesis may be put forward: (1) the water mass comes from the East and corresponds to bottom waters formed near Prydz Bay, (2) the water is formed in the Weddell sea and acquires an intermediate density between the WSBW and the CDW; then, following the topography, this water flows back to the Weddell basin by the Enderby basin. However, further physical measurements would be necessary for a better understanding of the origin and circulation of this deep water mass in this area.



The carbon isotope composition

The variations of $\delta^{13}\text{C}$ values in the uppermost surface waters (Fig. 7) follow the main frontal oceanic structures. Strong oscillations are observed north of the Polar Front, separating the ^{13}C -rich waters ($1.15\text{‰} < \delta^{13}\text{C} < 1.6\text{‰}$) of the polar front and subantarctic zones from the ^{13}C -depleted waters ($0.5\text{‰} < \delta^{13}\text{C} < 0.9\text{‰}$) of the antarctic zone. The major drop of $\delta^{13}\text{C}$ values down to 0.5‰ characterizes the antarctic divergence and is due to the upwelling of ^{13}C -depleted waters.

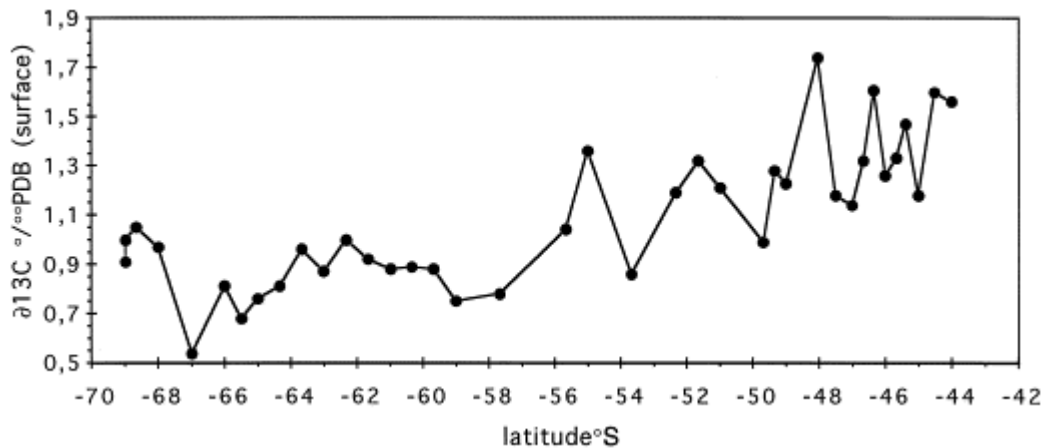


Fig. 7. Latitudinal variations of $\delta^{13}\text{C}$ of ΣCO_2 in the uppermost surface waters of the Southern Ocean.

Along the 30°E meridional section, the $\delta^{13}\text{C}$ values exhibit strong latitudinal and vertical variabilities (Fig. 8). In the subantarctic and polar front zones, the $\delta^{13}\text{C}$ values decrease to $\approx 0.6\text{‰}$ at 1000 m due to organic matter remineralization; the $\delta^{13}\text{C}$ values then remain close to 0.4‰ downward in the water column. The progressive upwelling of ^{13}C -depleted CDW (0.3 to 0.4‰) at the antarctic divergence is responsible for the decrease of about 0.7‰ in surface waters between the subantarctic and the antarctic zone. Around 4000 m, a local maximum of $\delta^{13}\text{C}$ values (0.45 to 0.5‰) is observed and corresponds with a local CFC-12 maximum; identifying a deep lense of better ventilated water.

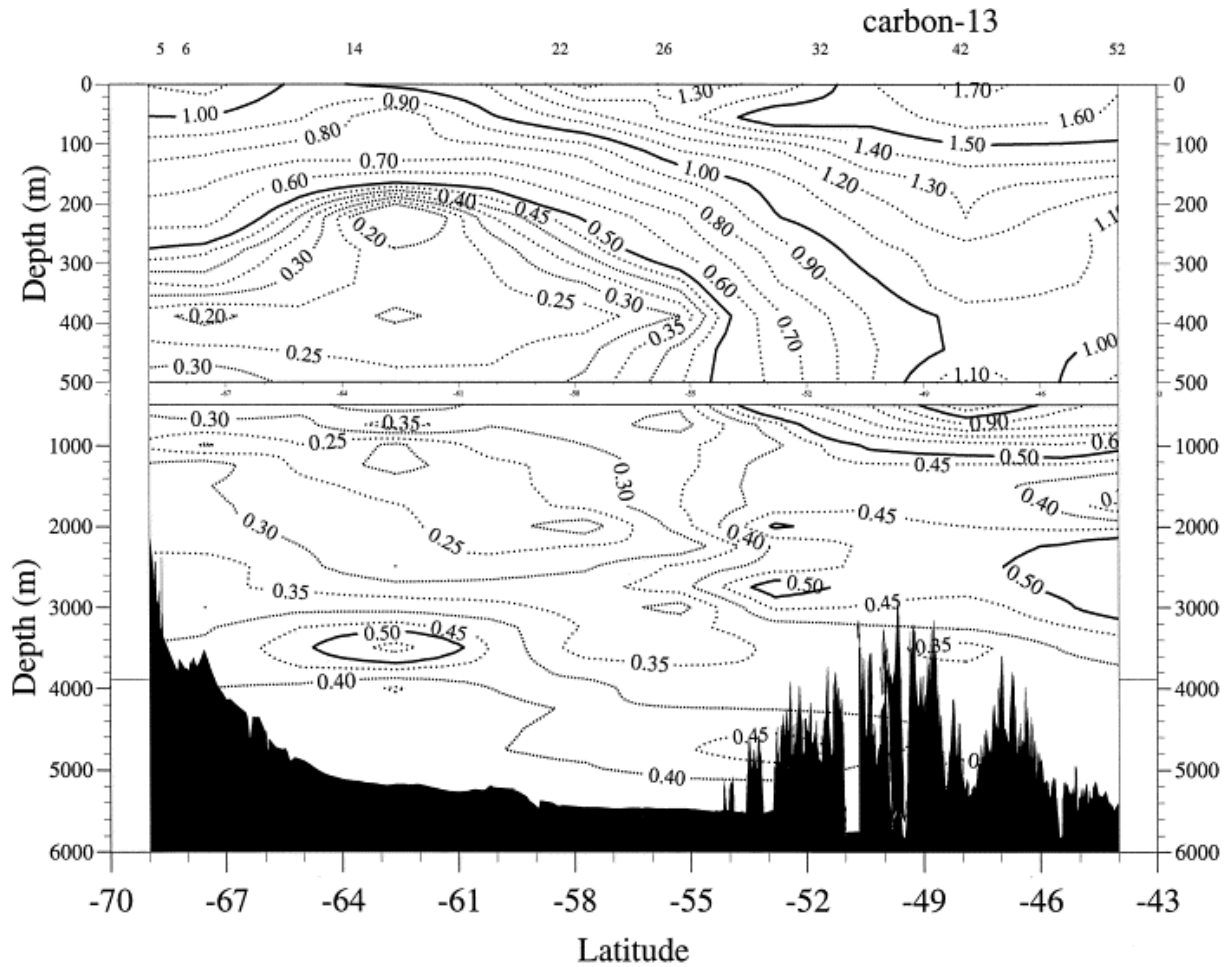


Fig. 8. Latitudinal distribution of the $\delta^{13}\text{C}$ values of ΣCO_2 .

Conclusions

The distributions of $\delta^{13}\text{C}$, $\delta^{18}\text{O}$ and CFC-12 along the meridian 30°E follow the general pathways of the main water masses described by temperature and salinity sections. The AASW and SASW are well discriminated by the uppermost $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ surface values and their mixing occurs in the PF-SAF zone. The sinking of AAIW north of the SAF and the upwelling of CDW at the AD are also well identified by the distribution of these tracers.

These tracers give also new information on the circulation of the water masses in this area. A relative maximum of CFC-12 appears around 4000 m between 66°S and 56°S and is correlated with maximum $\delta^{13}\text{C}$ values; both observations indicate a higher ventilation rate than in the surrounding deep waters. Along the continental margin of Antarctica, the bottom waters, between 2500 and 4000 m, are identified by relatively high $\delta^{18}\text{O}$ values ($-0.2\text{‰} < \delta^{18}\text{O} < 0\text{‰}$) and relatively high CFC-12 concentrations, indicating that this water mass is well ventilated, newly formed and does not come from the downwelling of the superficial waters. This water mass which is not identified by the temperature and salinity parameters may originate either from the Weddell Sea or from Prydz Bay. An extensive study of this region, including physical measurements and geochemical tracer data, would give information on the extension of this newly formed water mass and provide more details on its origin.

Acknowledgements

The CIVA1 cruise took place on board the R/V Marion Dufresne, thanks to the support of the Institut Français des Terres Polaires (IFRTP). We are grateful to all the cruise and to Y. Ballut and B. Ollivier for their valuable assistance. We thank L. Rever and F. Thomas in helping for the CFC measurements and N. Metzl for all the instructive discussions. This work was supported by the WOCE-PNEDC programm.

References

- Bullister, J.L., 1984. Atmospheric Chlorofluoromethans as Tracers of Ocean Circulation and Mixing: Measurement and Calibration Techniques and Studies in the Greenland and Norwegian Seas. PhD Thesis, University of California, San Diego, 172 pp.
- J.L. Bullister and R.F. Weiss, Determination of CCl_3F and CCl_3F_2 in seawater and air, *Deep-Sea Research* 35 (1988) (5), pp. 839–853.
- Carmack, E.C., 1977. Water characteristics of the Southern Ocean south of the Polar Front. In: Angel, M. (Ed.), *A Voyage of Discovery*. Pergamon, pp. 15–41.
- H. Craig, Isotopic standards for carbon and oxygen and correction factors for mass-spectrometric analysis of carbon dioxide, *Geochimica et Cosmochimica Acta* 12 (1957), pp. 133–149.
- H. Craig, Standards for reporting concentrations of deuterium and oxygen-18 in natural waters, *Science* 133 (1961), pp. 1833–1834.
- Craig, H., Gordon, L.I., 1965. Deuterium and oxygen 18 variations in the ocean and the marine atmosphere. In: Tongiorgi, E. (Ed.), *Stable Isotopes in Oceanographic Studies and Paleotemperatures*. Spoleto, Italy, CNR, pp. 9–130.
- G.E.R. Deacon, The Weddell gyre, *Deep-Sea Research A* 26 (1979), pp. 981–995.
- S. Epstein and T.K. Mayeda, Variations of the $^{18}\text{O}/^{16}\text{O}$ ratio in natural waters, *Geochimica et Cosmochimica Acta* 4 (1953), pp. 213–224.
- T.D. Foster and E.C. Carmack, Frontal zone mixing and antarctic bottom water formation in the Southern Weddell Sea, *Deep-Sea Res.* 23 (1976), pp. 301–317.
- Gordon, A.L., 1971. Oceanography of antarctic waters. In: Reid, J.L. (Ed.), *Antarctic Oceanography 1: Antarctic Research Series*, Vol. 15. American Geophysical Union, 169–203.
- V.V. Gouretski and A.I. Danilov, Weddell gyre: structure of the eastern boundary, *Deep-Sea Research* 40 (1993) (3), pp. 561–582.
- Jacobs, S.S., Georgi, D.T., 1977. Observations on the southwest Indian/Antarctic ocean. In: Angel, M. (Ed.), *A Voyage of Discovery*. Pergamon, pp. 43–84.
- P. Jean-Baptiste, F. Mantsi, L. Mémery and D. Jamous, ^3He and chlorofluorocarbons (CFC) in the Southern Ocean: tracers of water masses, *Marine Chemistry* 35 (1991), pp. 137–150.
- P.M. Kroopnick, The dissolved $\text{O}_2\text{--CO}_2\text{--}^{13}\text{C}$ system in the eastern equatorial Pacific, *Deep-Sea Research* 21 (1974), pp. 211–227.

P.M. Kroopnick, The distribution of ^{13}C of ΣCO_2 in the world oceans, *Deep-Sea Research* 32 (1985) (1), pp. 57–84.

J.R.E. Lutjeharms and H.R. Valentine, Southern Ocean thermal fronts south of Africa, *Deep-Sea Research* 31 (1984) (12), pp. 1461–1475.

F. Mantsi, C. Beauverger, A. Poisson and N. Metzl, Chlorofluoromethans in the western Indian sector of the Southern Ocean and their relations with geochemical tracers, *Marine Chemistry* 35 (1991), pp. 151–167.

A.H. Orsi, W.D. Nowlin Jr. and T. Whitworth III, On the circulation and stratification of the Weddell Gyre, *Deep-Sea Res.* 40 (1993) (1), pp. 169–203.

Y.H. Park, L. Gamberoni and E. Charriaud, Frontal structure, water masses, and circulation in the Crozet Basin, *J. Geophys. Res.* 98 (1993) (C7), pp. 12361–12385.

P. Schlosser, R. Bayer, A. Foldvik, T. Gammelsrød, G. Rohardt and K.O. Münnich, Oxygen-18 and helium as tracers of Ice Shelf Water and water/ice interaction in the Weddell Sea, *J. Geophys. Res.* 95 (1990) (C3), pp. 3253–3263.

P. Schlosser, J.L. Bullister J. and R. Bayer, Studies of deep water formation and circulation in the Weddell Sea using natural and anthropogenic tracers, *Mar. Chem.* 35 (1991), pp. 97–122.

N.R. Smith, D. Zhaoqian, K.R. Kerry and S. Wright, Water masses and circulation in the region of Prydz Bay, Antarctica, *Deep-Sea Res.* 31 (1984) (9), pp. 1121–1147.

R.F. Weiss, H.G. Östlund and H. Craig, Geochemical studies of the Weddell Sea, *Deep-Sea Res. A* 26 (1979), pp. 1093–1120.

T. Whitworth III and W.D. Nowlin Jr., Water masses and currents of the Southern Ocean at the Greenwich Meridian, *J. Geophys. Res.* 92 (1987) (C6), pp. 6462–6476.